

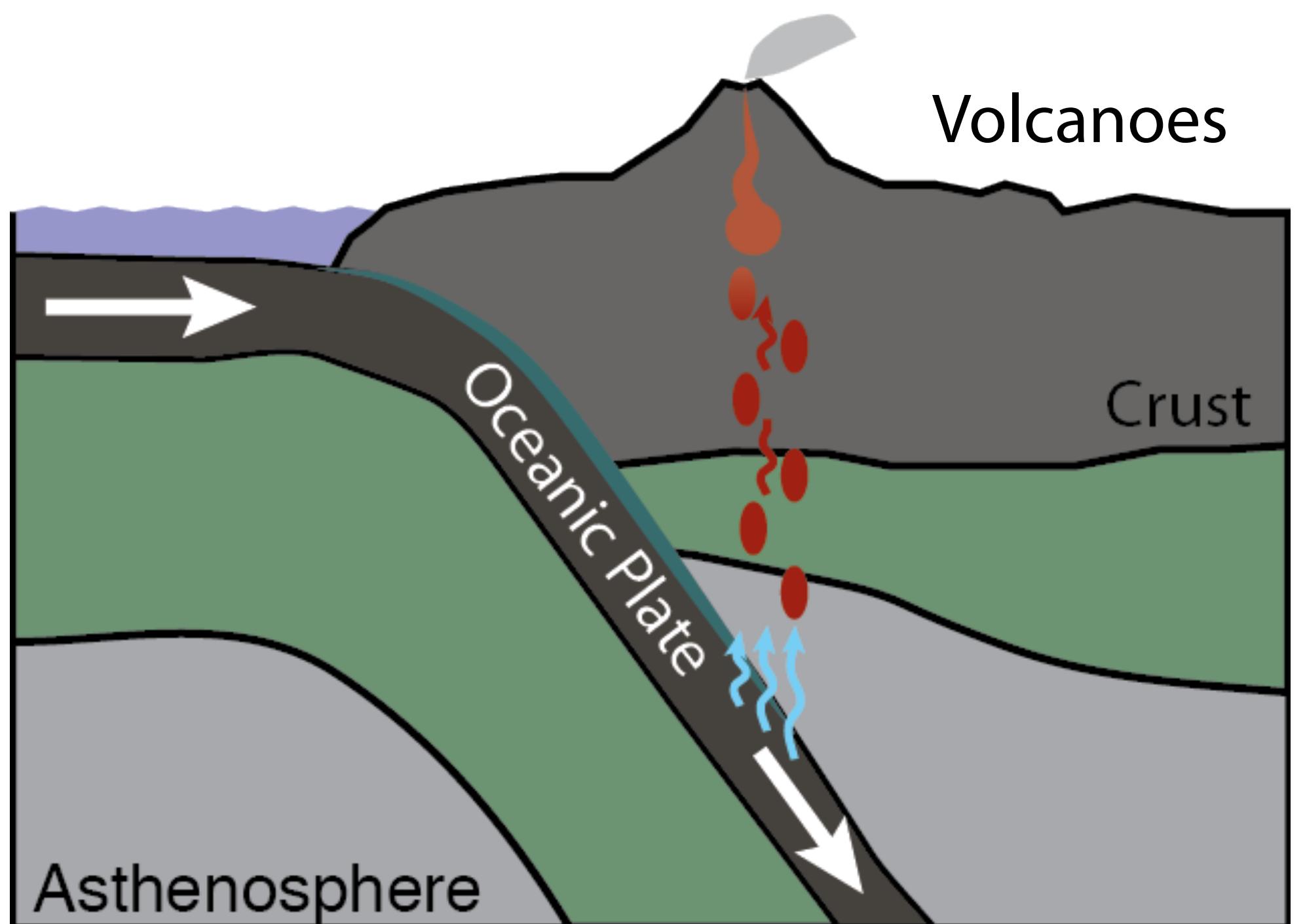
Fluid mediated deformation controls slip behaviors at the base of the subduction seismogenic zone

Cailey B. Condit, Eirini Poulaki, Jason Ott, Peter C. Lindquist, Will F. Hoover, Tshering Lama Sherpa

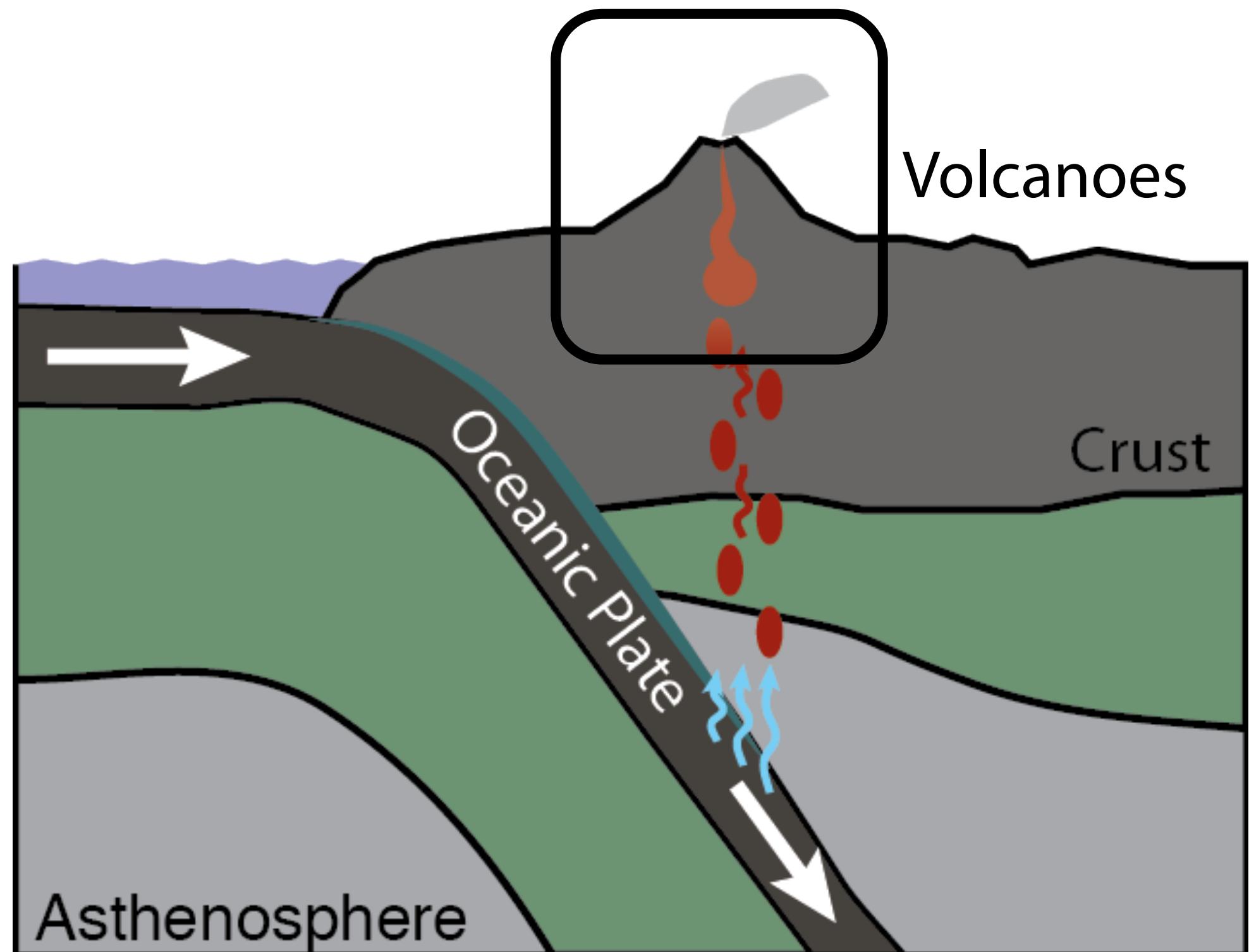


Research funding provided by: UW Royalty Research Fund; **National Science Foundation** CAREER Award to Condit (NSF EAR-2338181), NSF-EAR Award 221781 to Condit, and Postdoctoral Research Fellowship to Hoover at UW (NSF EAR-2053033)

Subduction zones

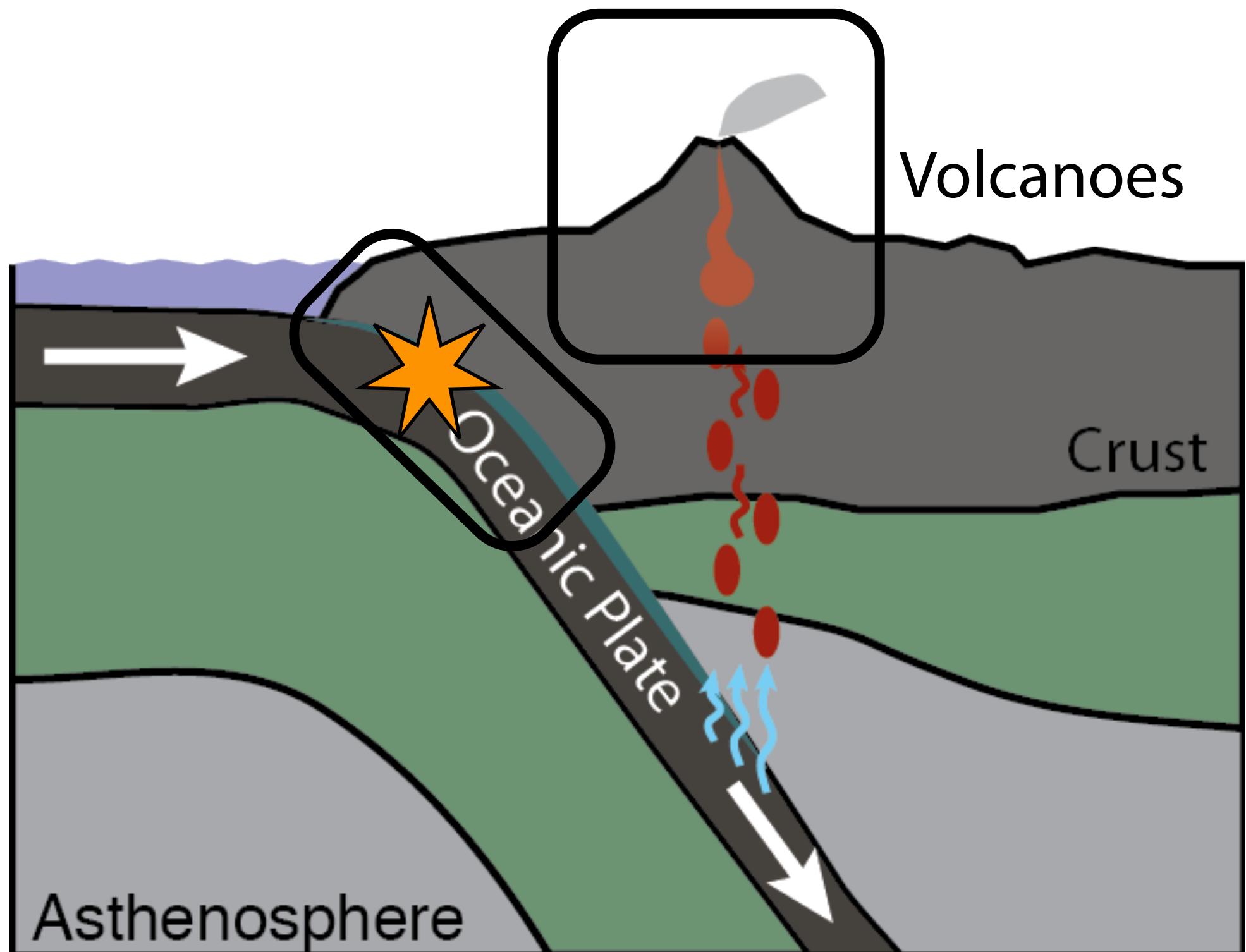


Subduction zones



Most devastating volcanic eruptions

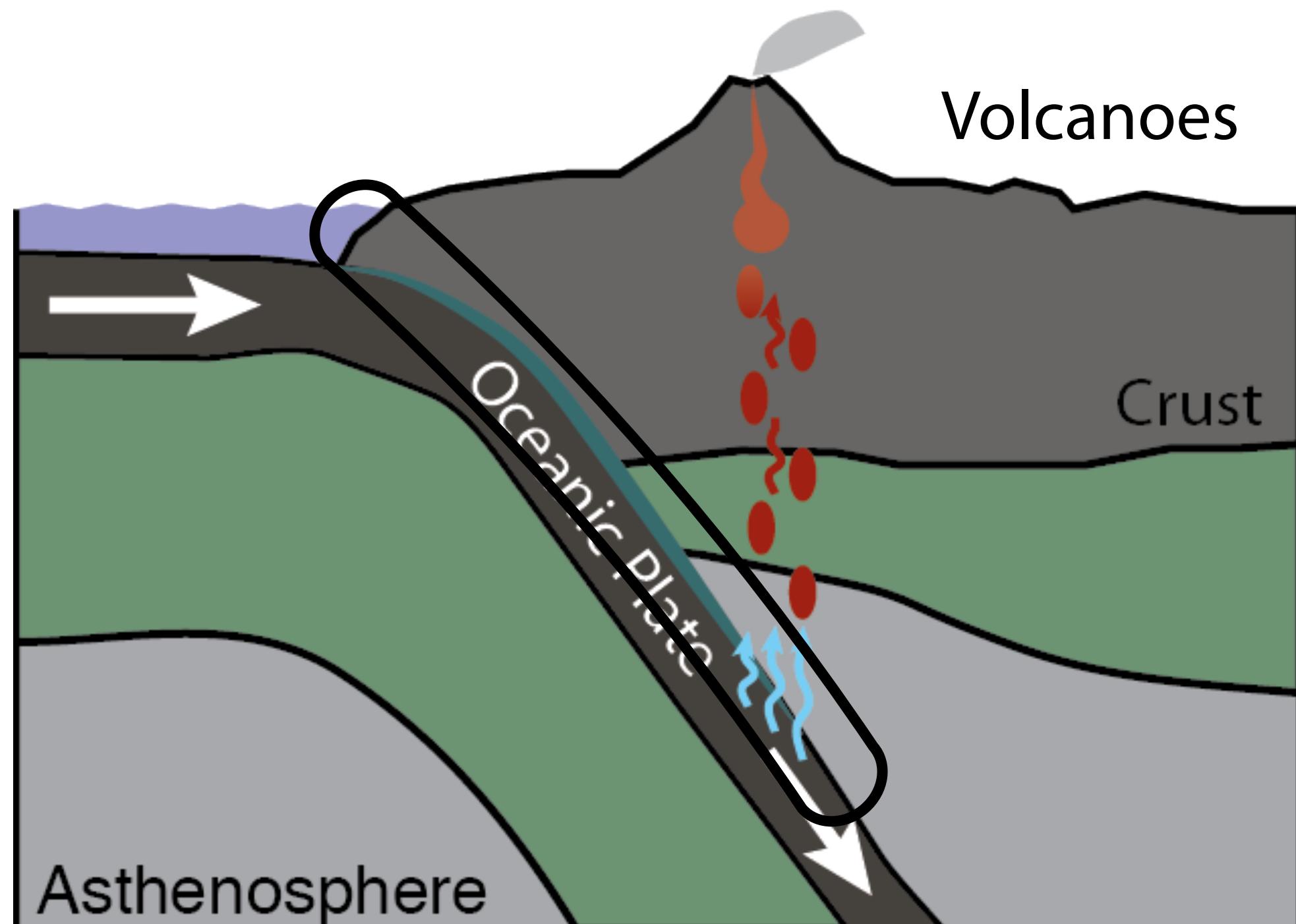
Subduction zones



Most devastating volcanic eruptions

Largest earthquakes on earth

Subduction zones: Plate boundary processes



Most devastating volcanic eruptions

Largest earthquakes on earth

Plate interface locus of *fluids*, chemical & mechanical transformations

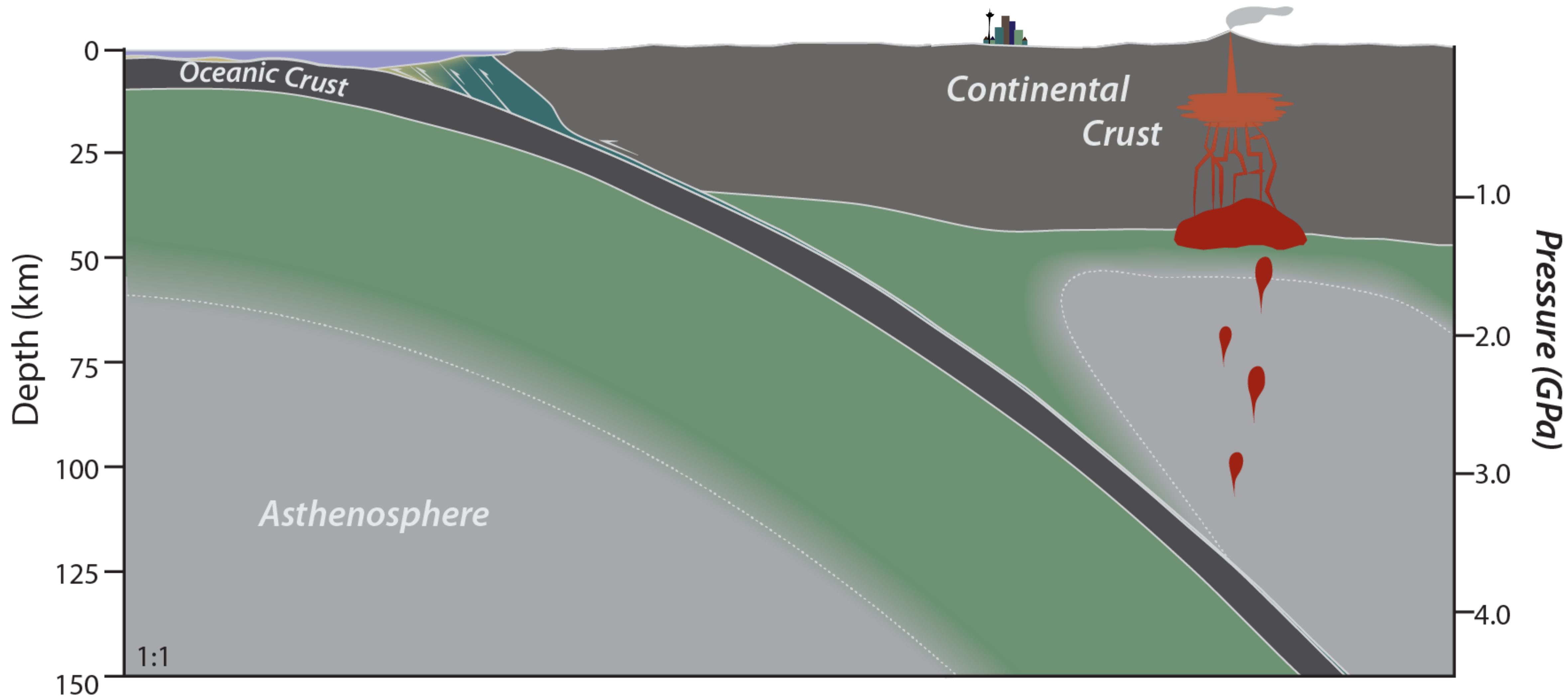


Plate interface locus of *fluids*, chemical & mechanical transformations

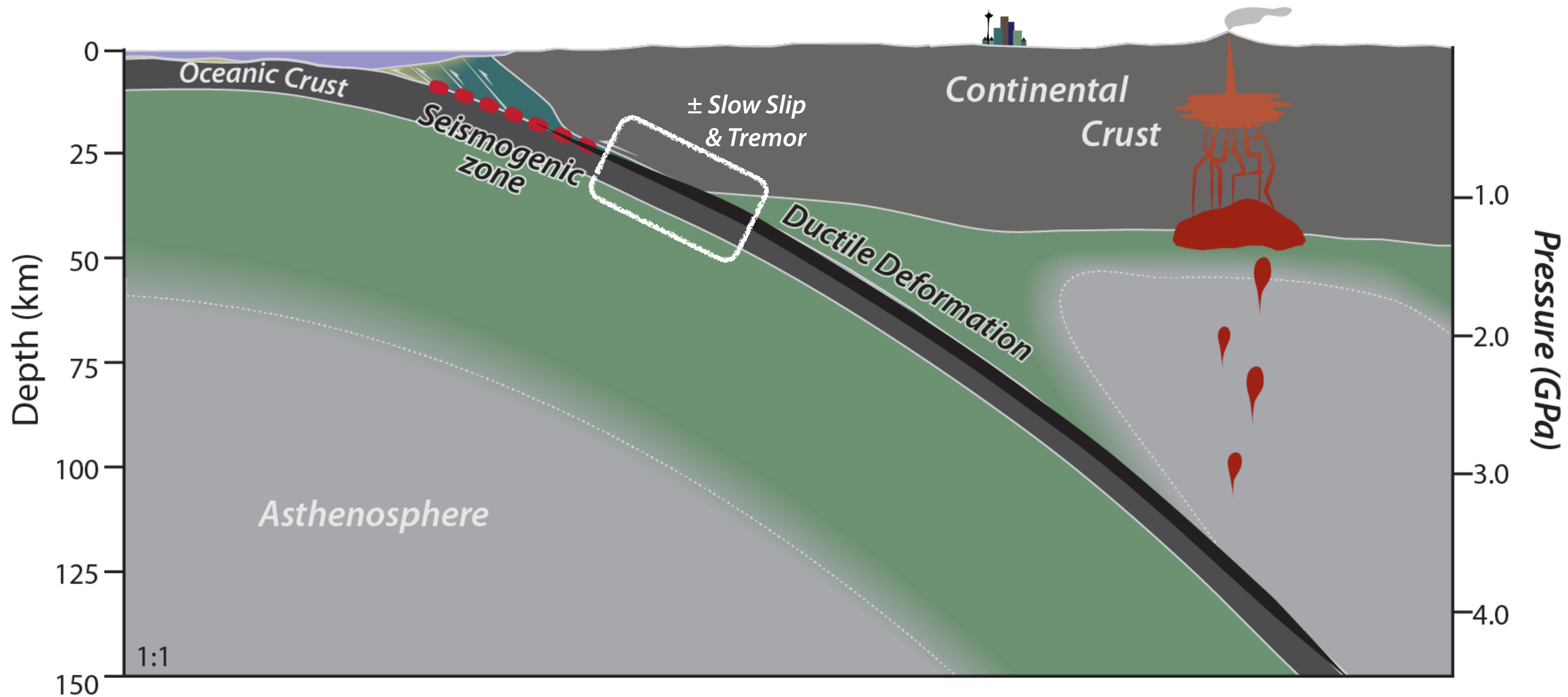
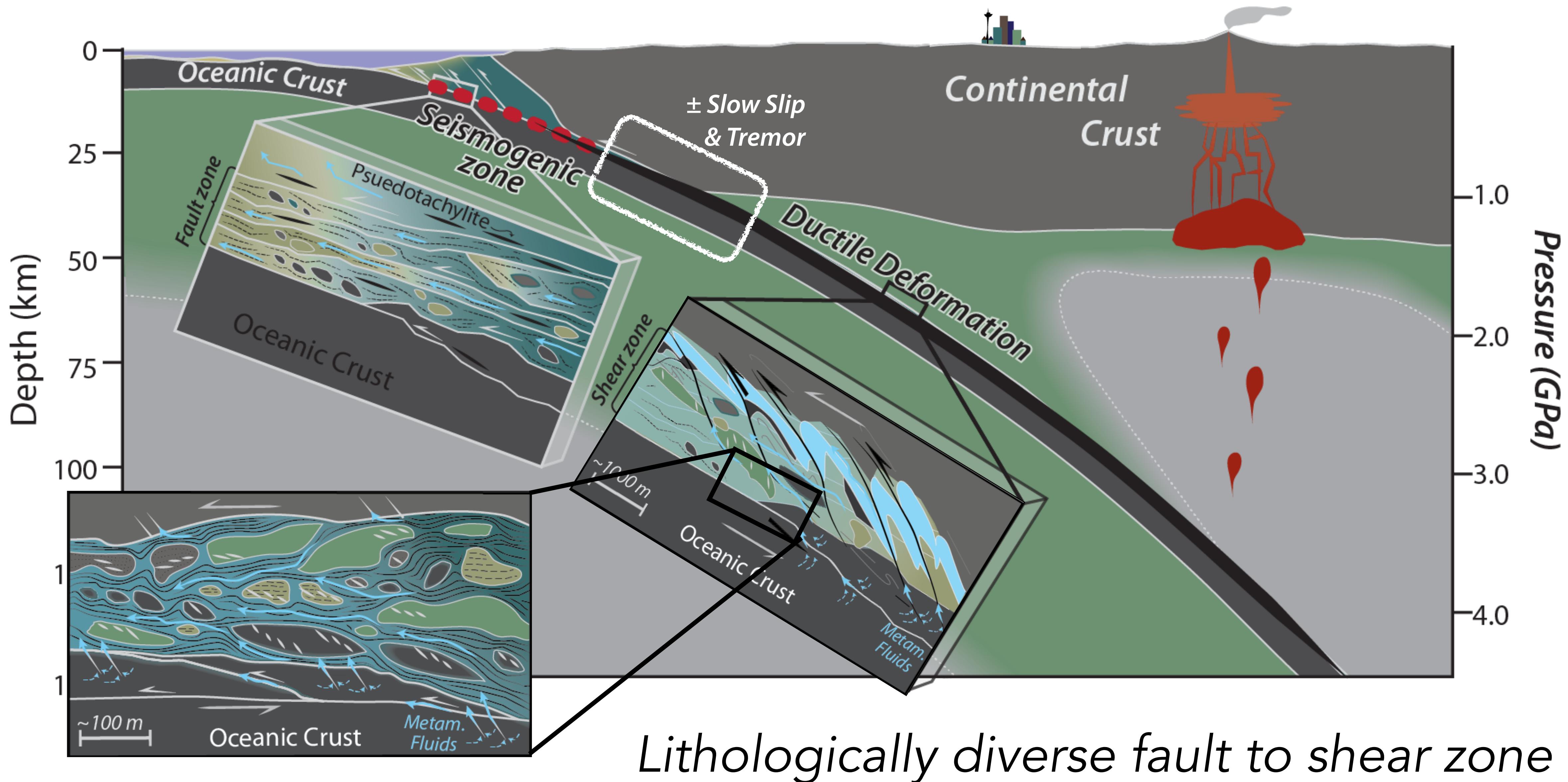
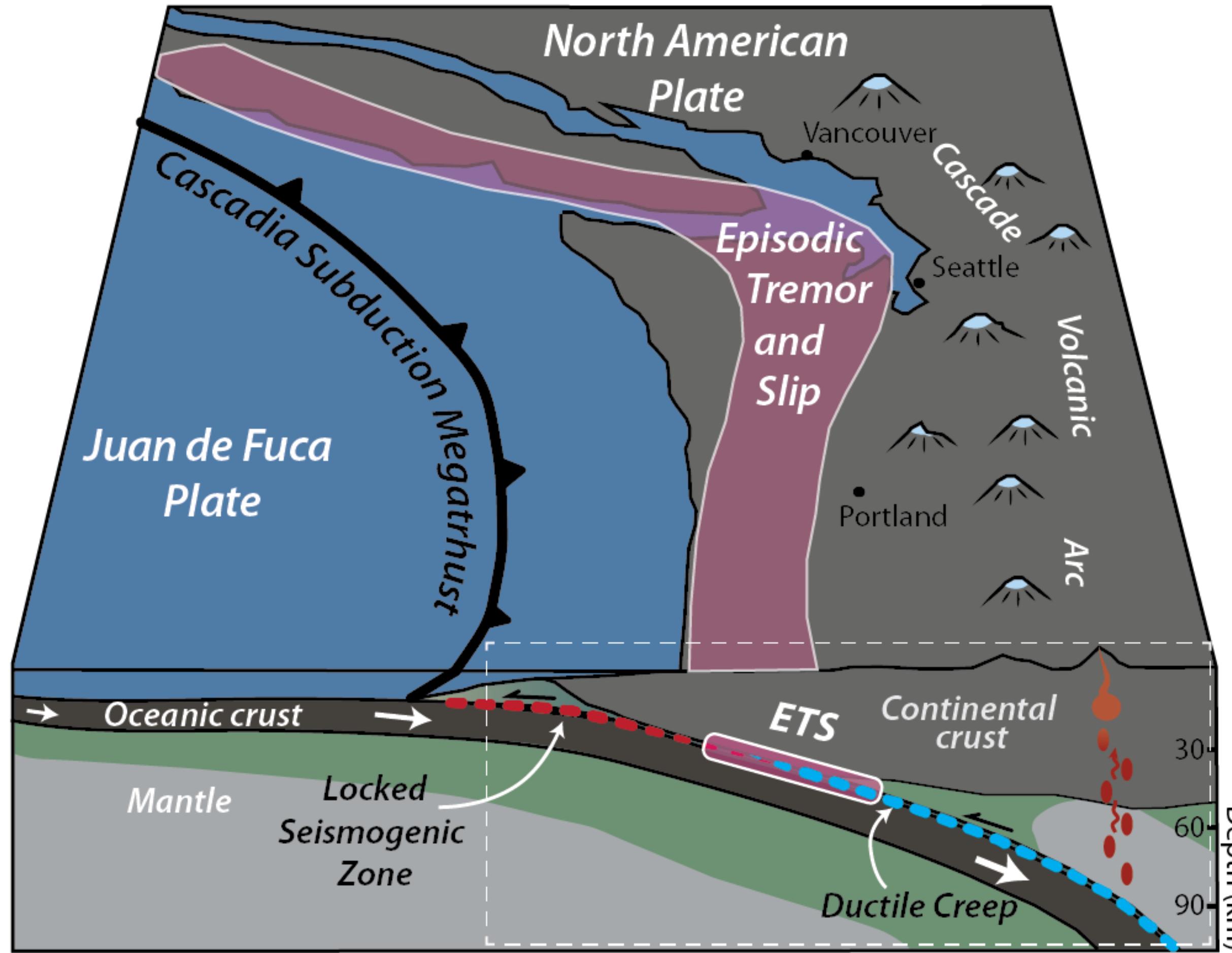


Plate interface locus of *fluids*, chemical & mechanical transformations

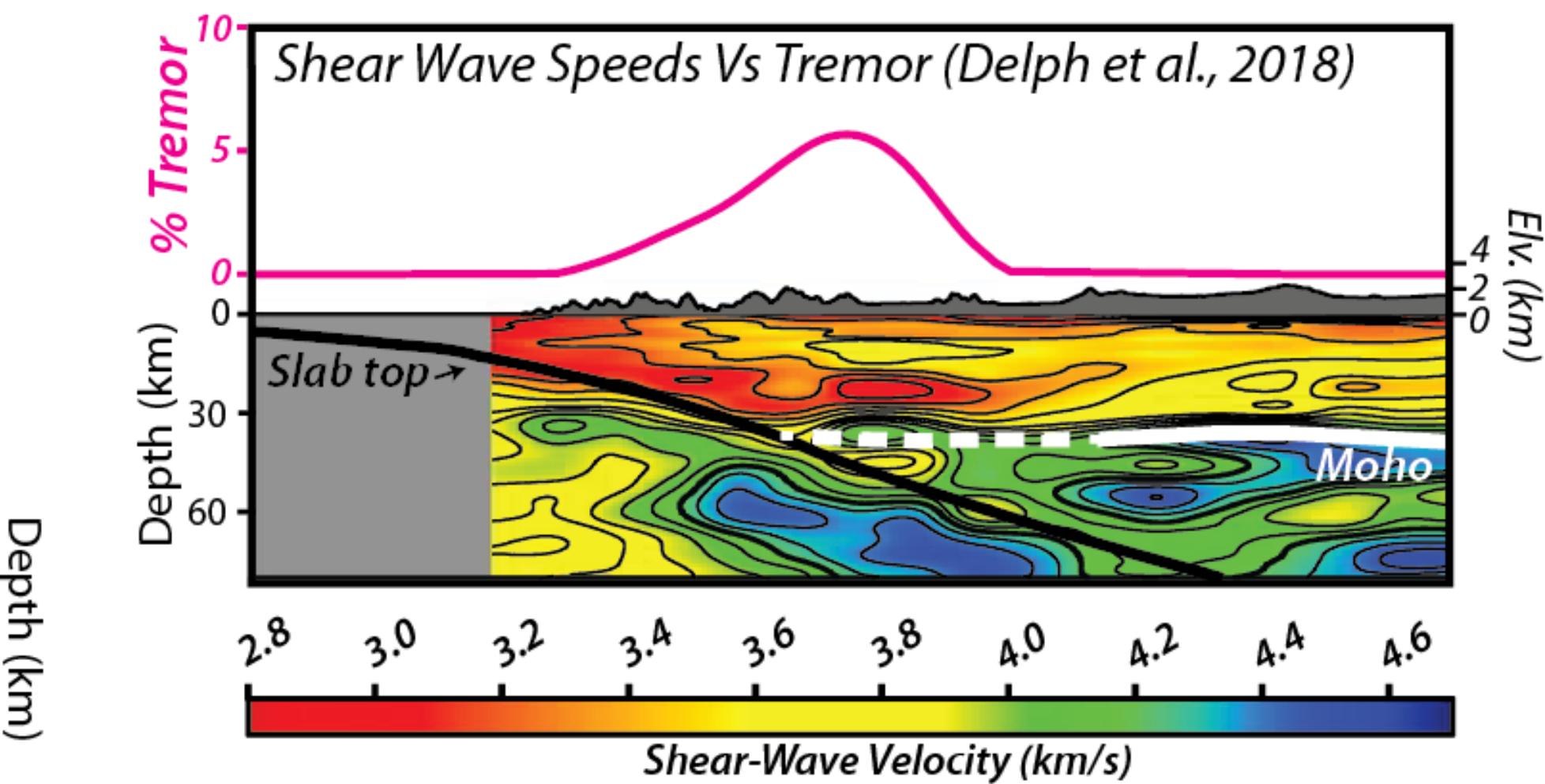
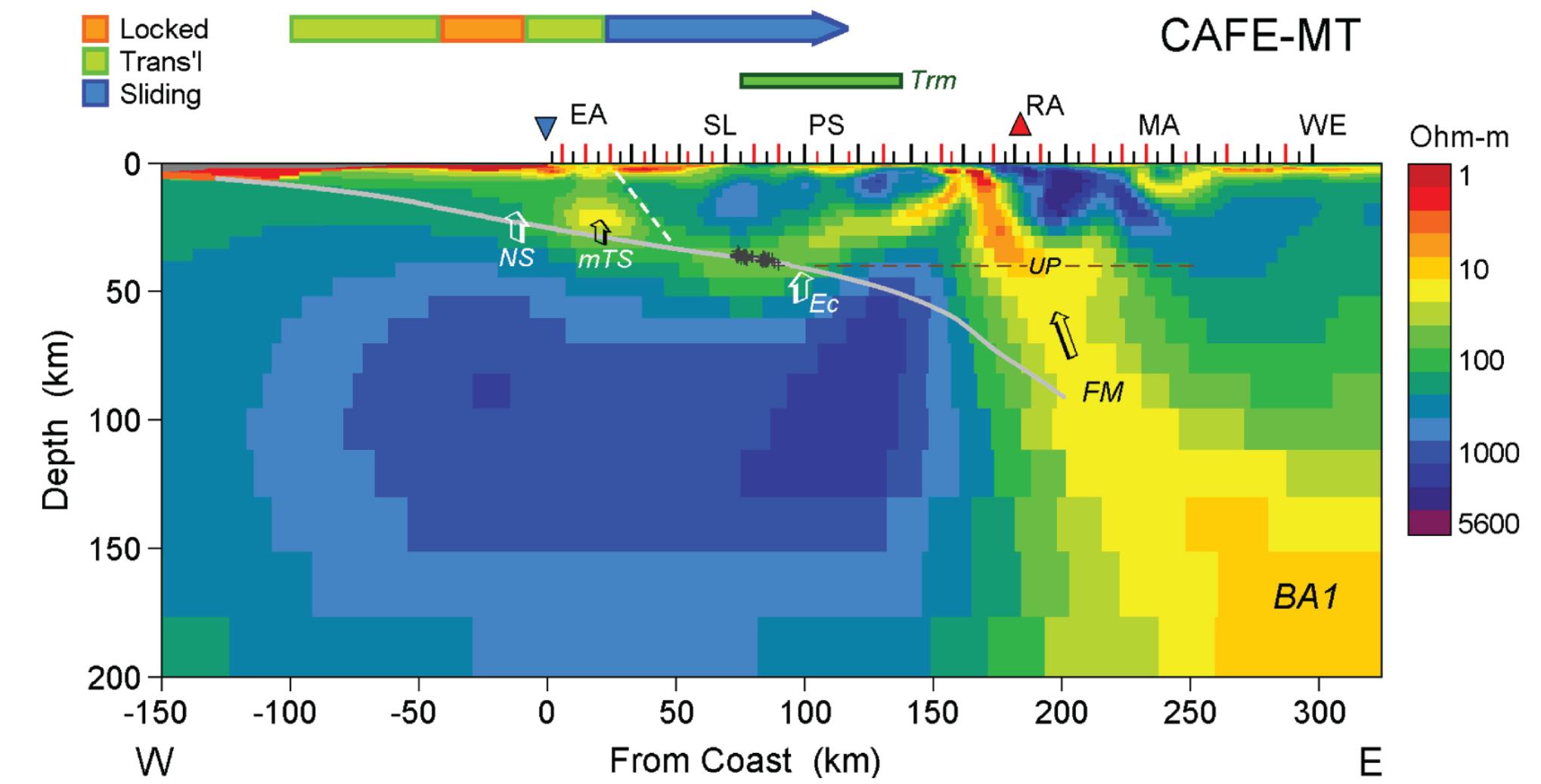


Base of the subduction seismogenic zone: EQs and slow slip events

Geophysical observations

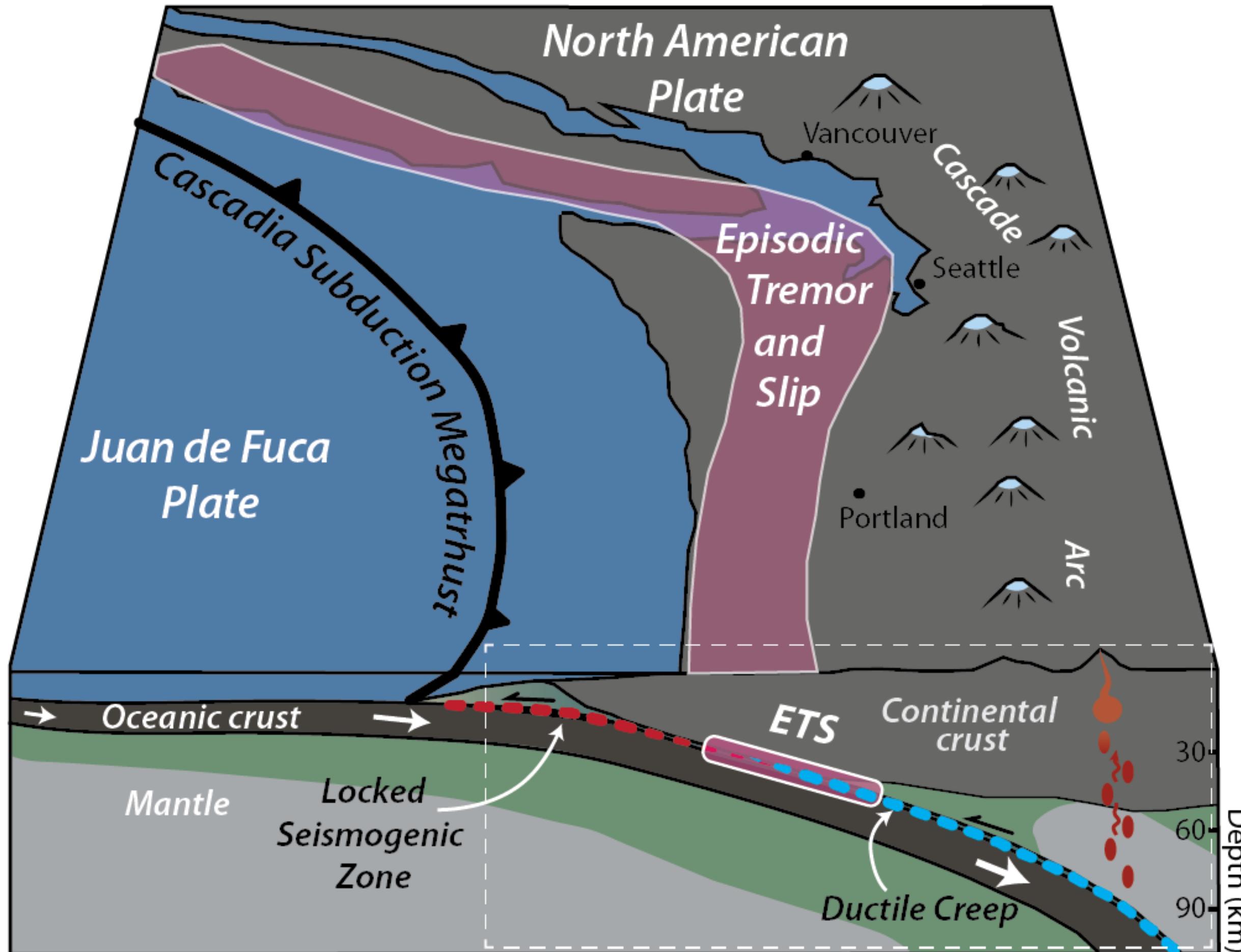


MT images conductors, inferred fluids (Wanamaker et al., 2014)



Base of the subduction seismogenic zone: EQs and slow slip events

Where deep slow slip occurs...



- Geophysical observations → **Fluid-rich** environment
 - slow V_s , conductive zones in MT, high V_p/V_s ratios
- Tidal trigger of slow slip events
 - very small stress perturbations (kPa) result in several order magnitude increase in slip rates
→ most consistent with low differential stress
- Inferred **high pore fluid pressures (P_f)**

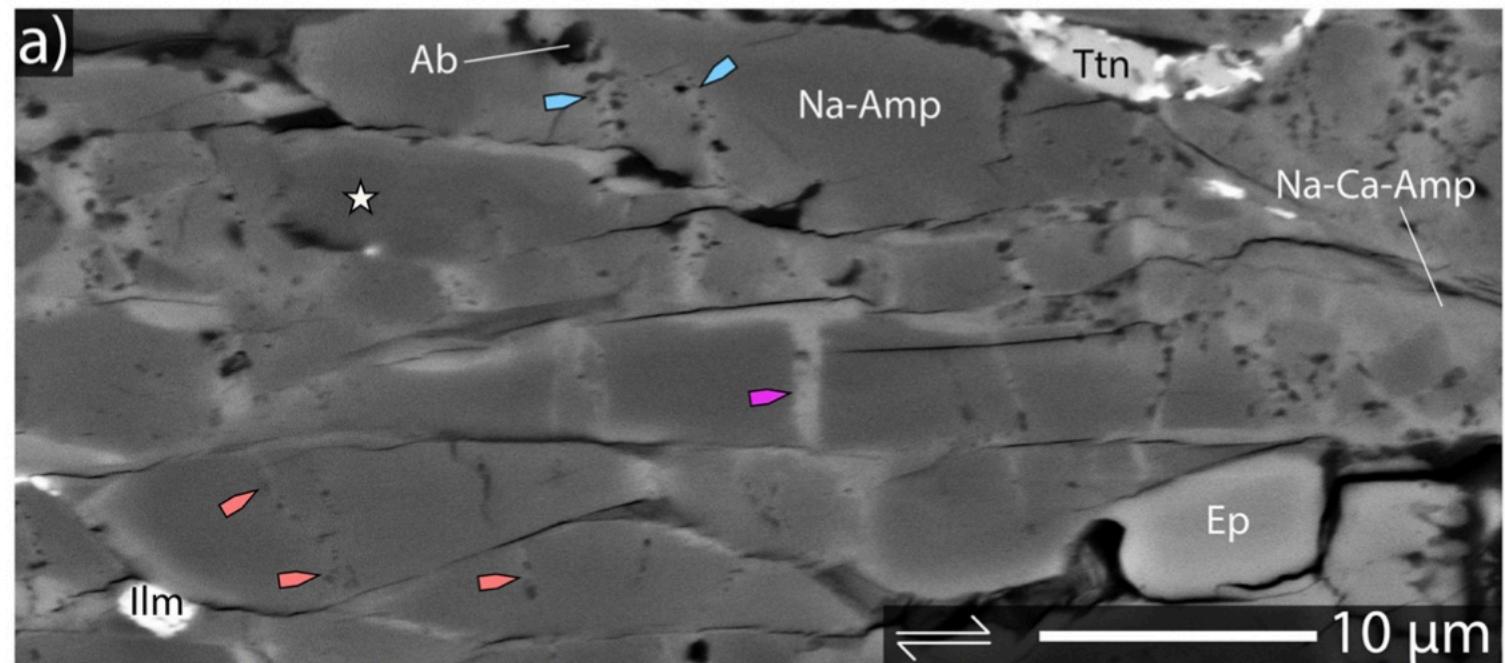
Deep slow slip events have preceded large megathrust earthquakes; contribute to the subduction zone slip budget; Occur down-dip* of megathrust hypocenters but in their slip patches

Strength of the plate interface dependent on **fluid behavior and volumes**

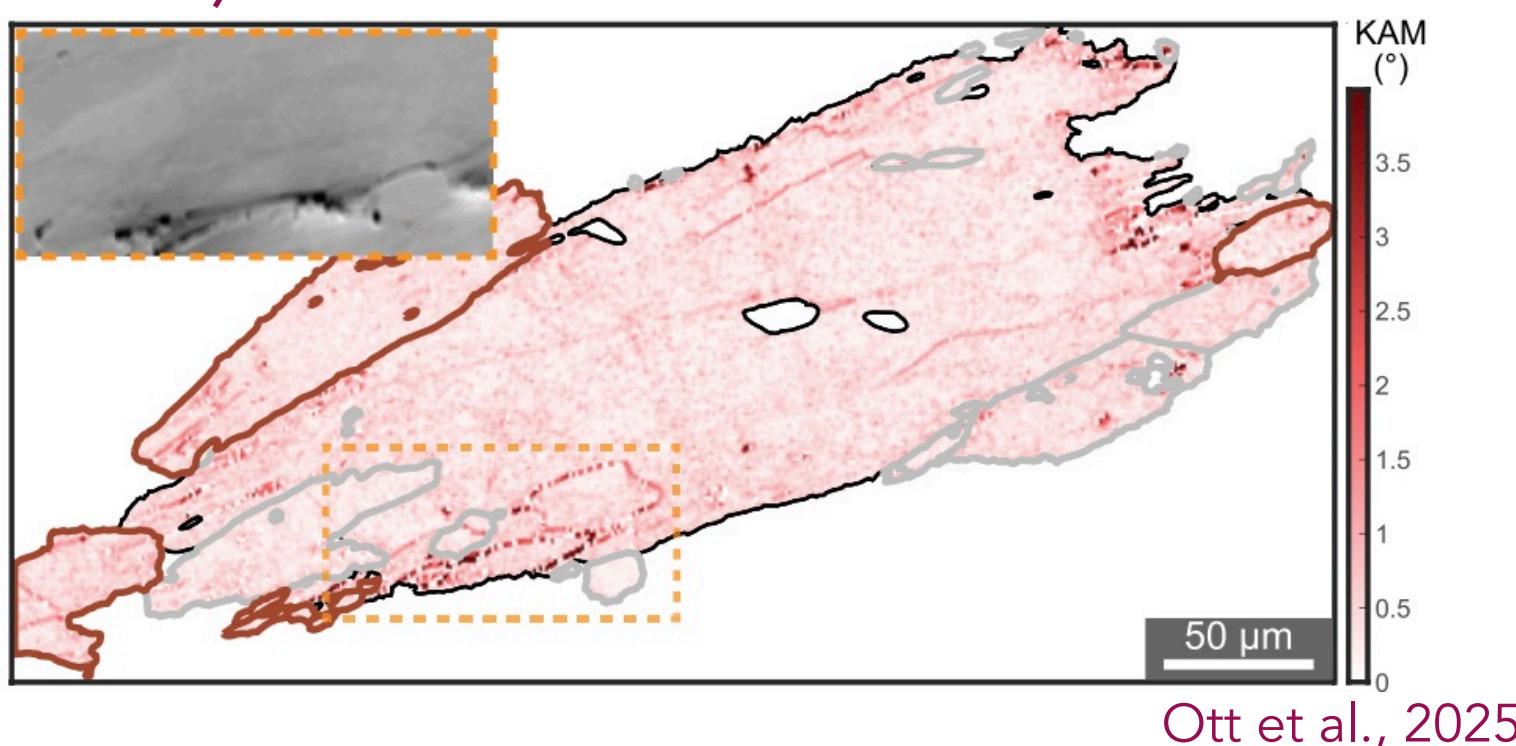
Fluids can change def. mechanism, mineralogy, and mode of deformation

Shift in mechanism

Microboudinage (diffusive mechanism; $n \sim 2$)



Dislocation creep (power law rheology;
 $n \sim 3-5$)

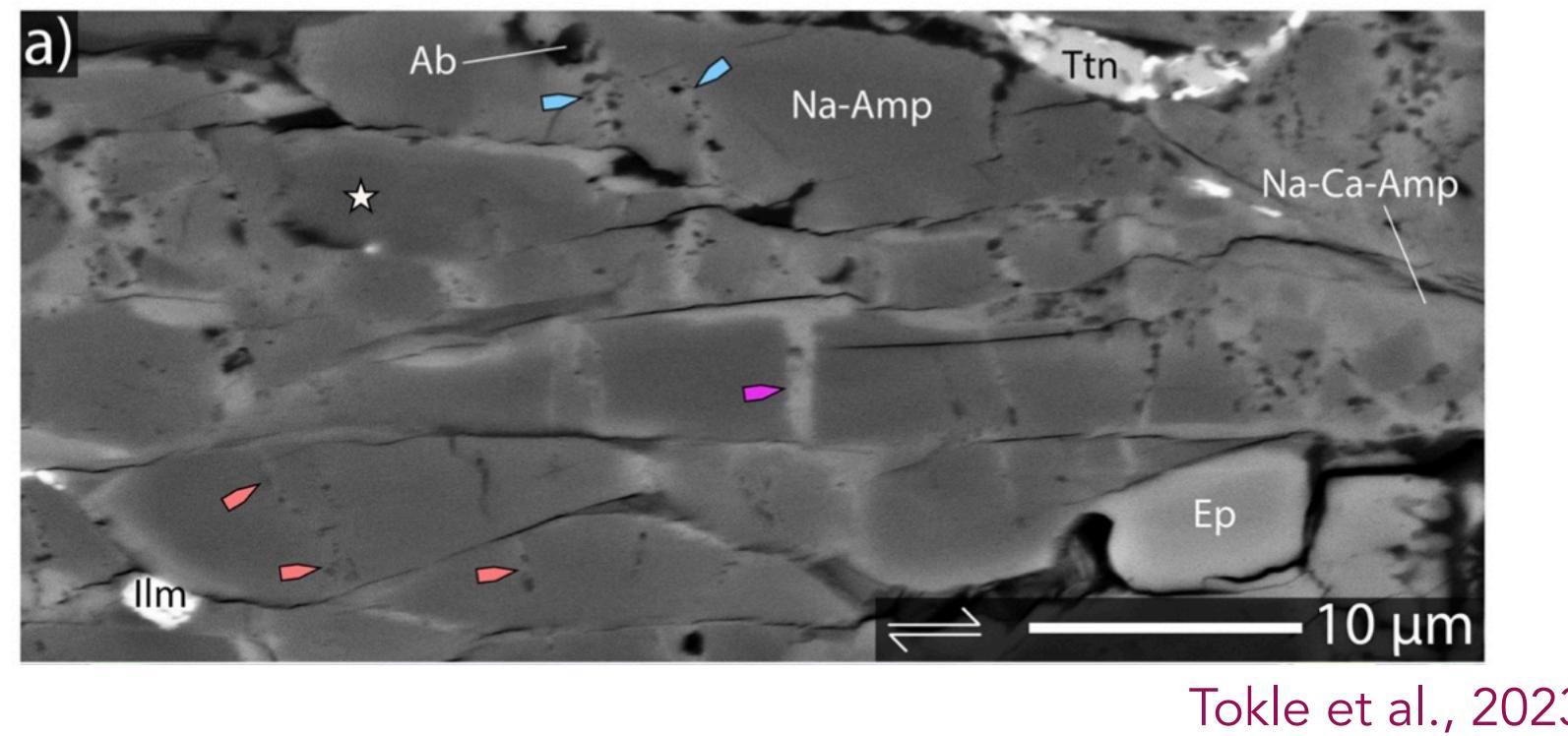


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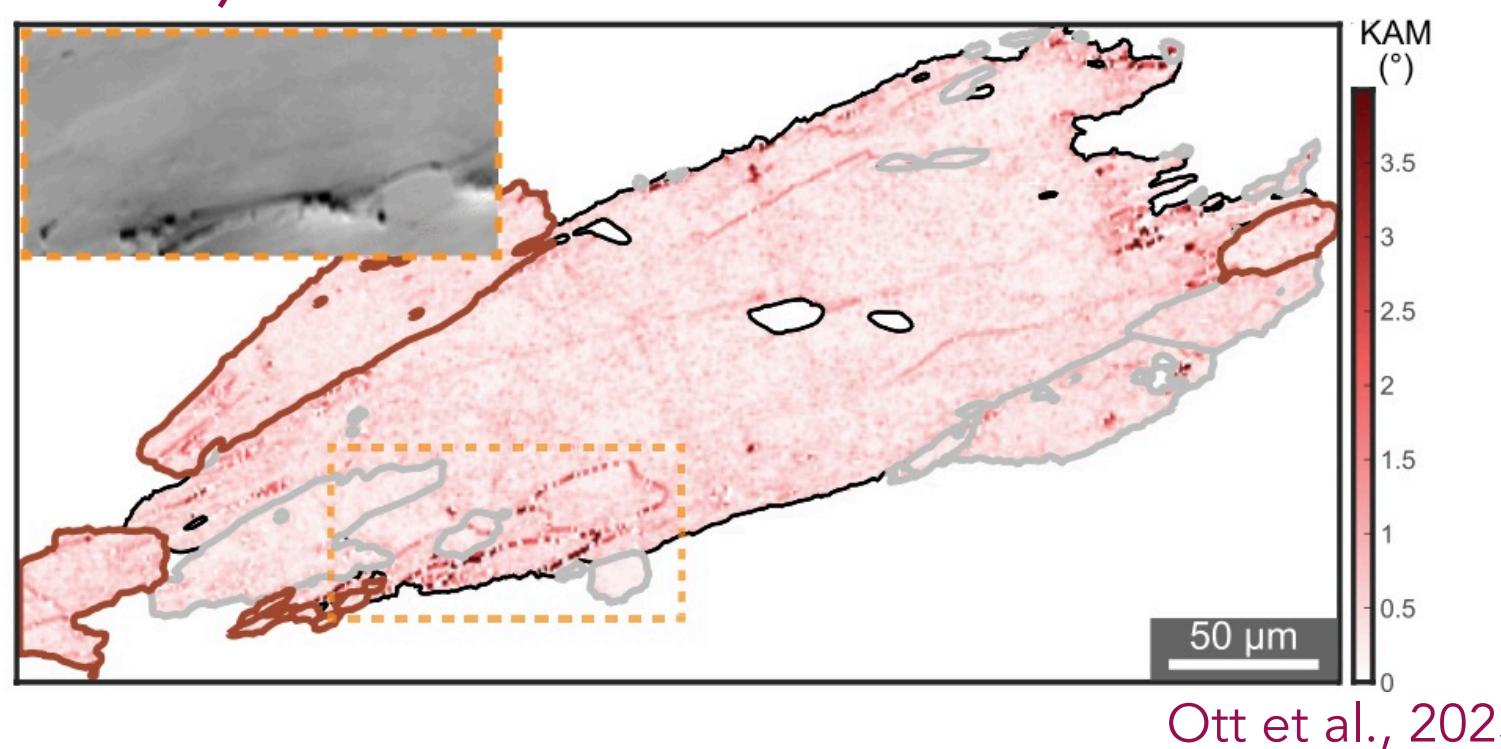
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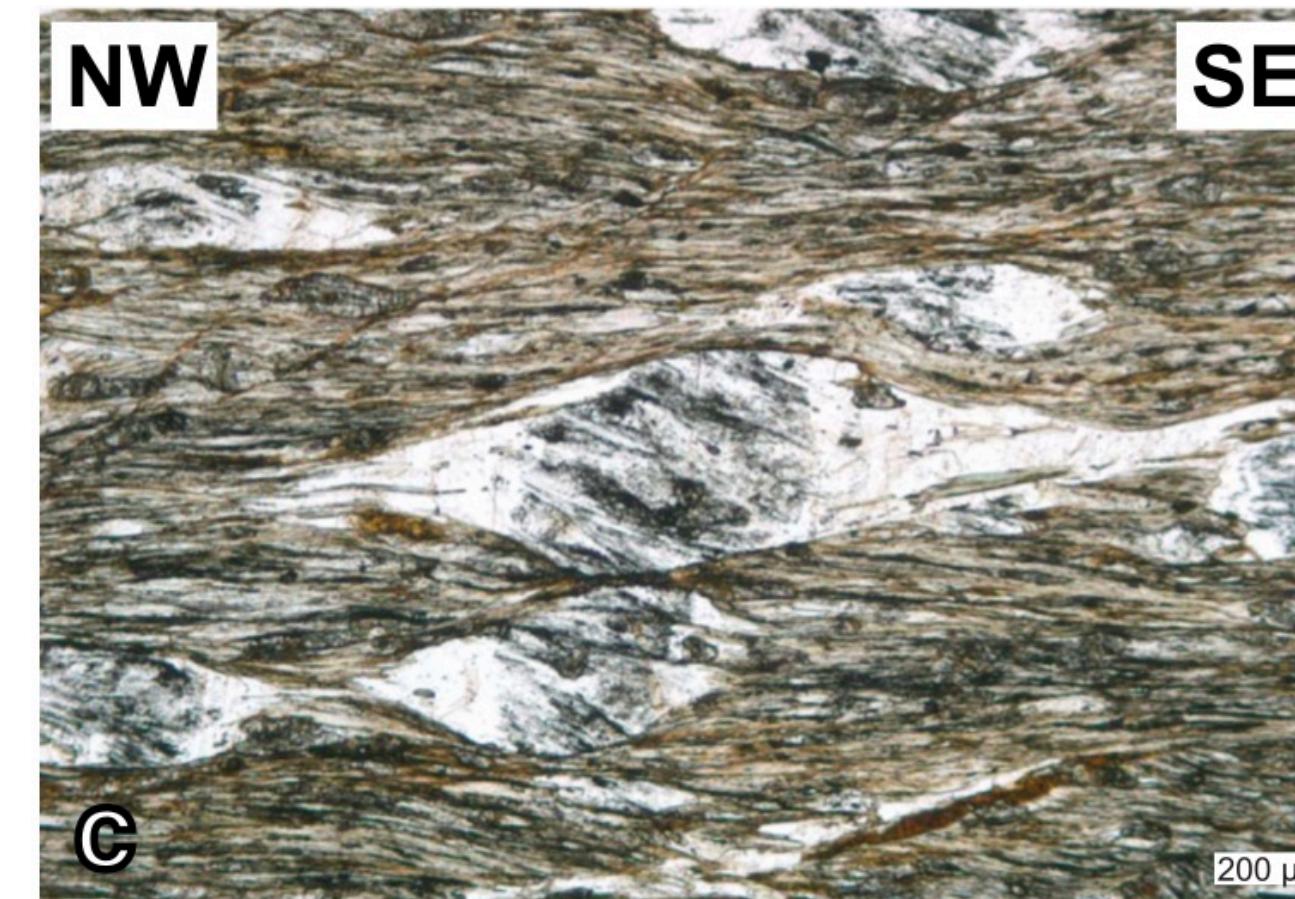
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Dislocation creep (power law rheology; $n \sim 3-5$)



New minerals through fluid-mediated metamorphic reactions



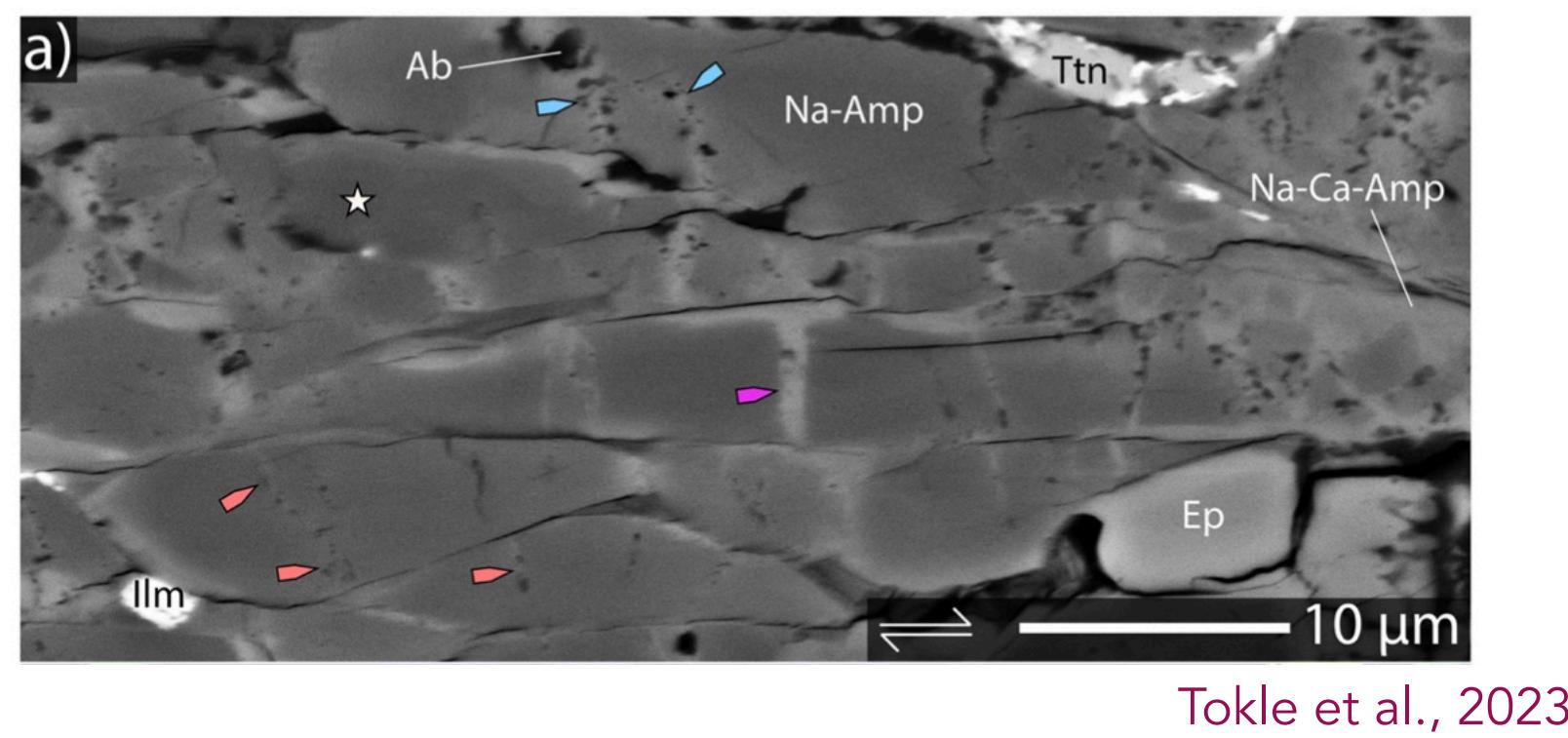
Transformation-Assisted Creep, Ab porphyroblast/clasts reacting to quartz+ w/ fluid in Pelona schist

Strength of the plate interface dependent on **fluid behavior and volumes**

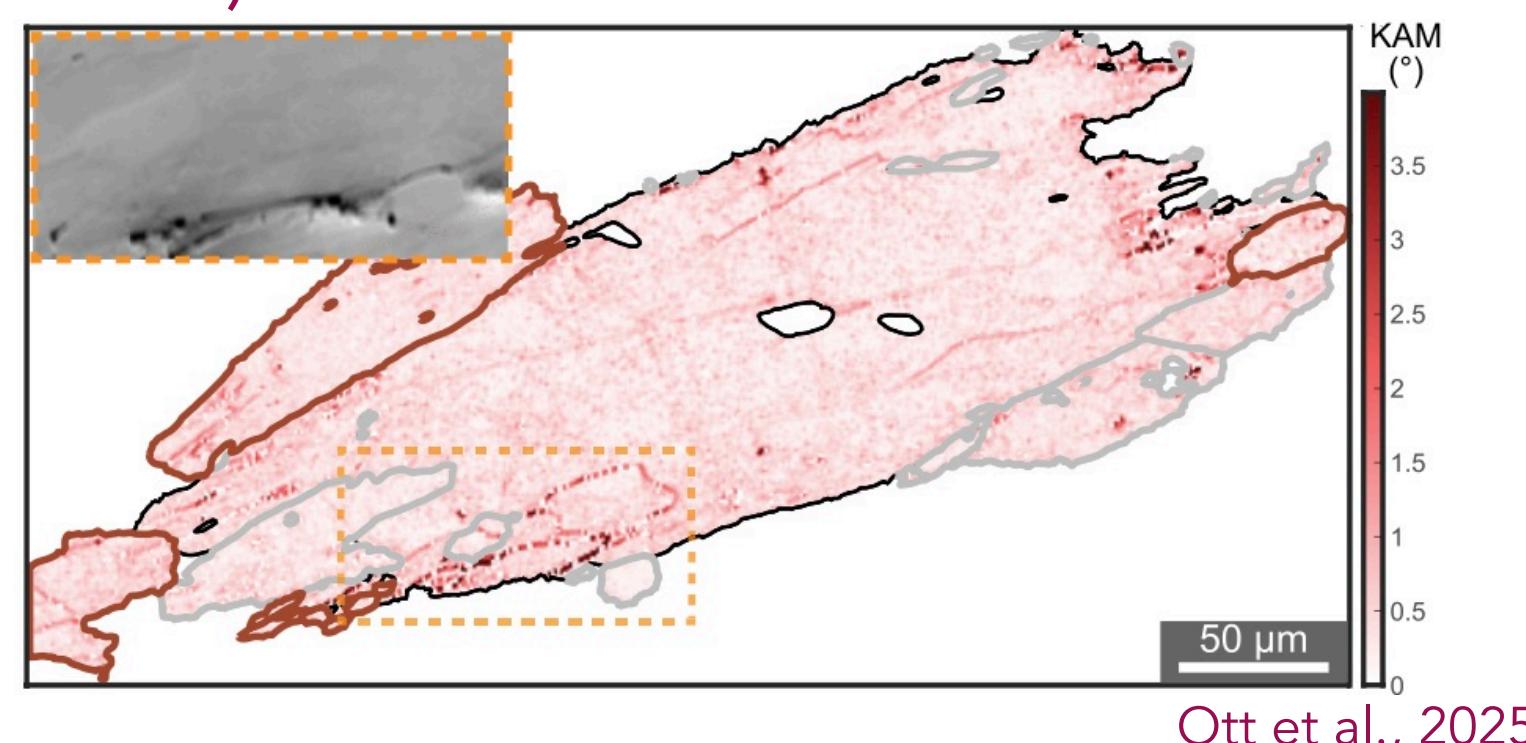
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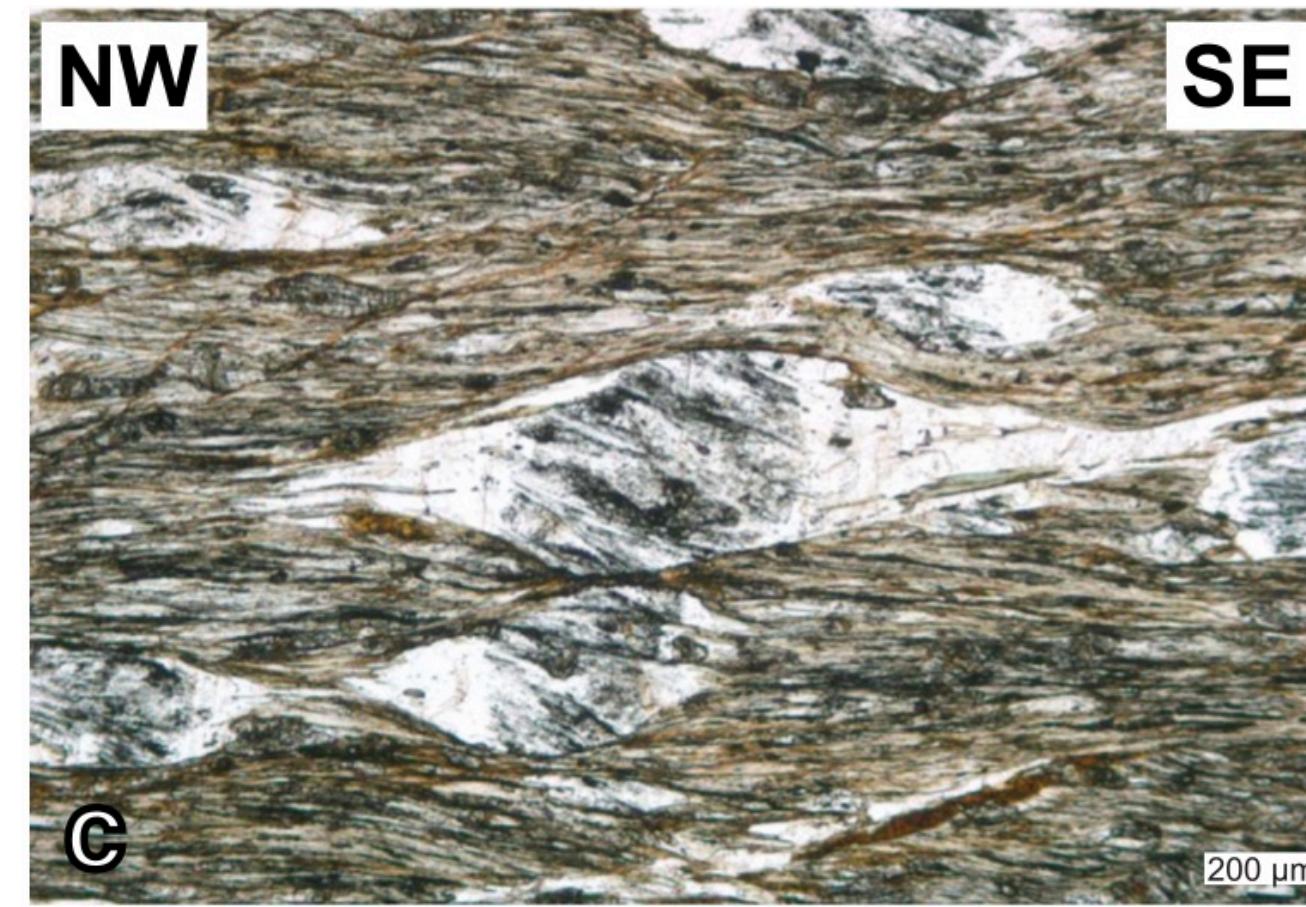
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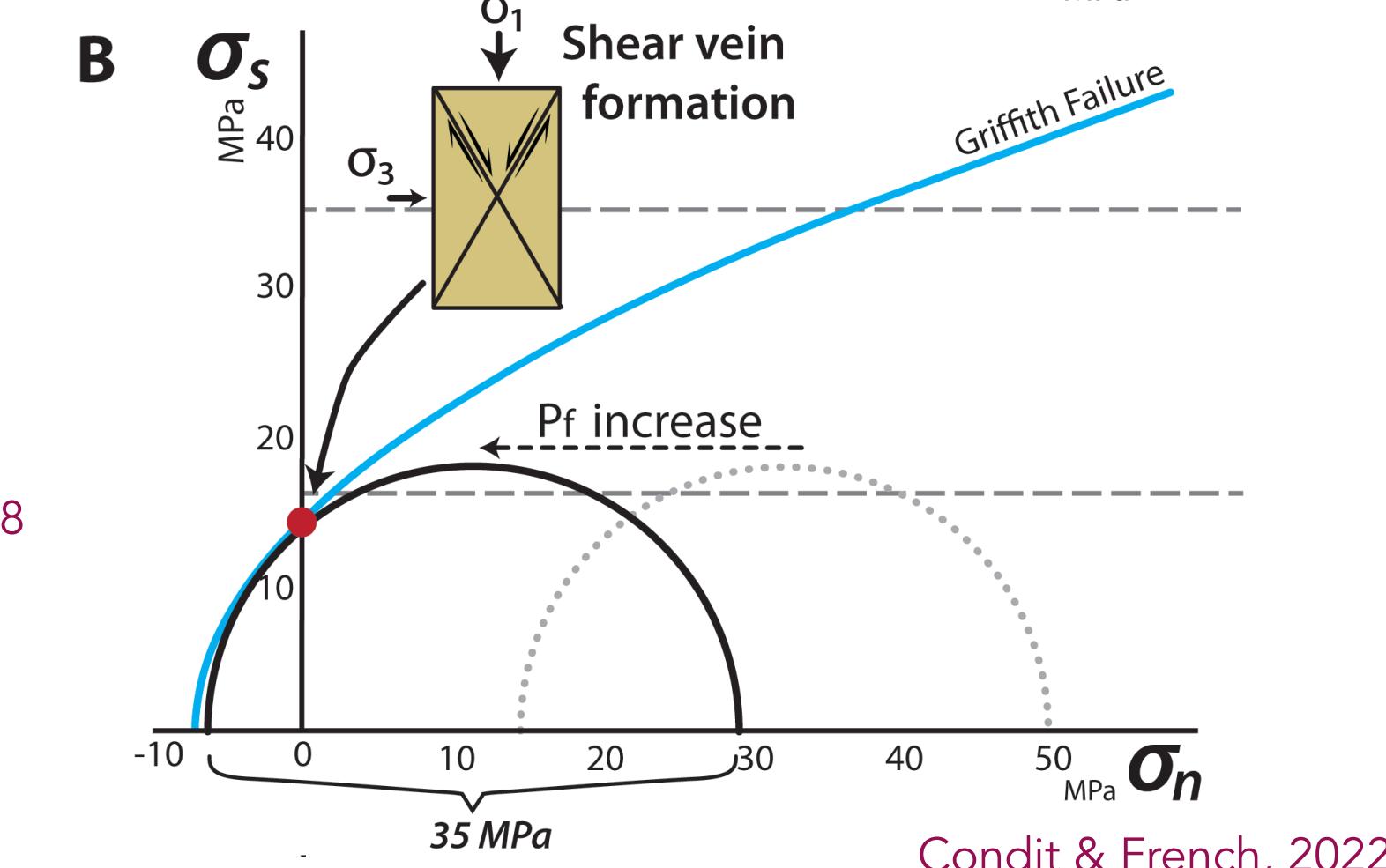
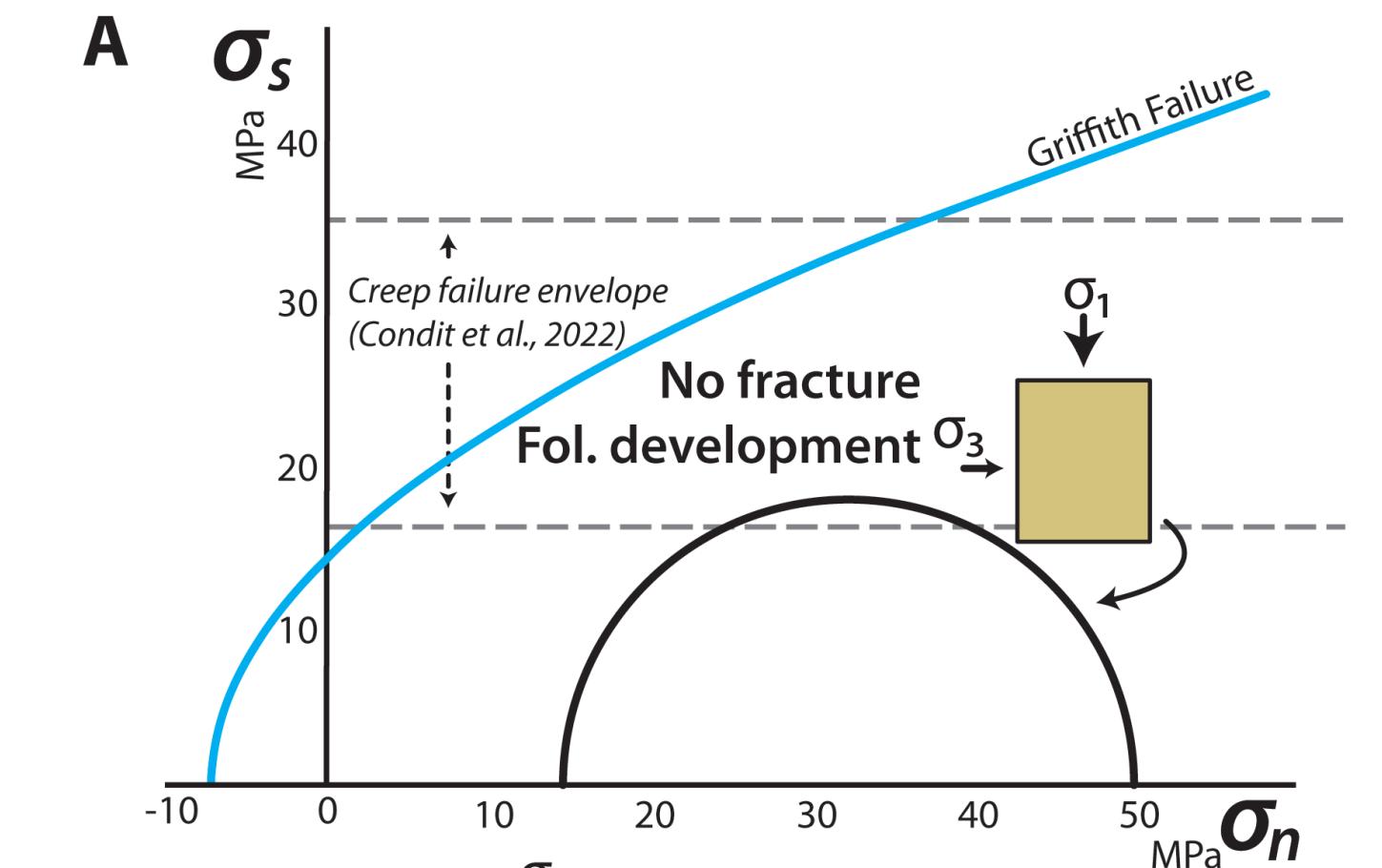
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Transformation-Assisted Creep, Ab porphyroblast/clasts reacting to quartz+ w/ fluid in Pelona schist



Activate frictional/brittle def.
Pore fluid pressure ($\sigma_n = \sigma - P_f$)

Observations from the exhumed rock record from the base of the subduction seismogenic zone

What are the phenomena we see in rocks that we need improved:

...Constitutive relations for?

...Dynamic mechanistic understanding of chemical-mechanical feedbacks?

Observations from the exhumed rock record from the base of the subduction seismogenic zone

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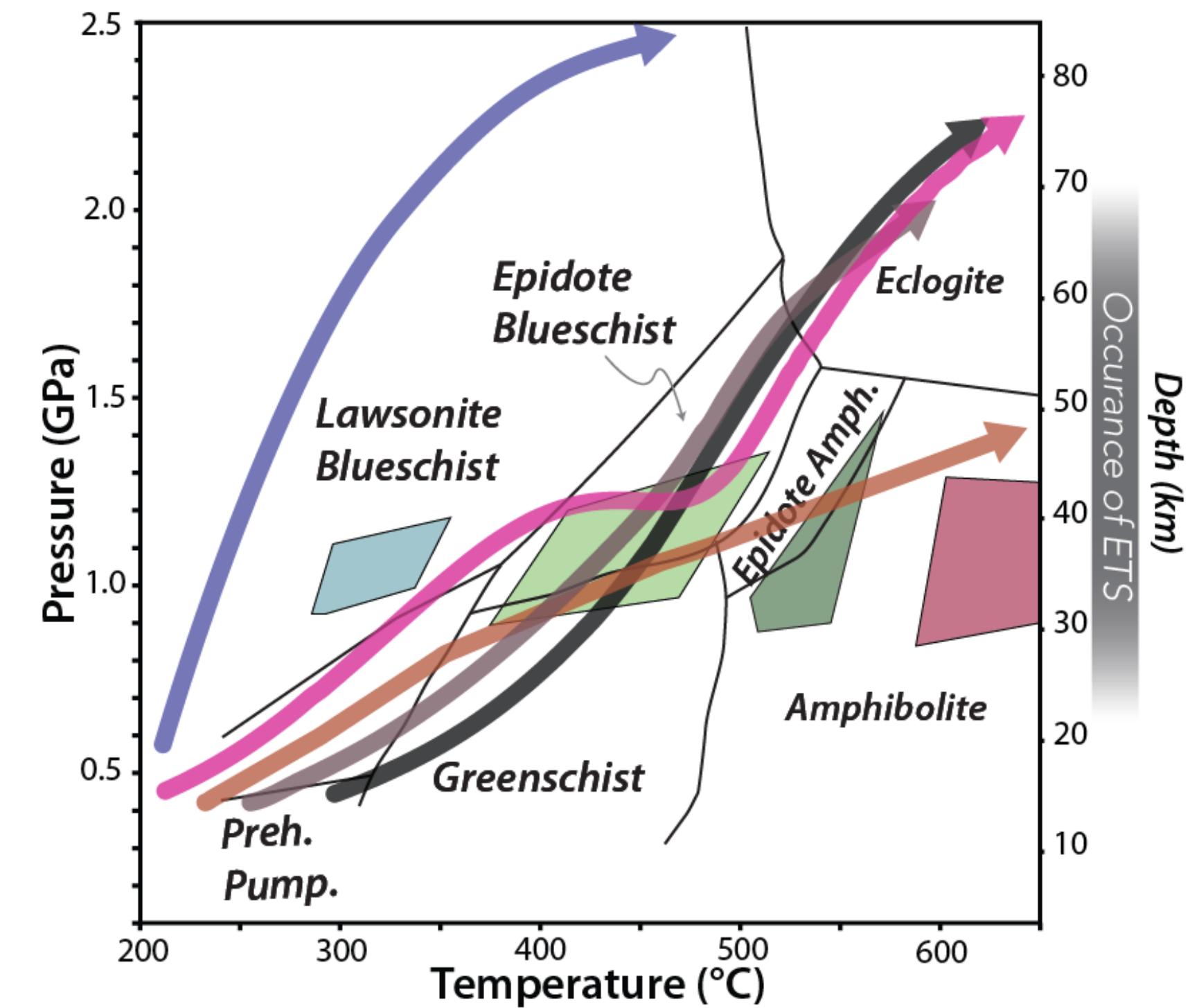
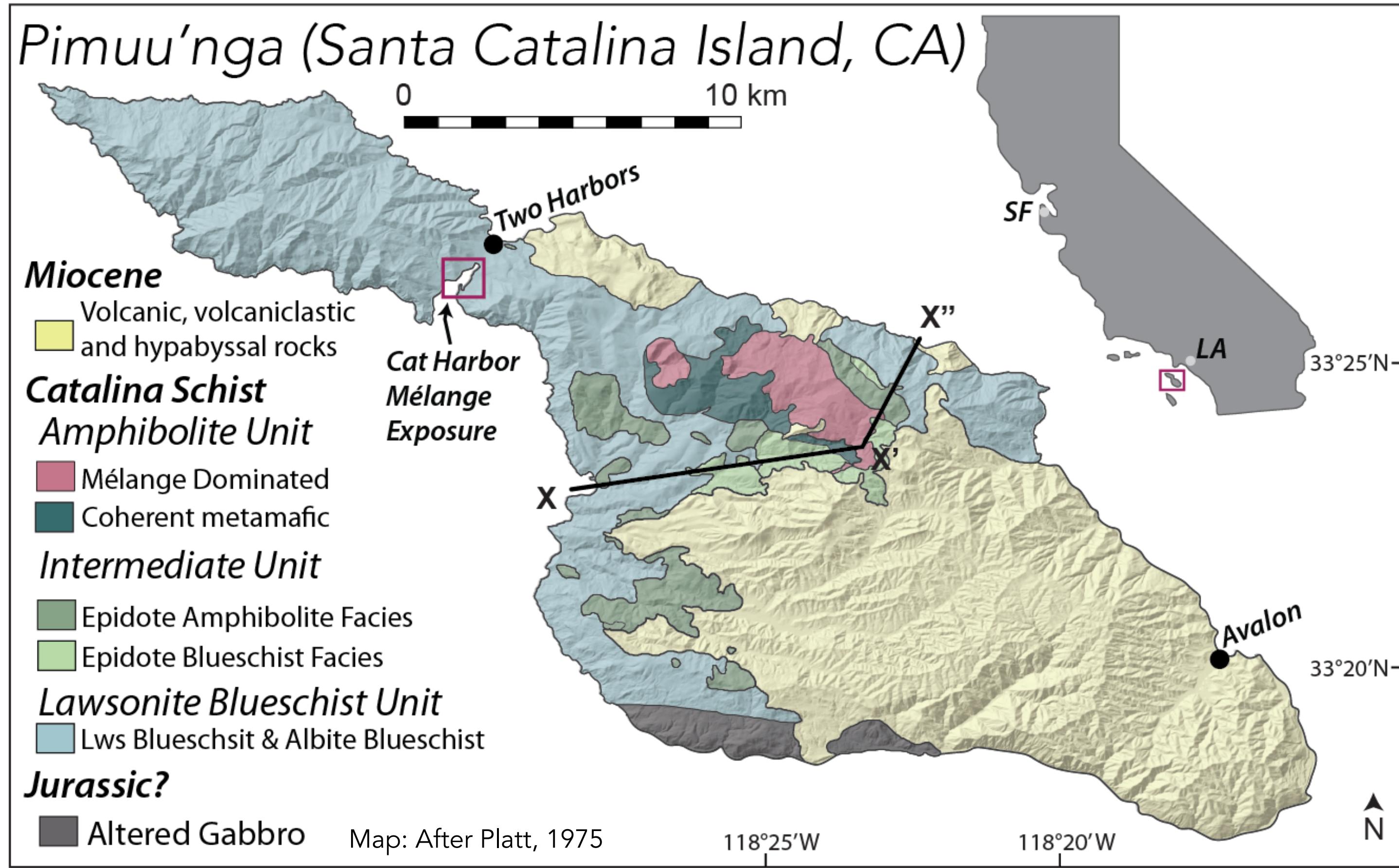
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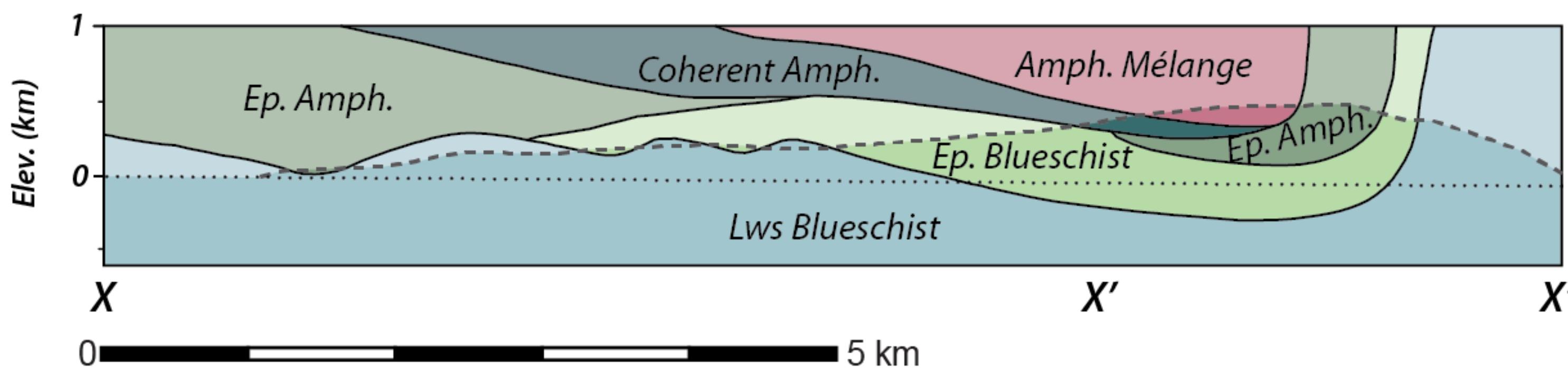
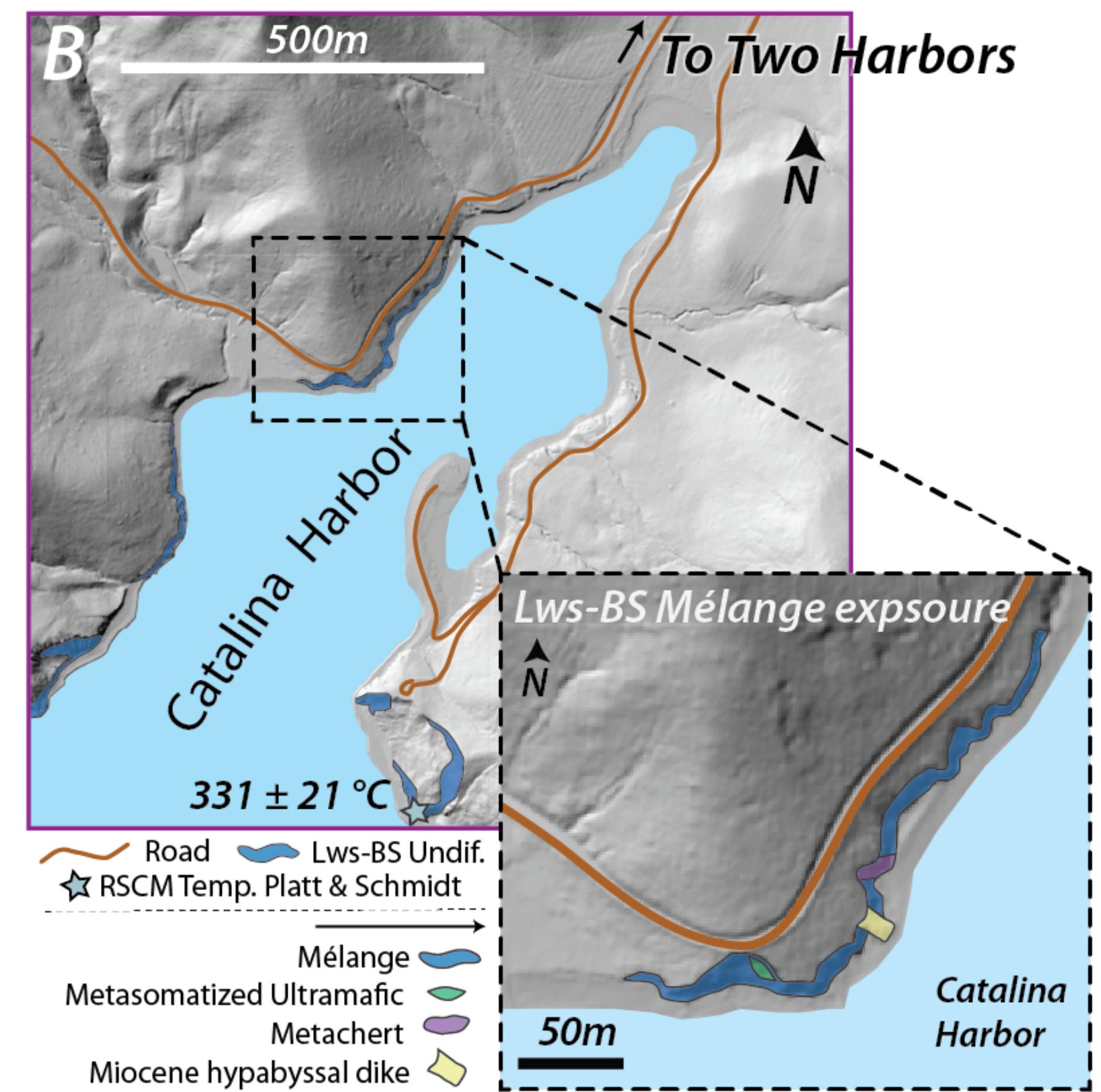
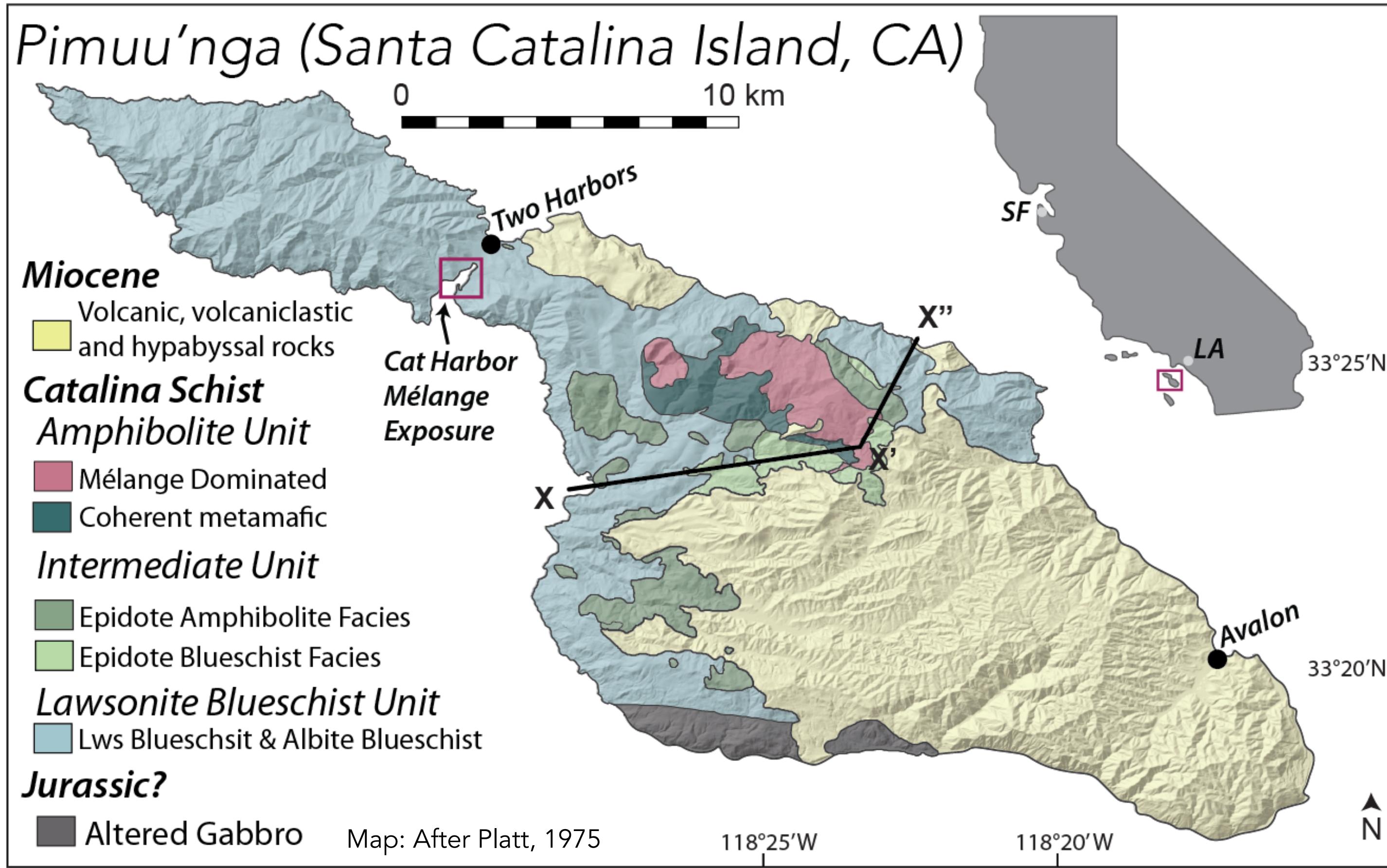
Pressure Solution Creep along the plate interface is not slow slip

Activation of frictional mechanisms with high Pf could be slow slip

Evidence for brittle deformation (**seismic slip?**) at these deep depths intimately related to fluid process



Exhumed set of terranes recording Cretaceous subduction beneath North America from 115 -95 Ma



Cross-section: Platt and Schmidt, 2024

Lawsonite blueschist unit - ~1.0 GPa (30-35 km) and ~320-330 °C

Temperatures: Platt and Schmidt, 2024

Looking to the WNW

Metasedimentary mélange

Block-in-matrix structure, quartz vein-rich mess



Cat Harbor Mélange, Pimuu'nga (Santa Catalina Island) · CA, USA

Looking to the WNW



Cat Harbor Mélange, Pimuu'nga (Santa Catalina Island) - CA, USA



Looking to SE

Condit et al., 2025 Geology

Metasedimentary mélange

Block-in-matrix structure, quartz vein-rich exposure

**How do these structures form? How do they deform?
and what slip behaviors do they accommodate?**



Looking to NW



Looking to NW



Metasedimentary mélange

Block-in-matrix structure, quartz vein-rich exposure

**How do these structures form? How do they deform?
and what slip behaviors do they accommodate?**

Extensional veins in fold noses



Outer-arc extension fractures precipitate quartz;
Transform into coarse grained quartz vein-mesh networks in
fold noses

Looking to NW

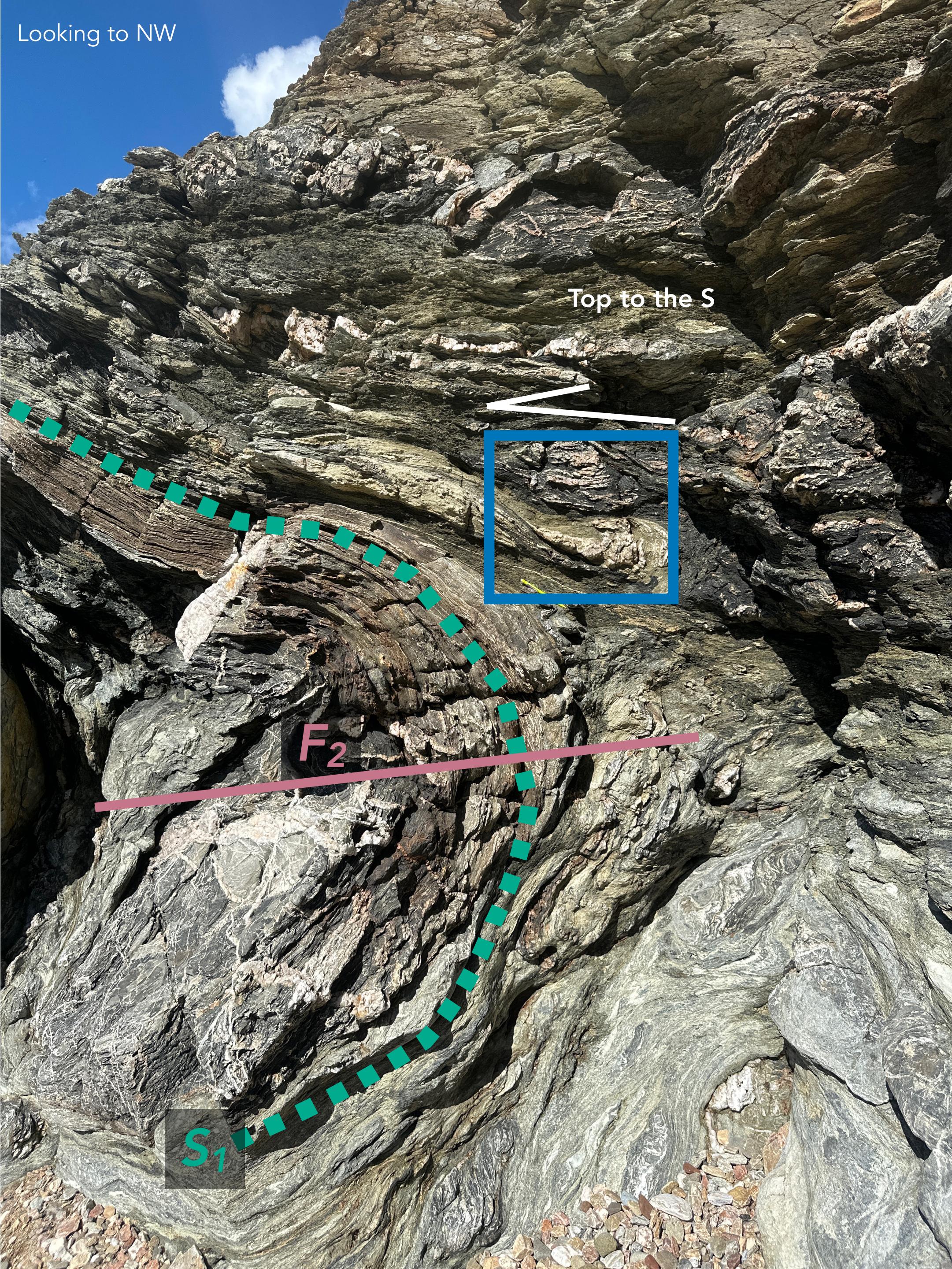


Extensional veins in fold noses

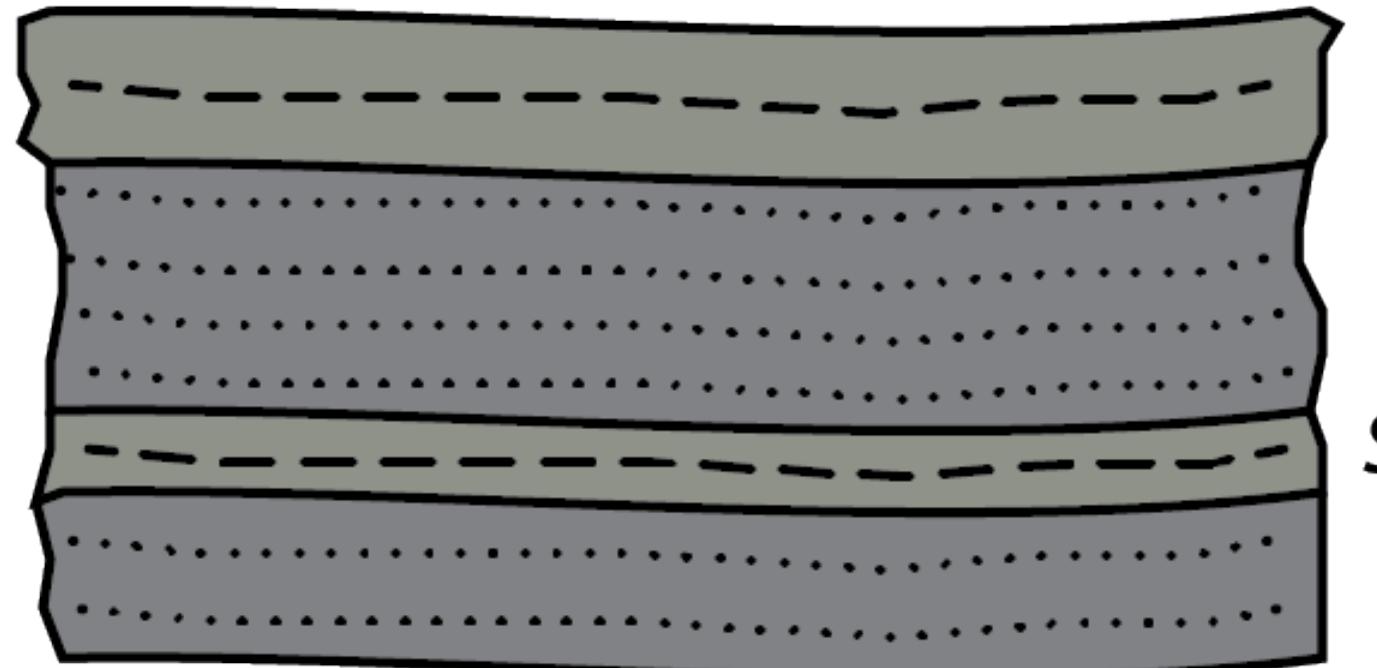


Vein-filled fold noses become blocks in the mélange

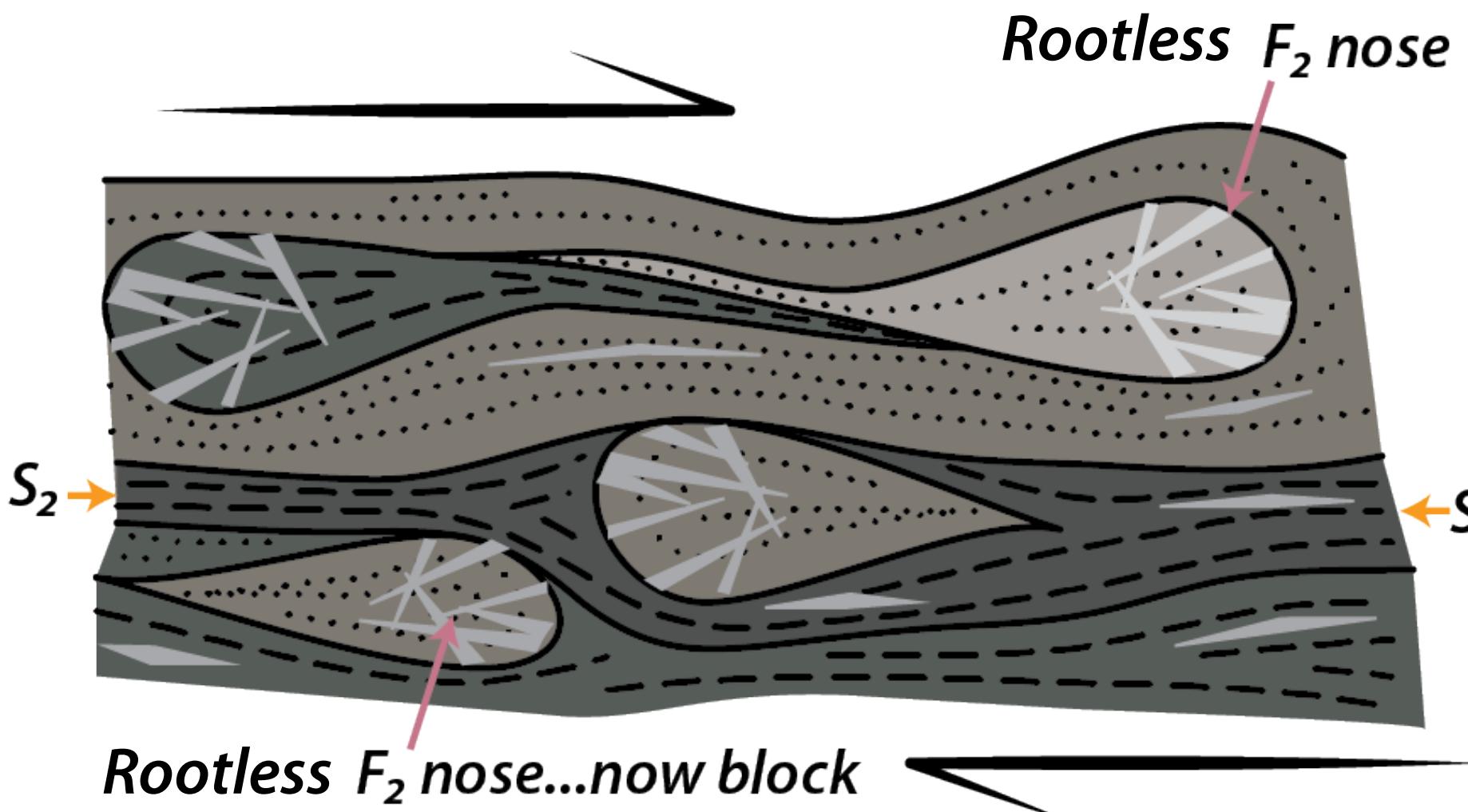
Looking to NW



Subduction, deformation fluid infiltration & metamorphism

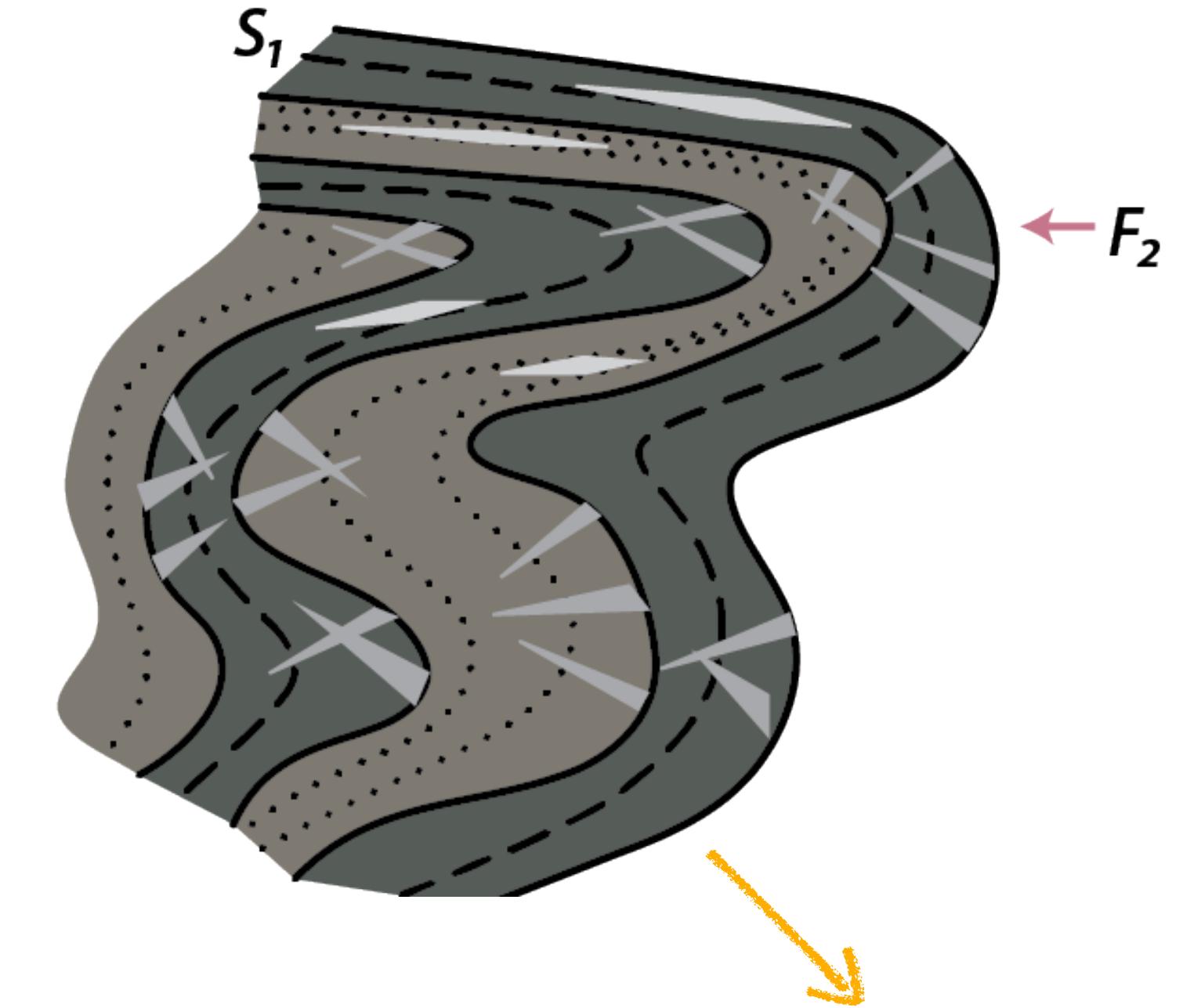


Deposition near the trench:
~100 Ma; graywackes, sandstones, shales

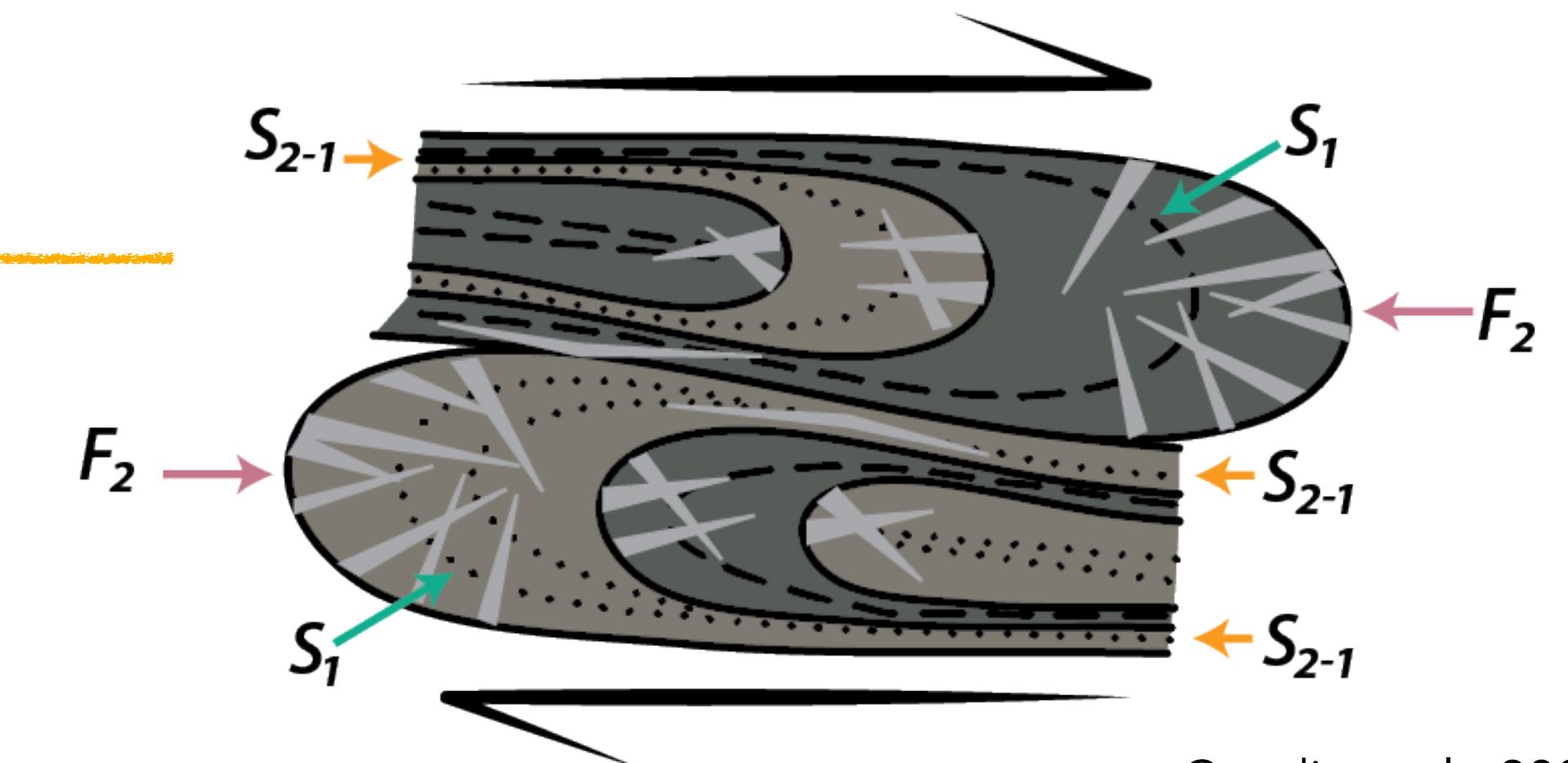


Rootless fold noses = strong blocks
Fold Limbs (F_2) = weak matrix

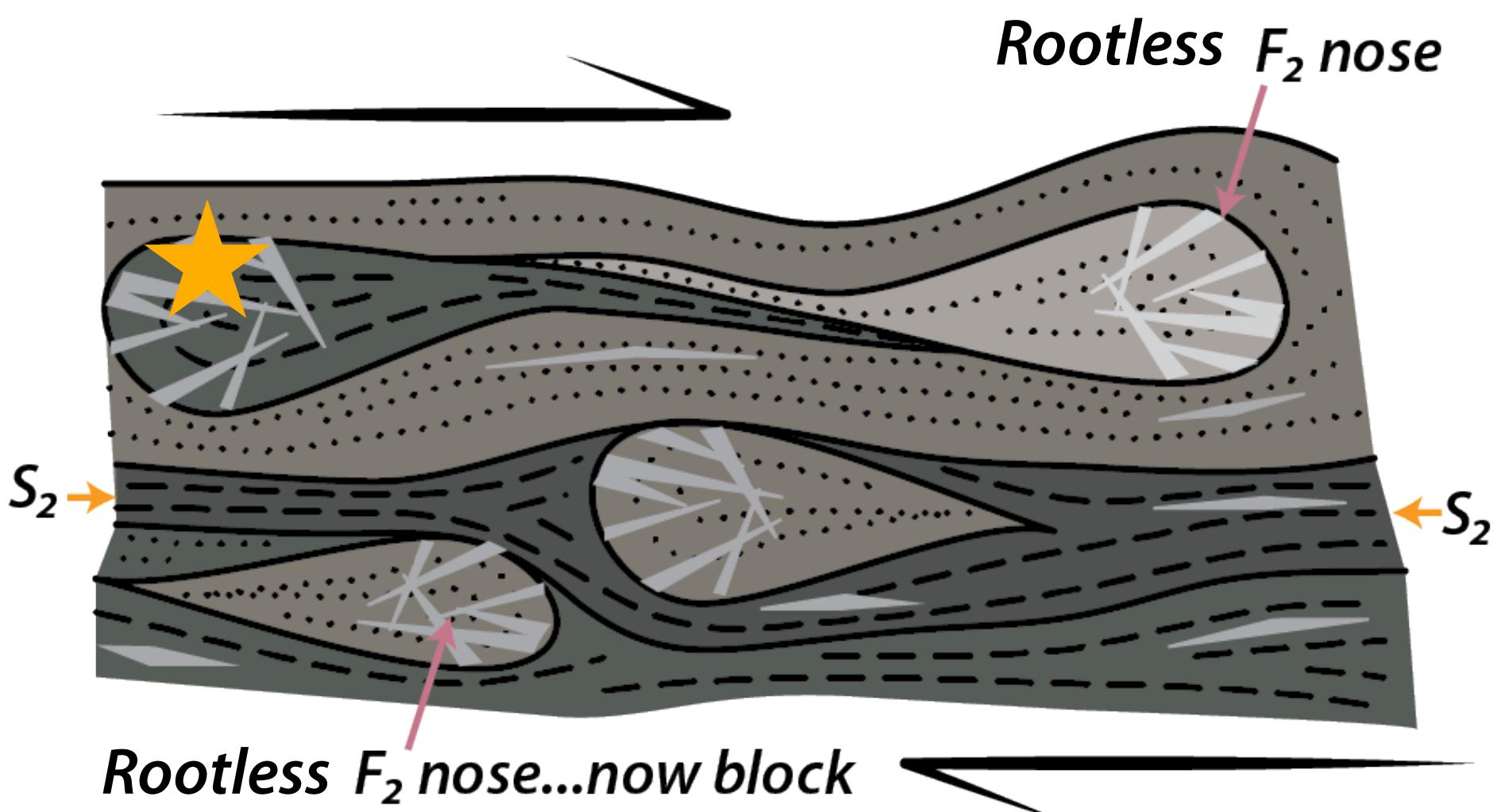
Folding, veining, during prograde subduction



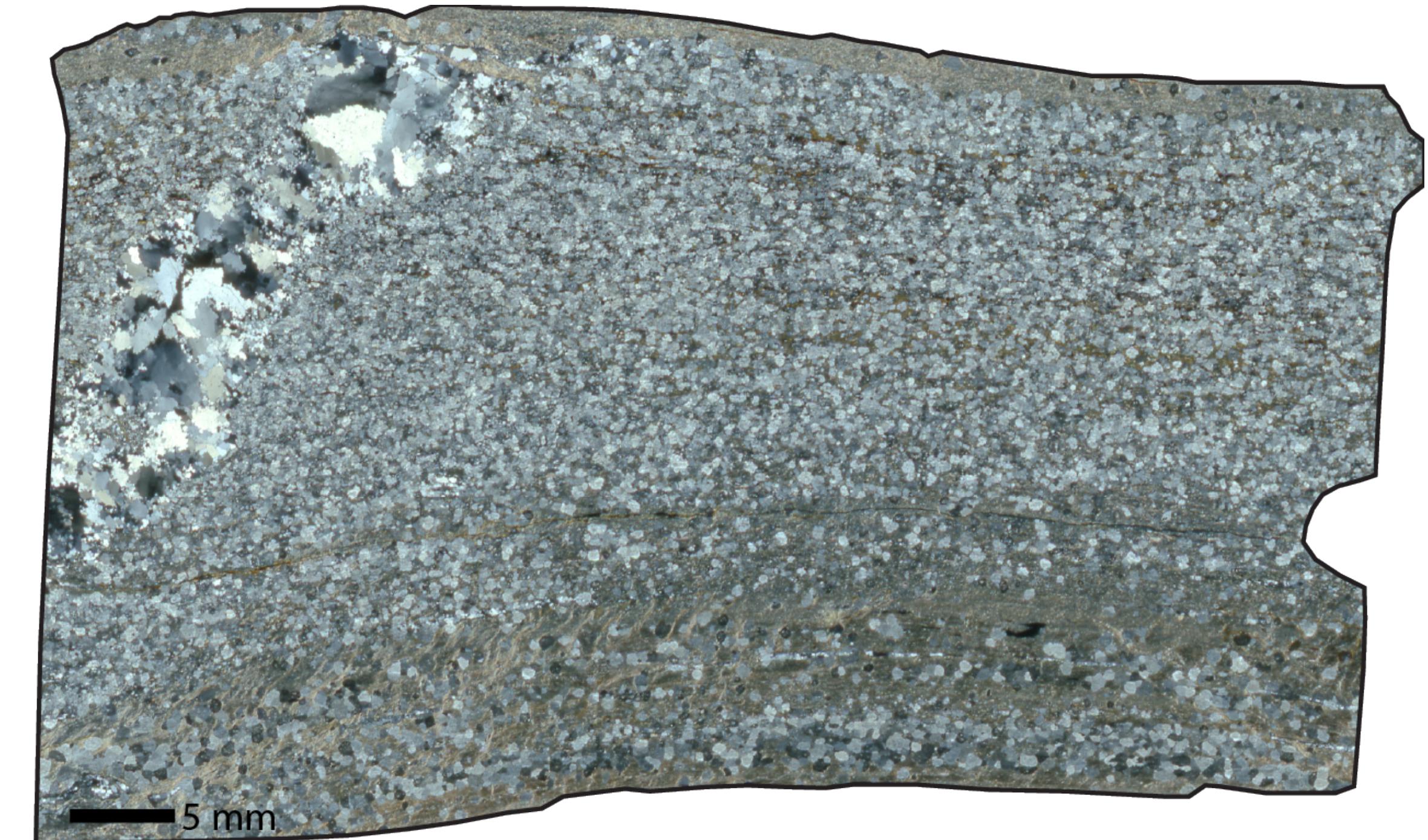
Veins strengthen fold noses, progressive shearing
with continued subduction, strain localization in limbs/matrix

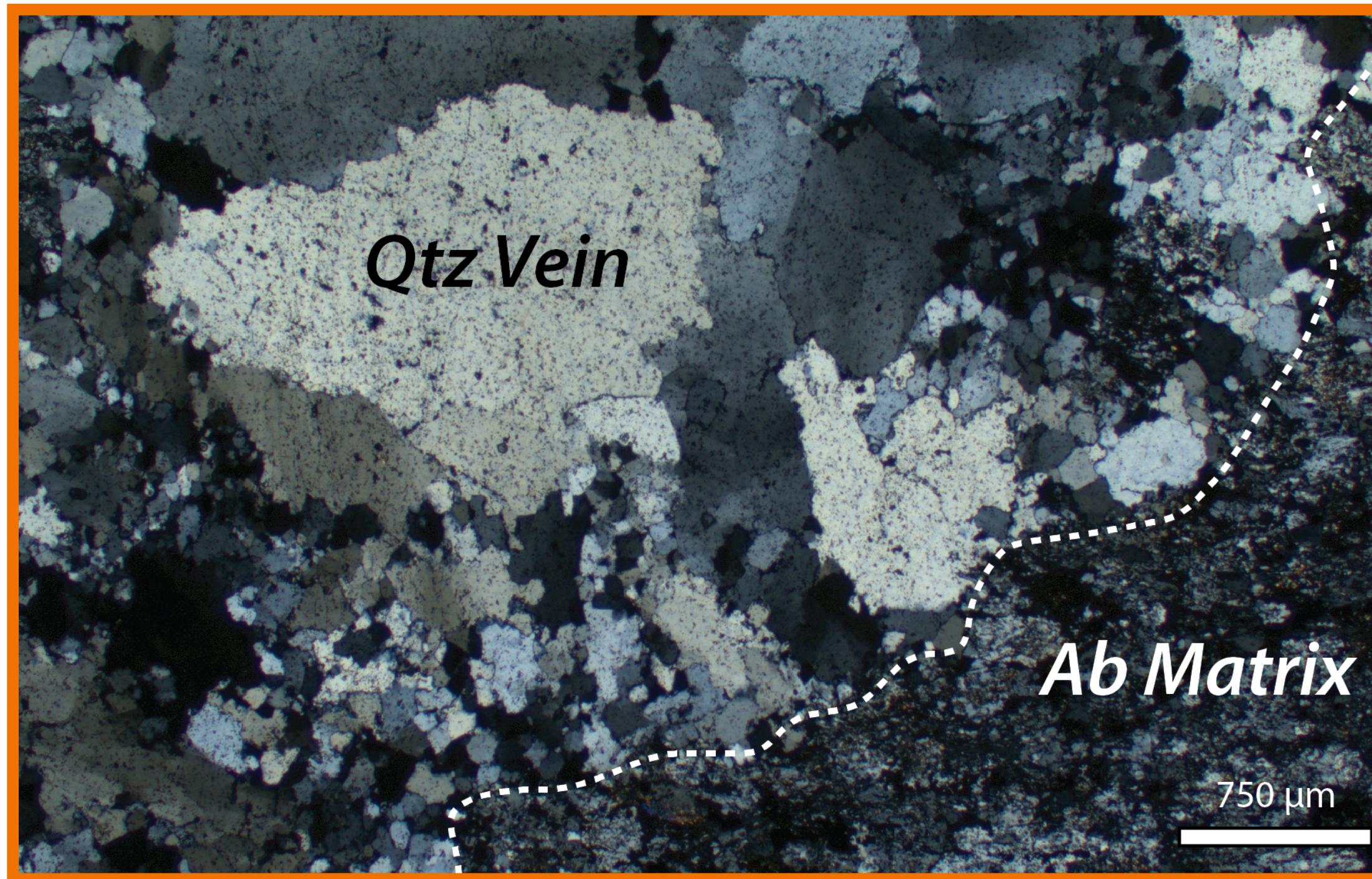


Rootless Fold Nose/Block

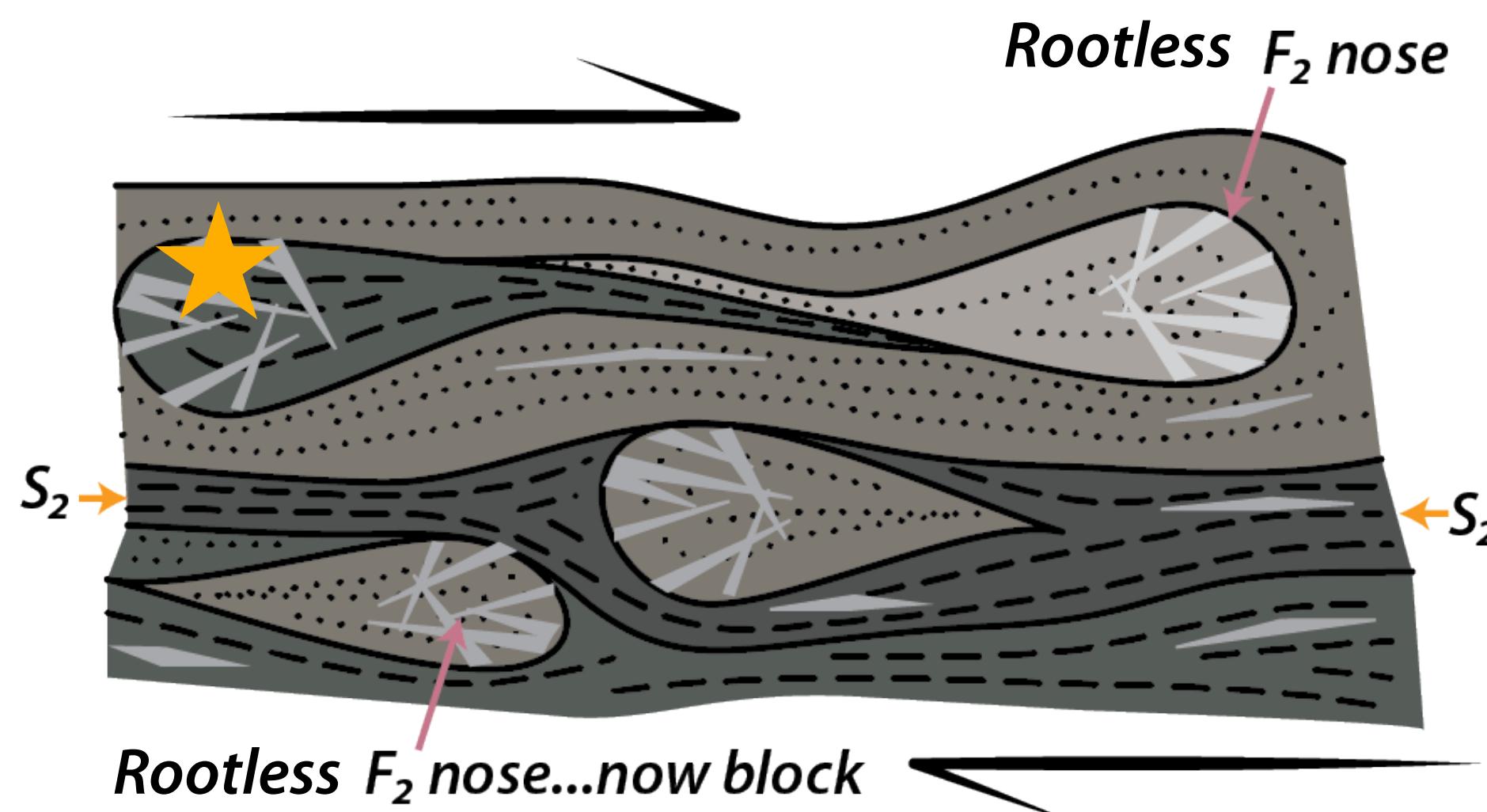
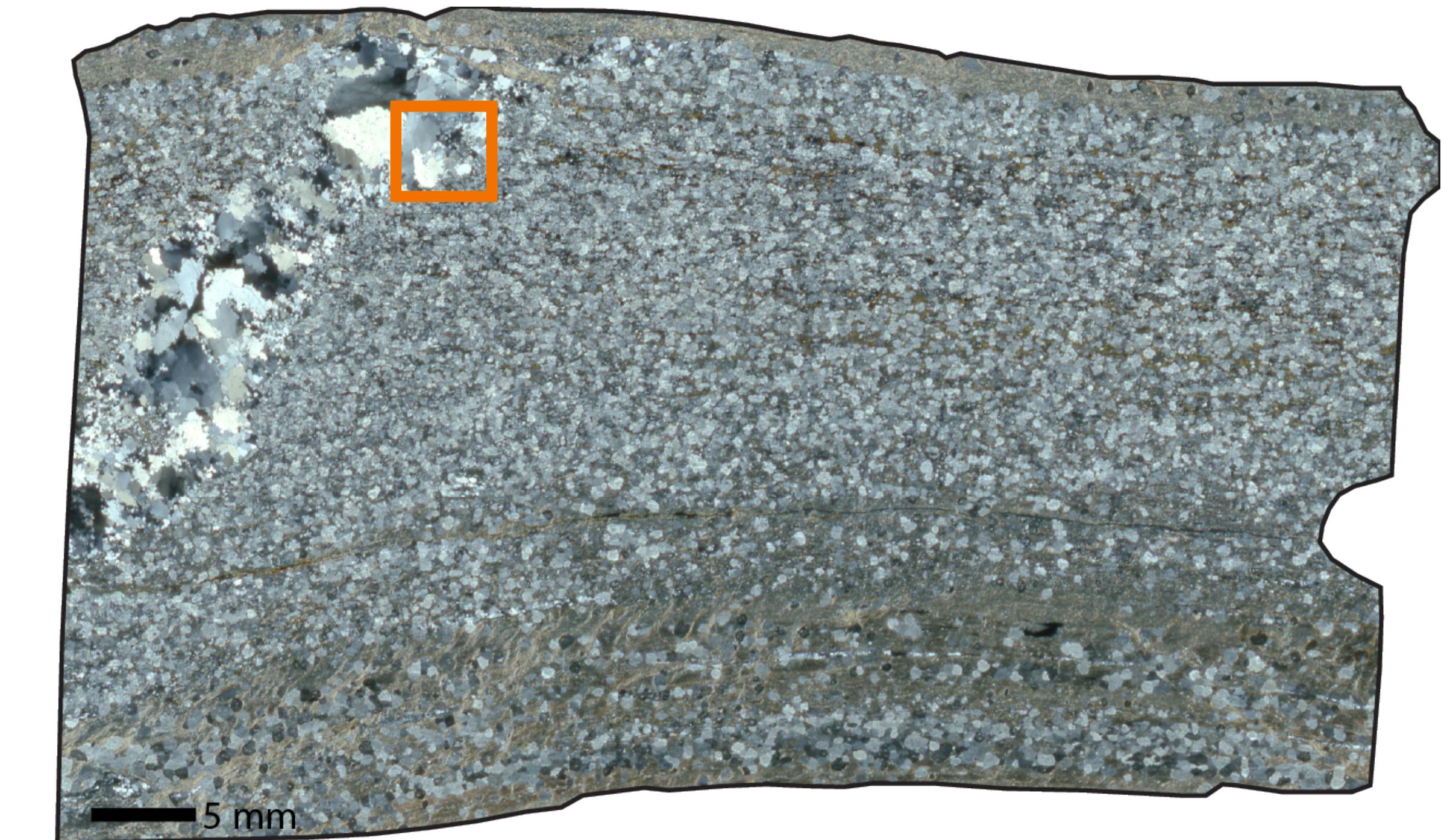


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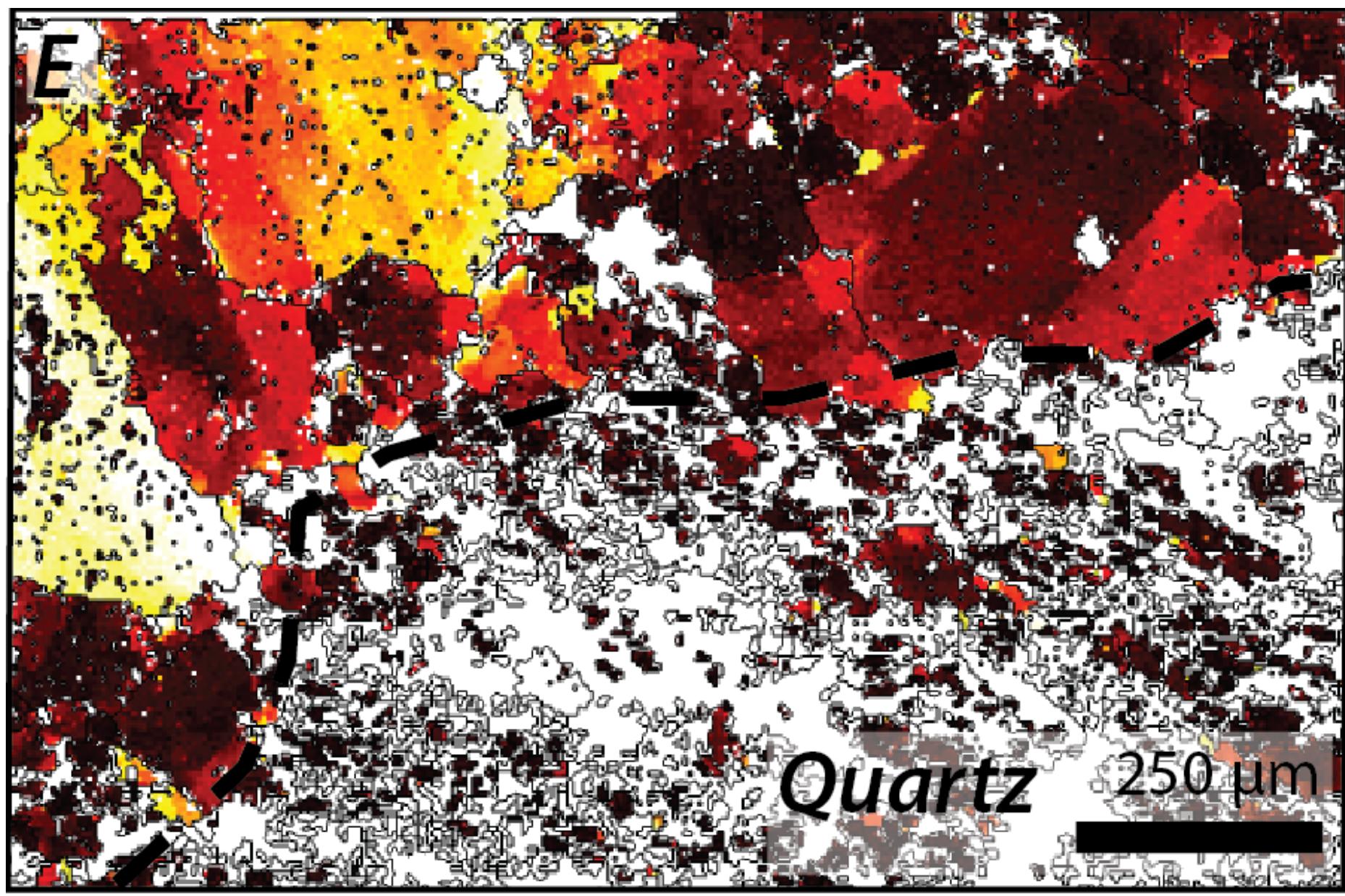
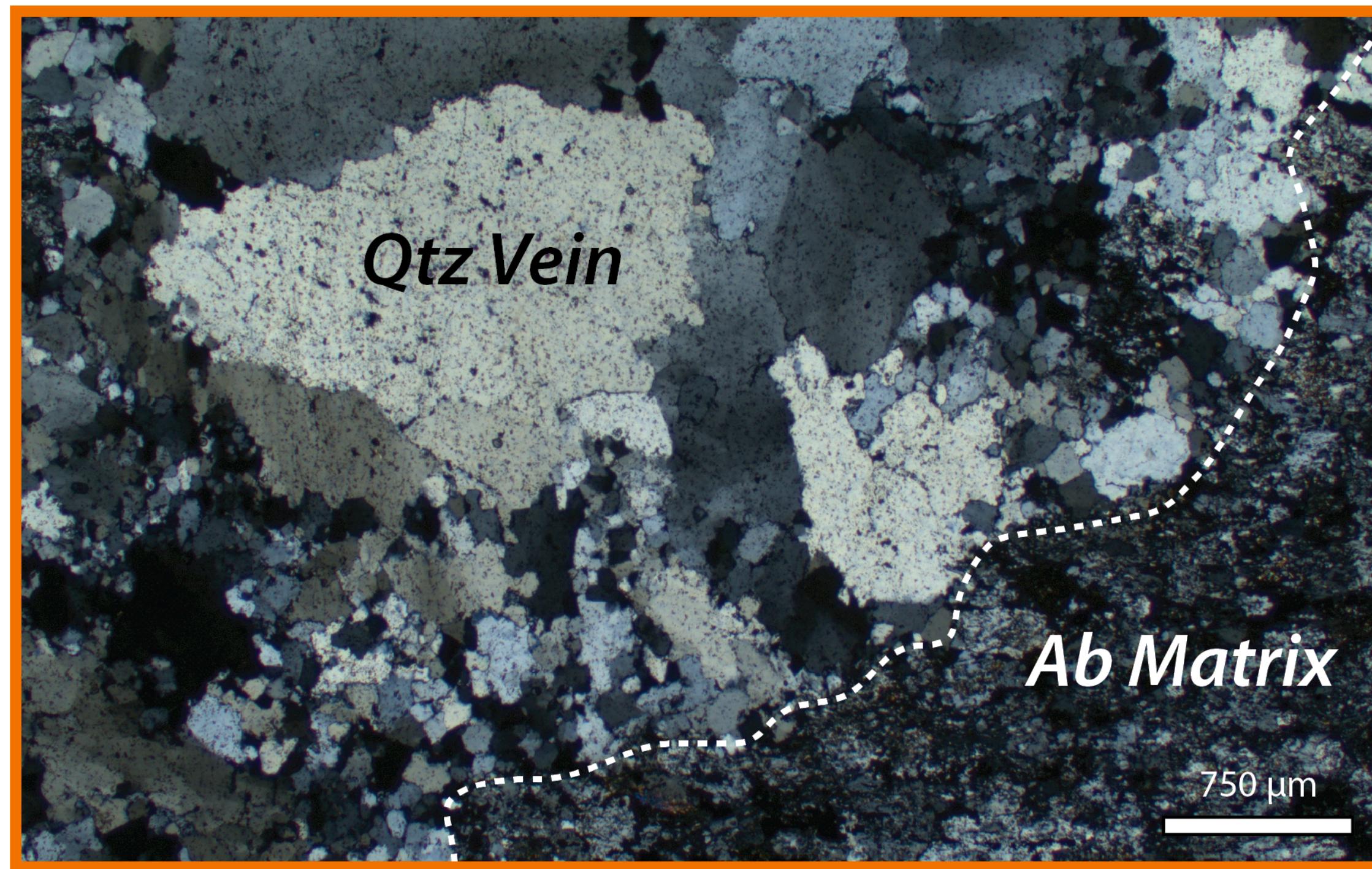


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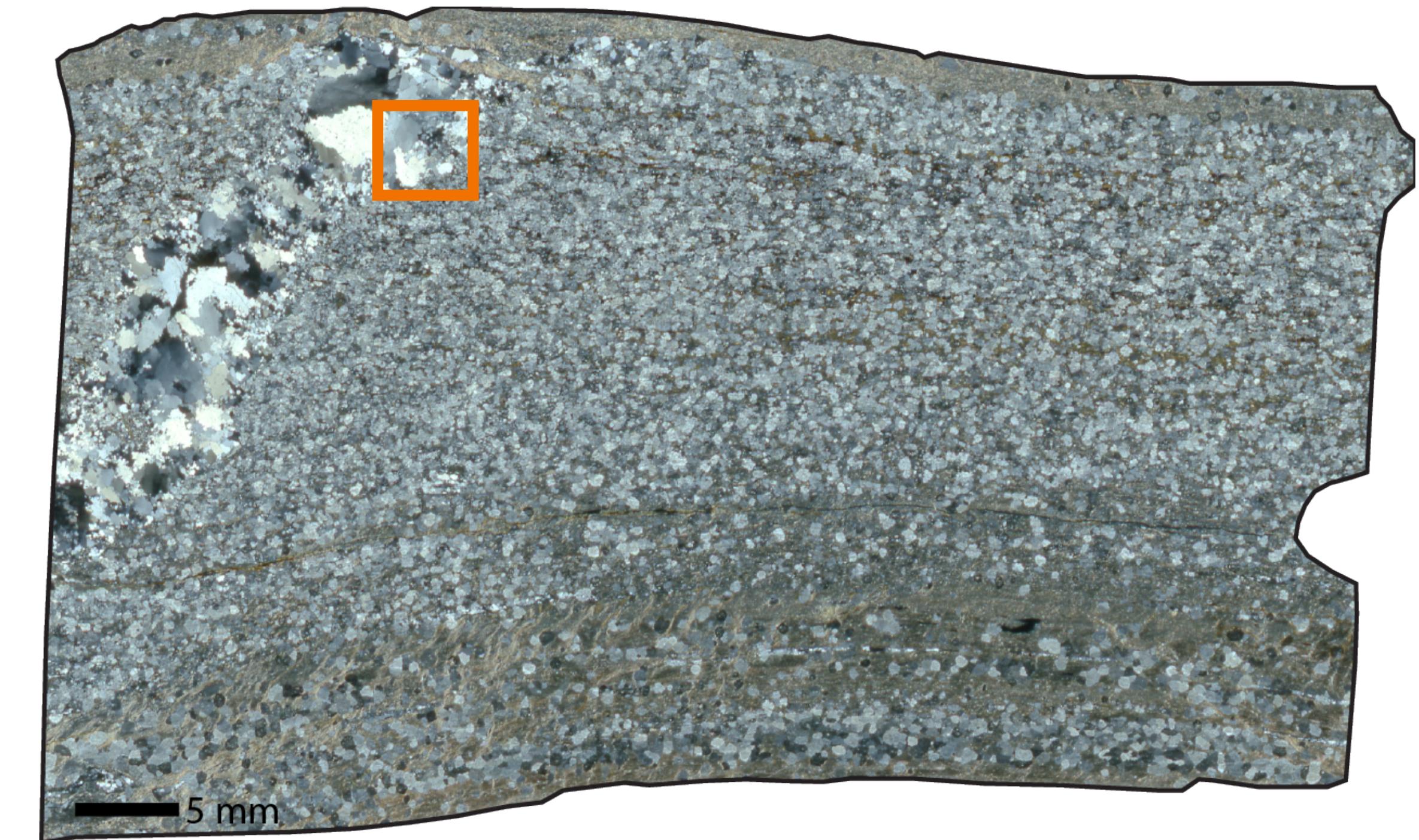
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Veins - Cross-cut early fabric (S_1); deform by BLG and incipient SGR; temps of 300-350°C

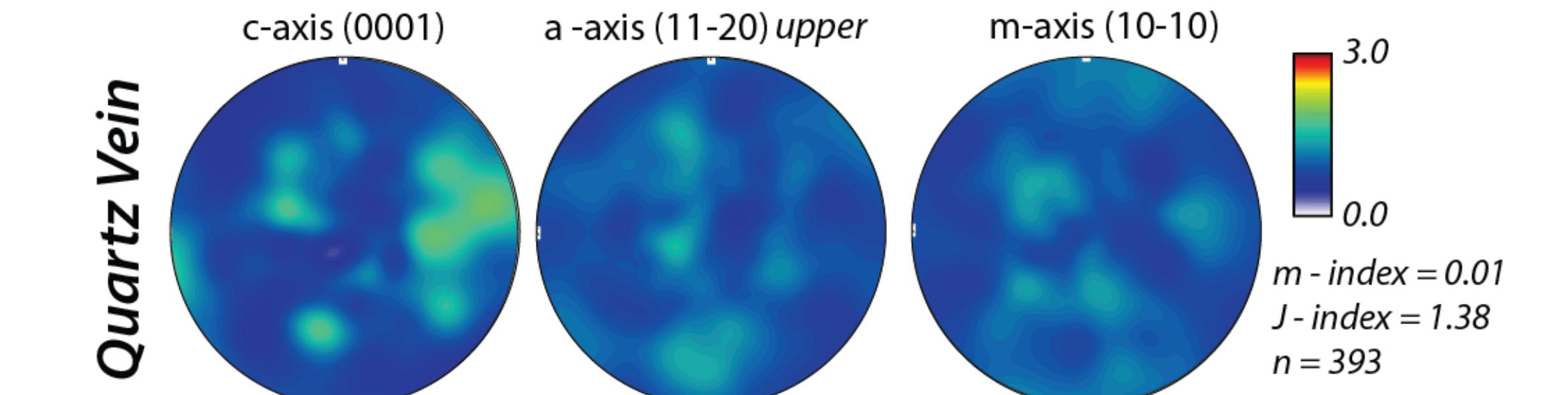


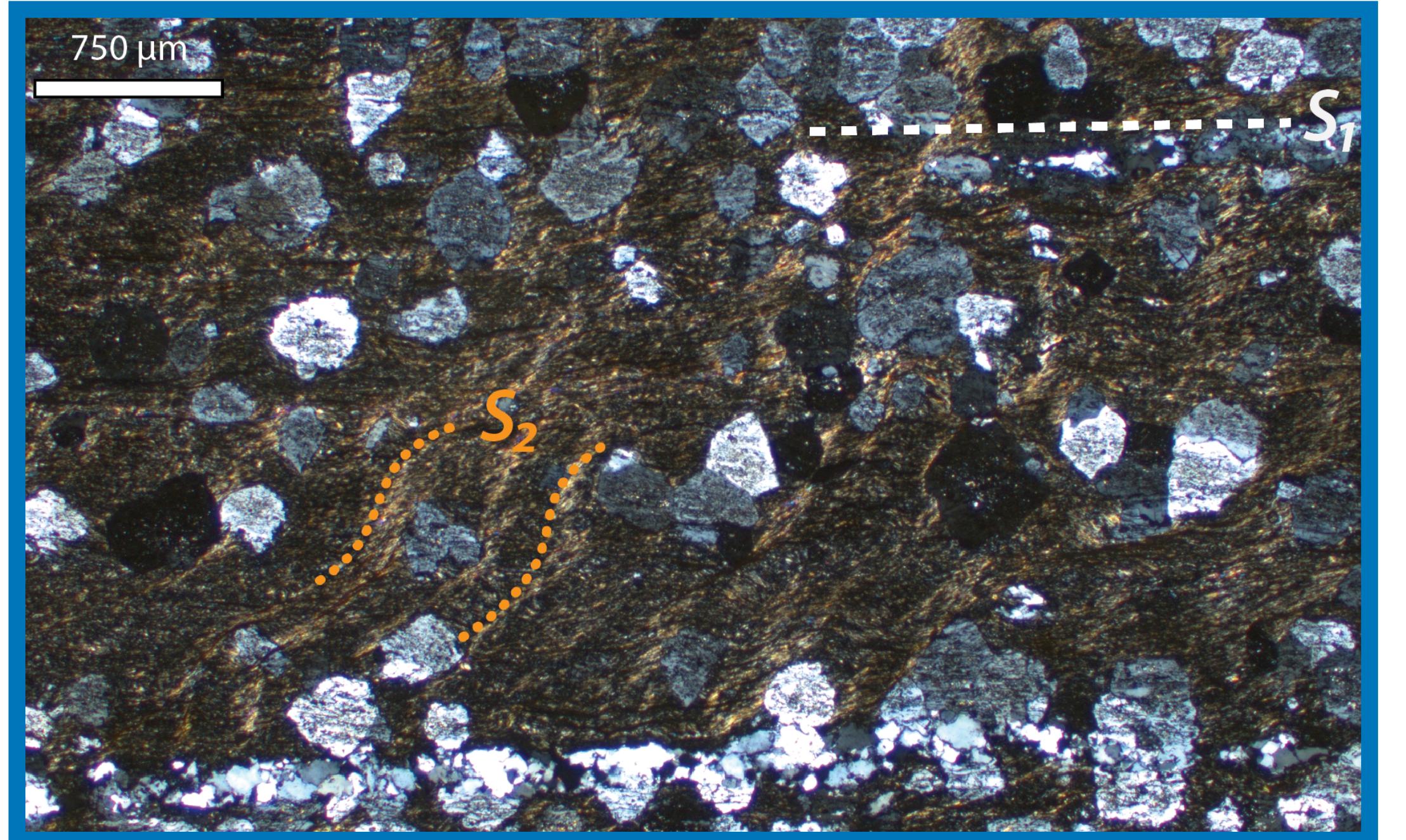
Quartz Mis2Mean EBSD map of vein

Rootless Fold Nose/Block

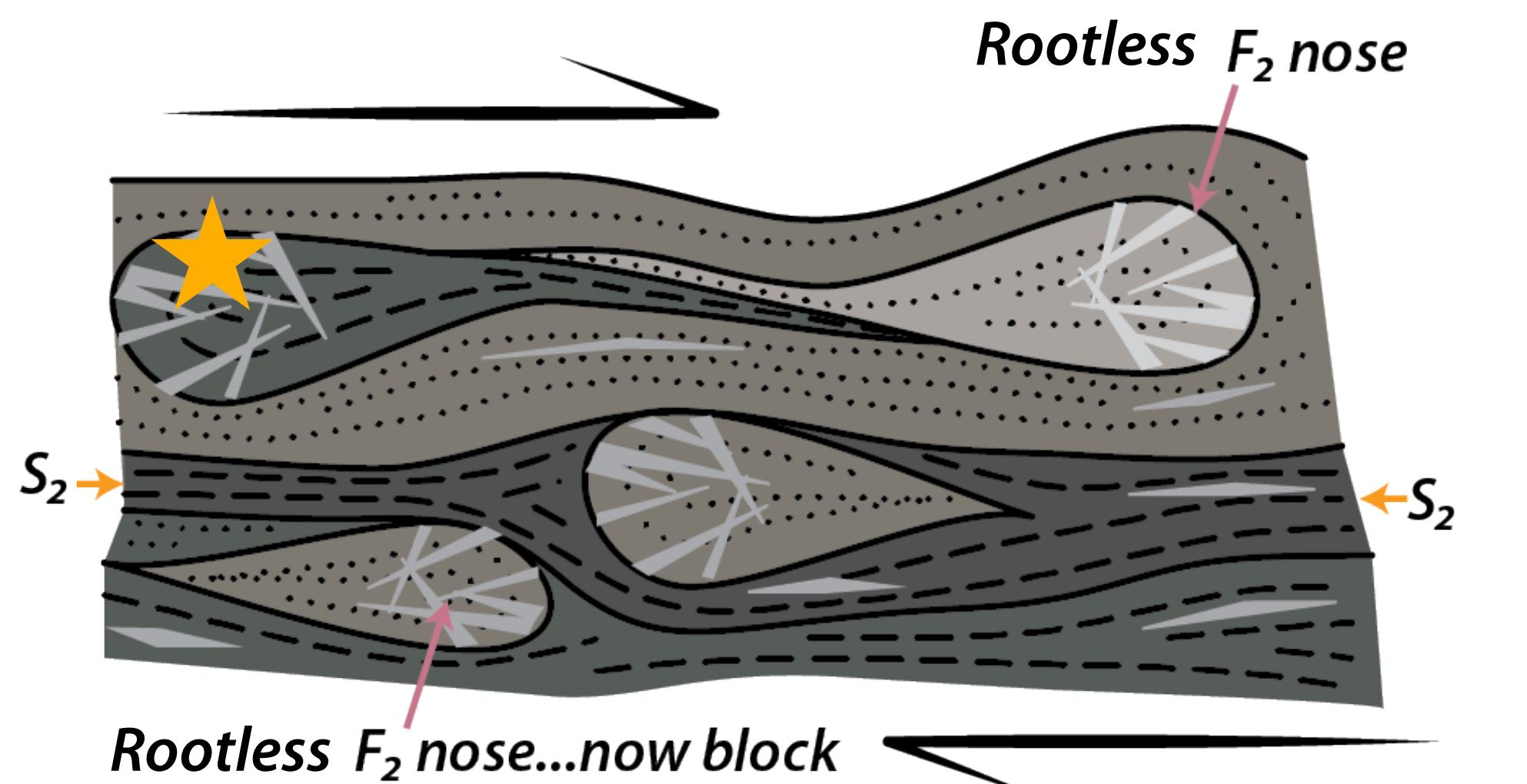
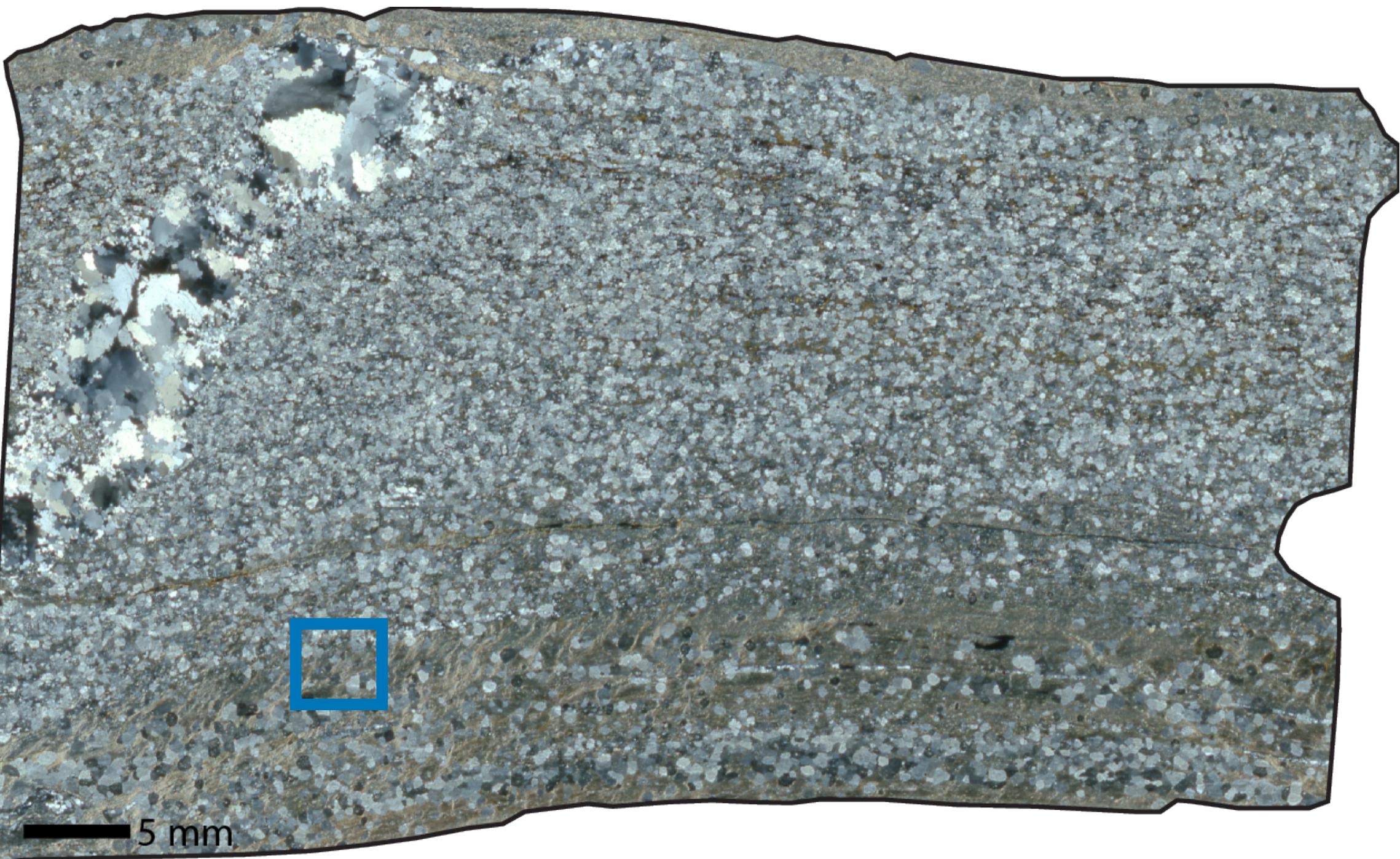


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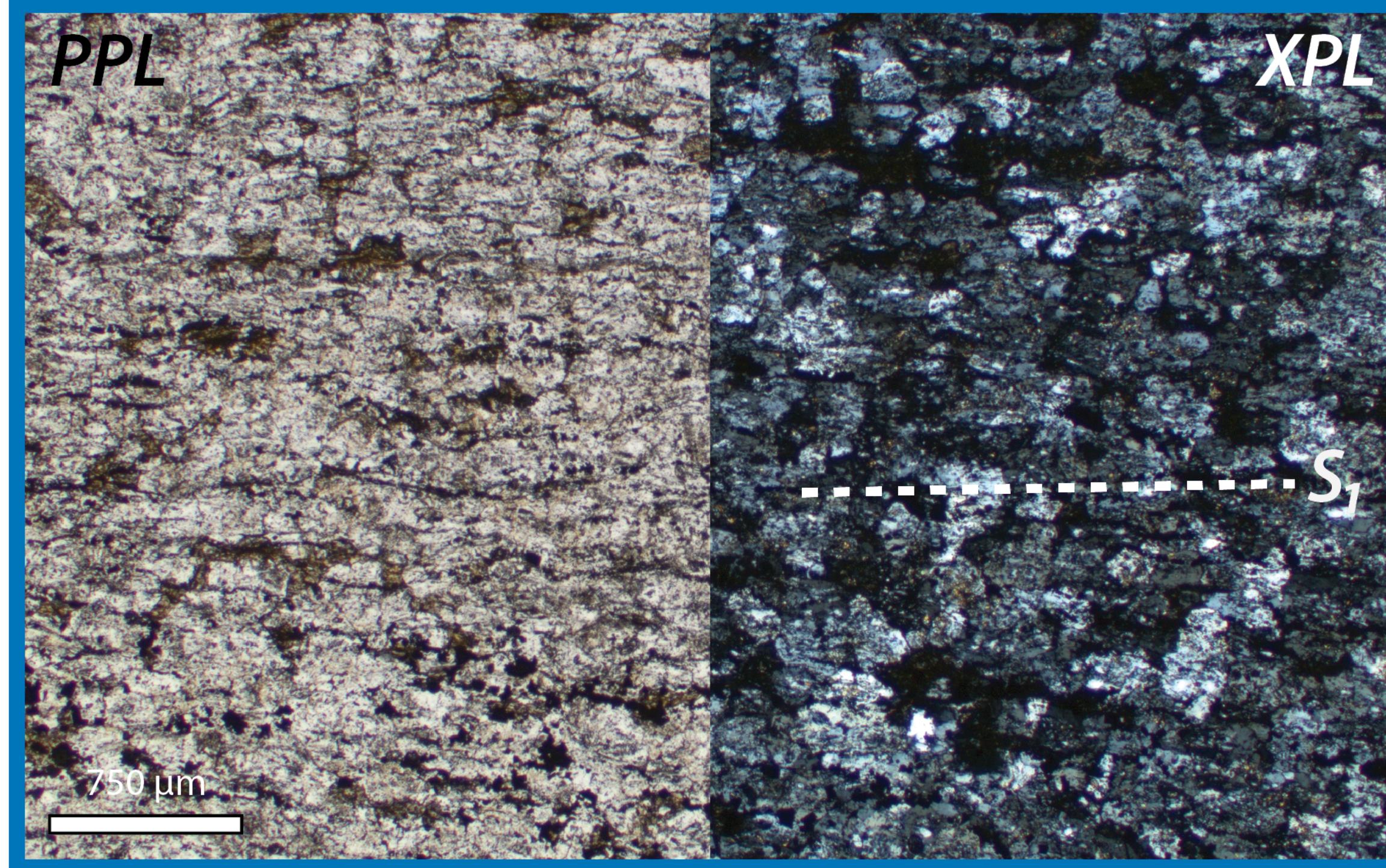
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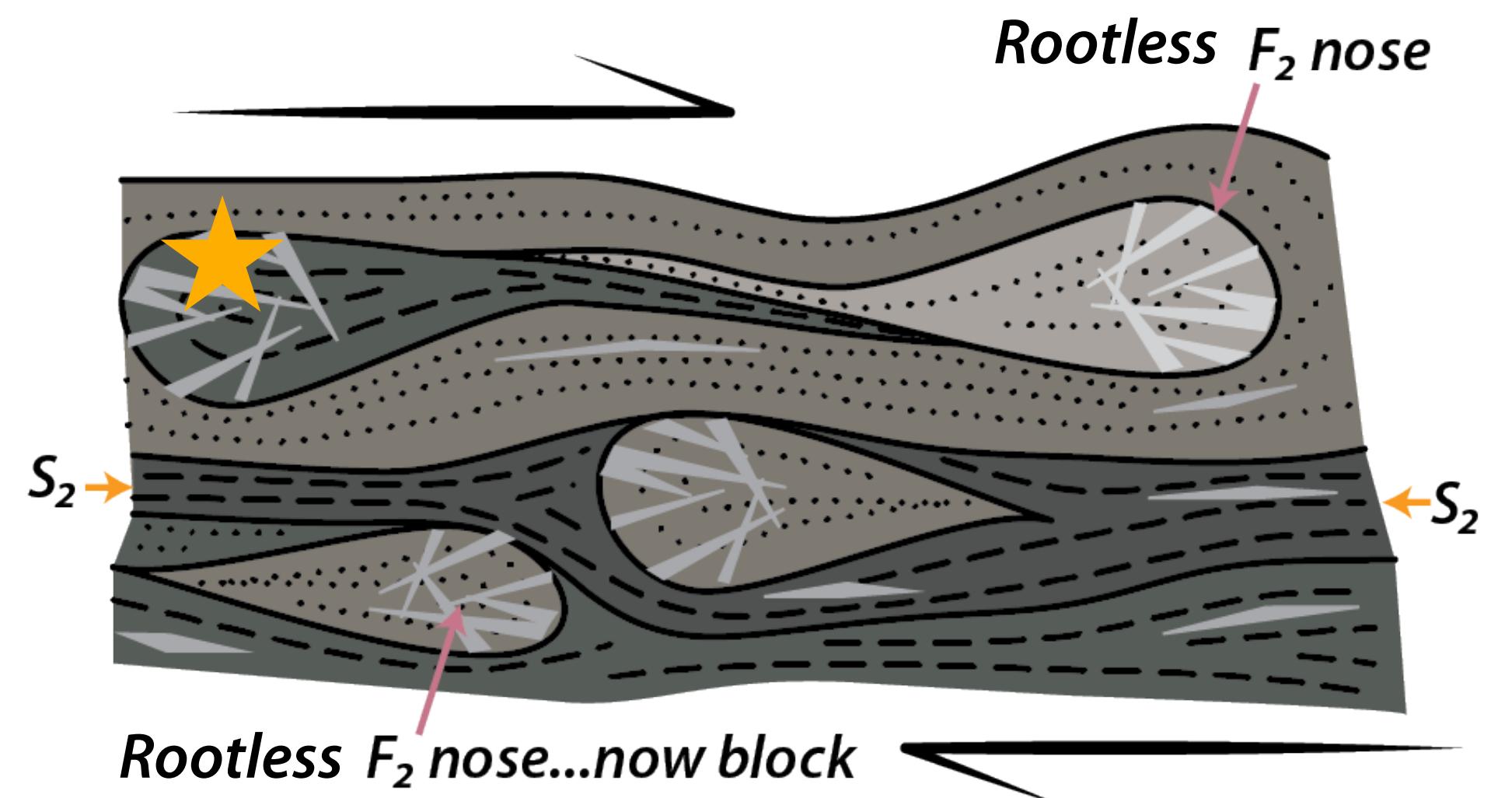
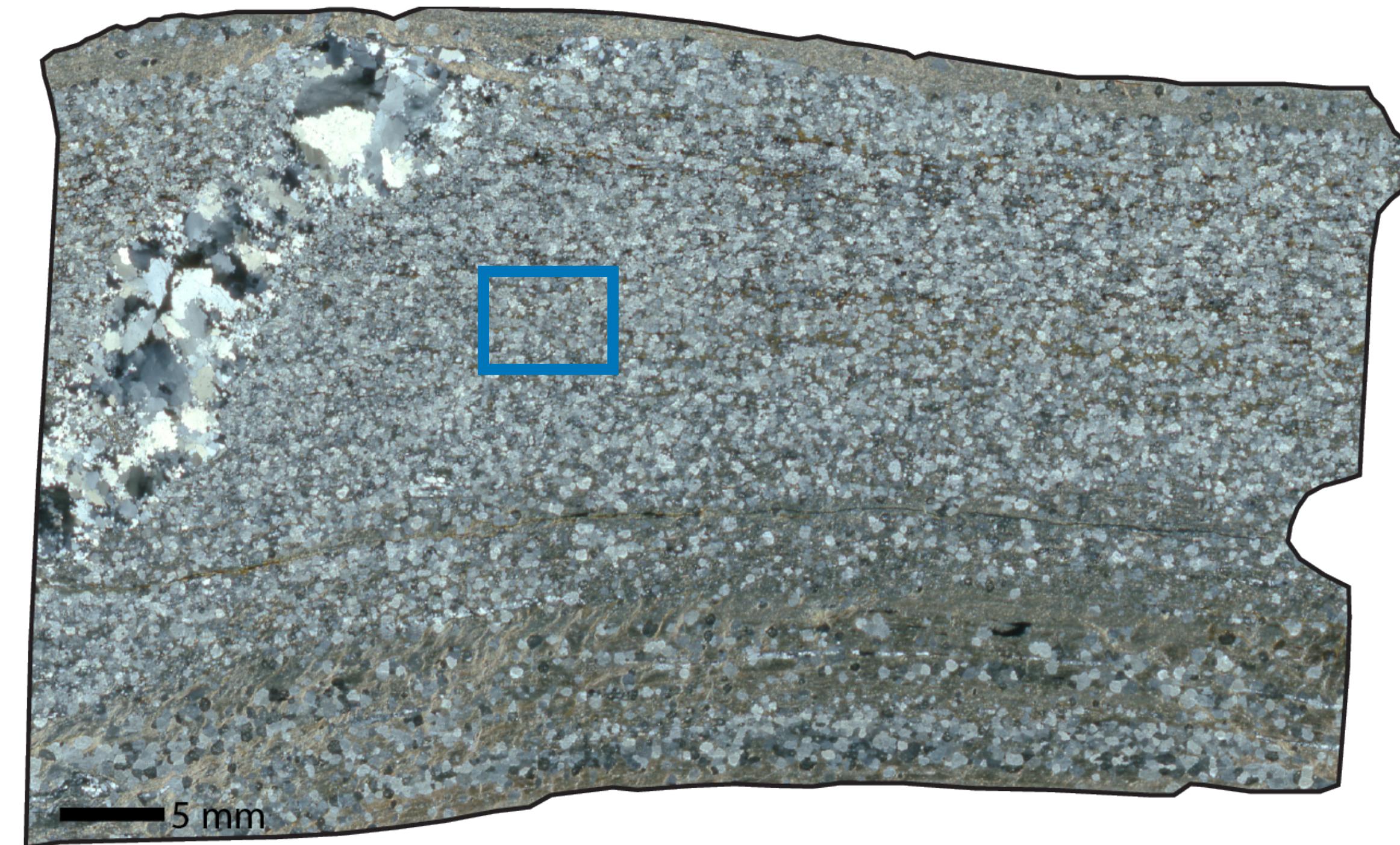
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F₂ Limbs = weak matrix

Veins - Cross-cut early fabric (S₁); deform by BLG and incipient SGR; temps of 300-350°C

Early fabric (S₁)- Kinking in sheet silicates; albite porphyroblasts/clasts with grain size >>50-70 μ m some pressure-solution creep microstructures but not dominant



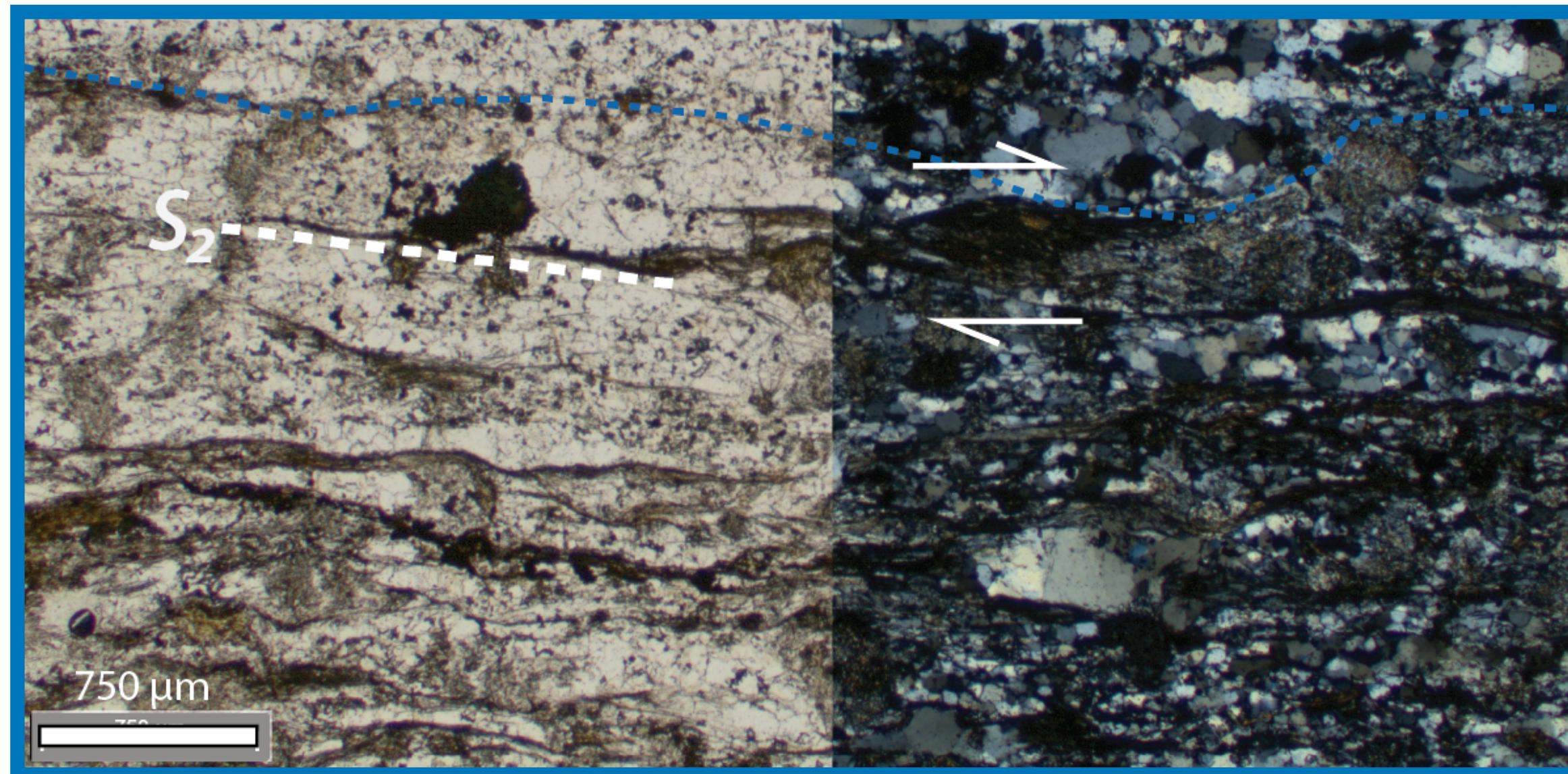
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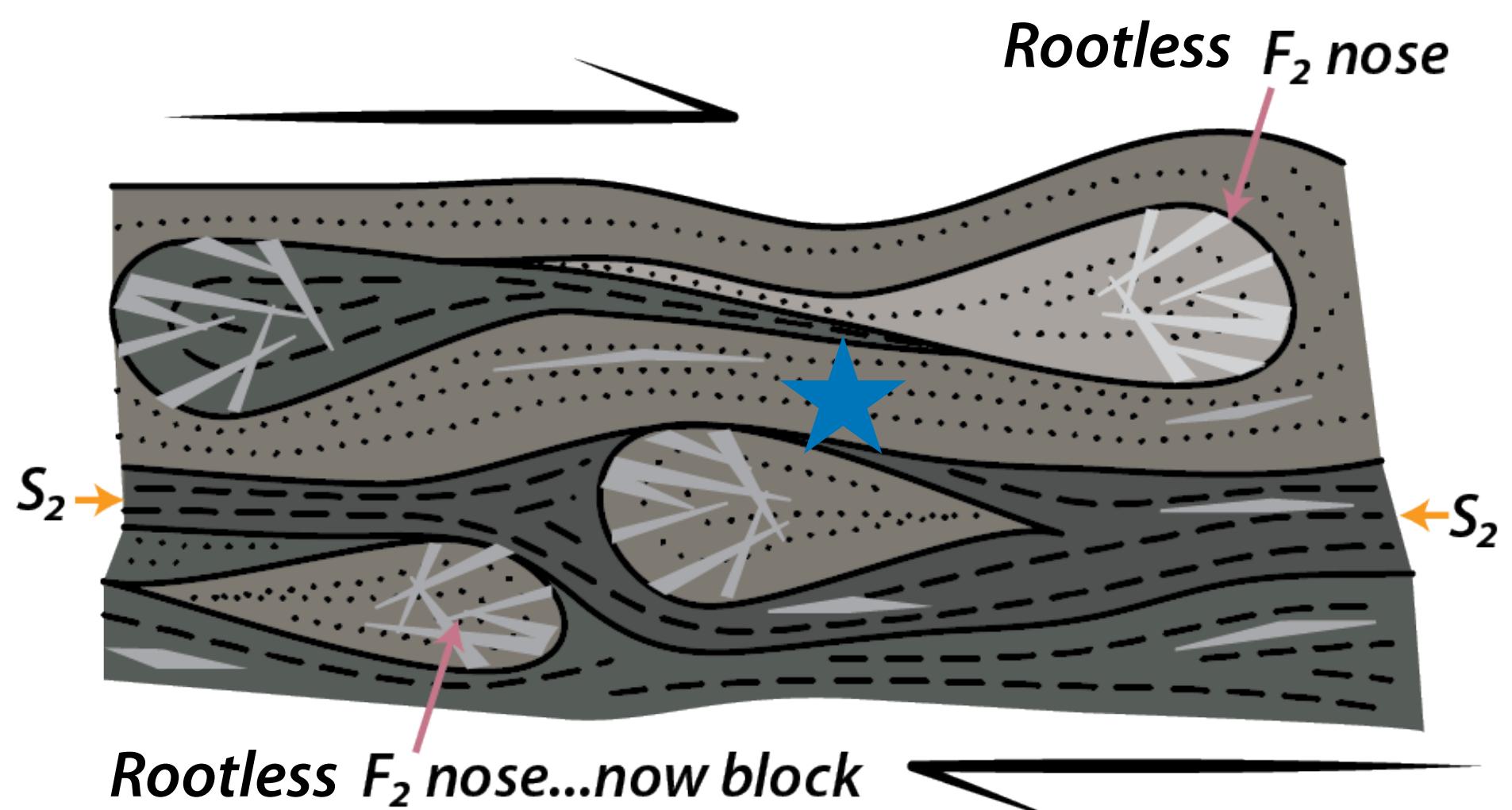
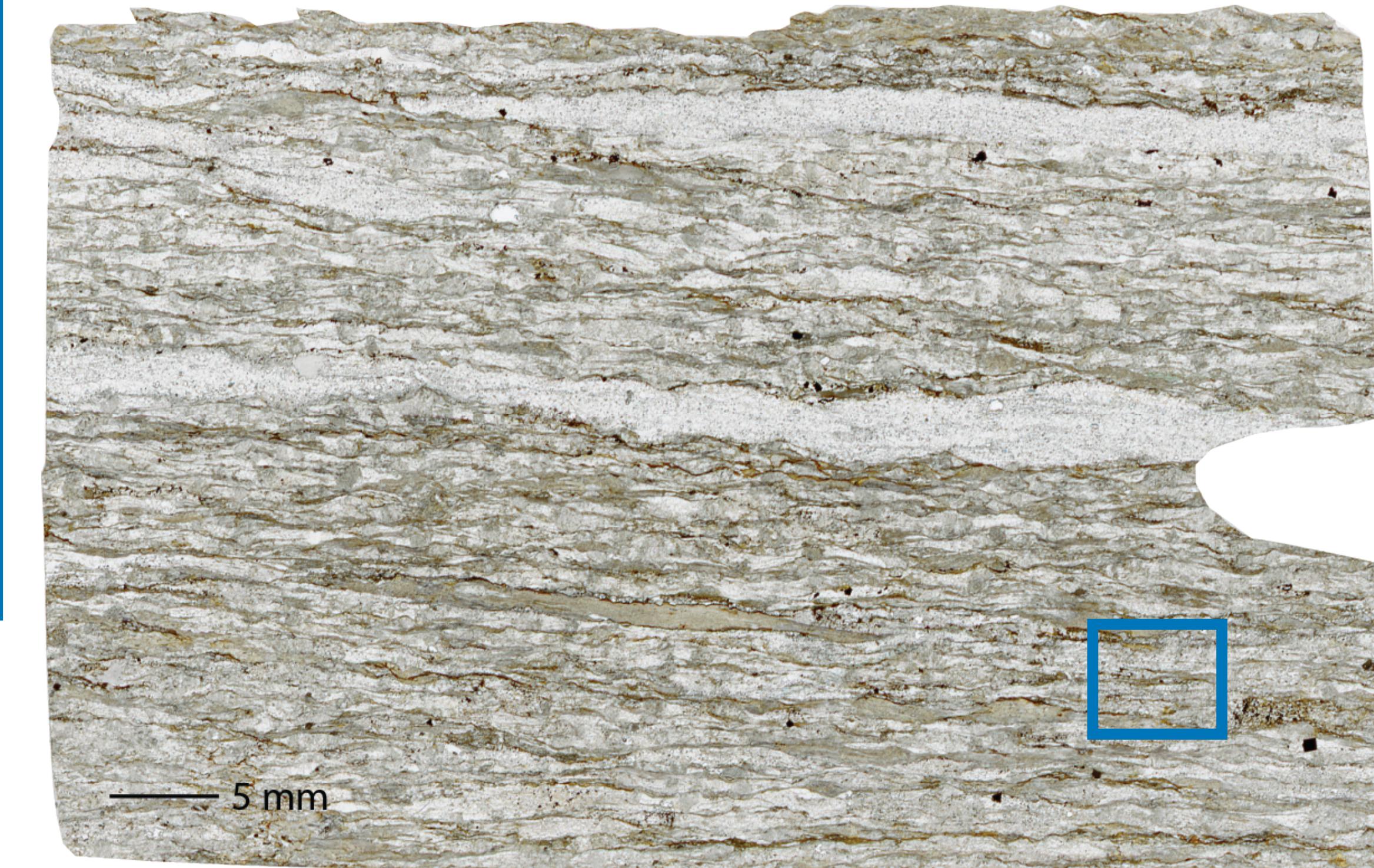
Rootless fold noses = strong blocks
 F_2 Limbs = weak matrix

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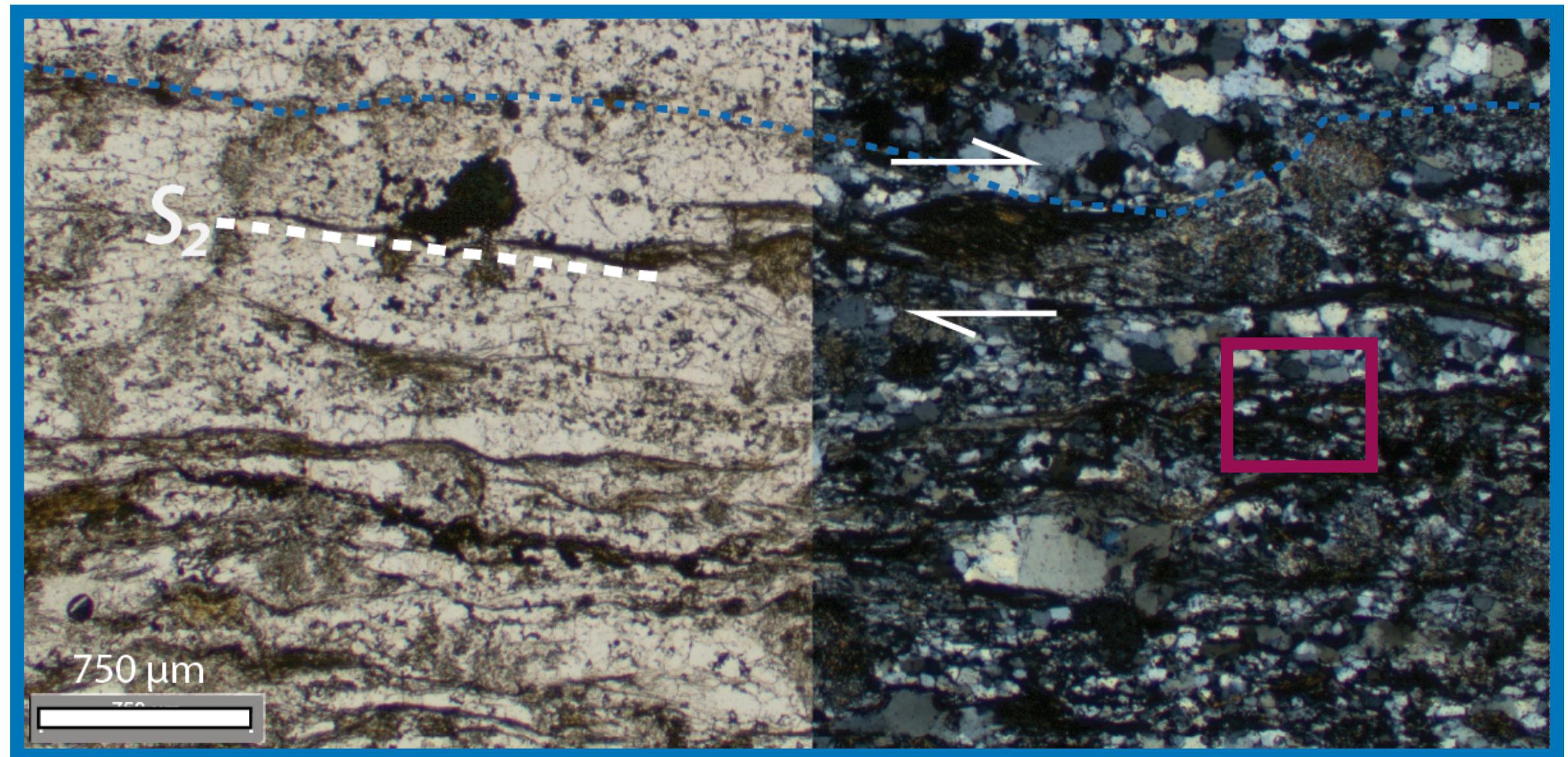


Fold Limb/Matrix

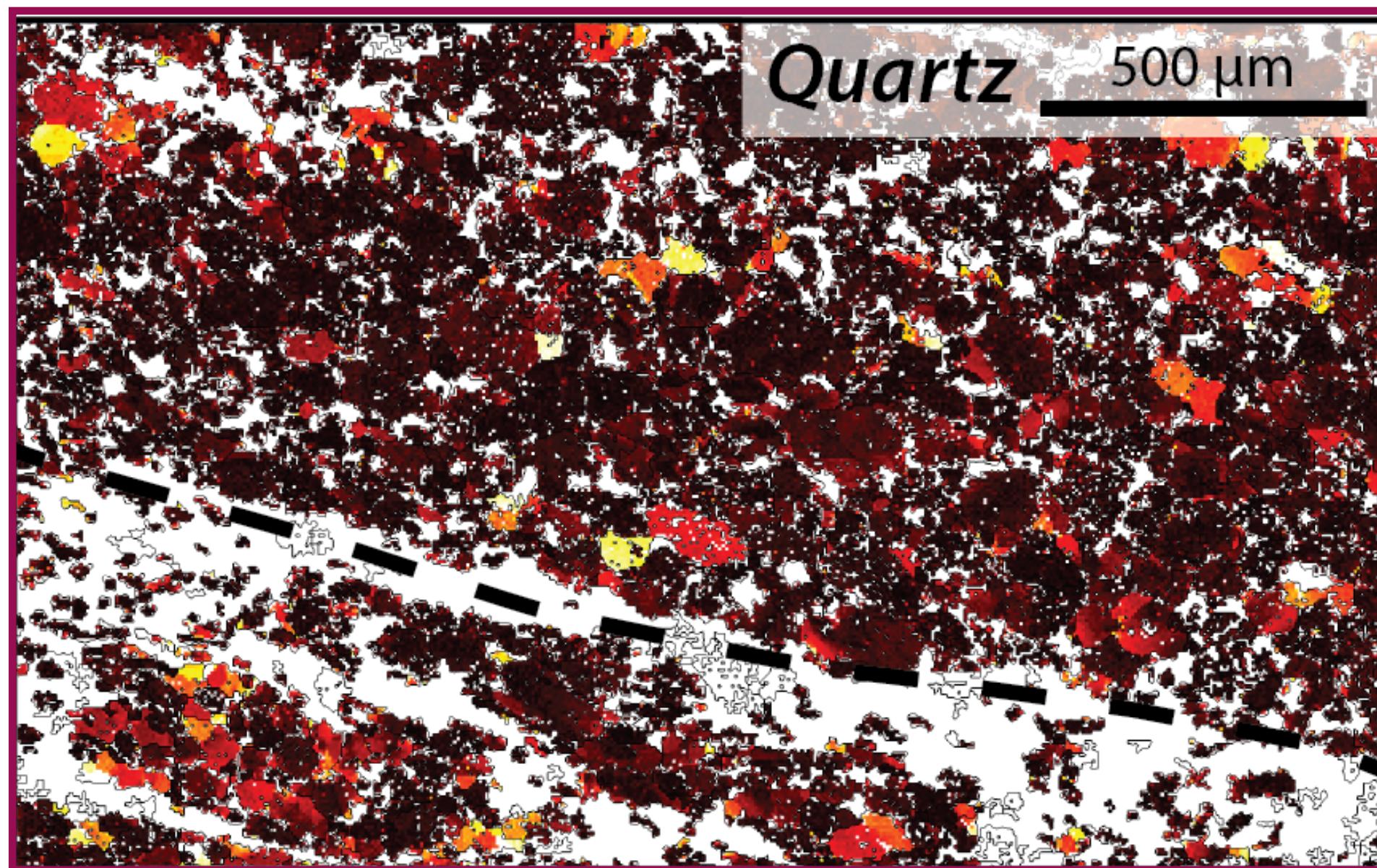
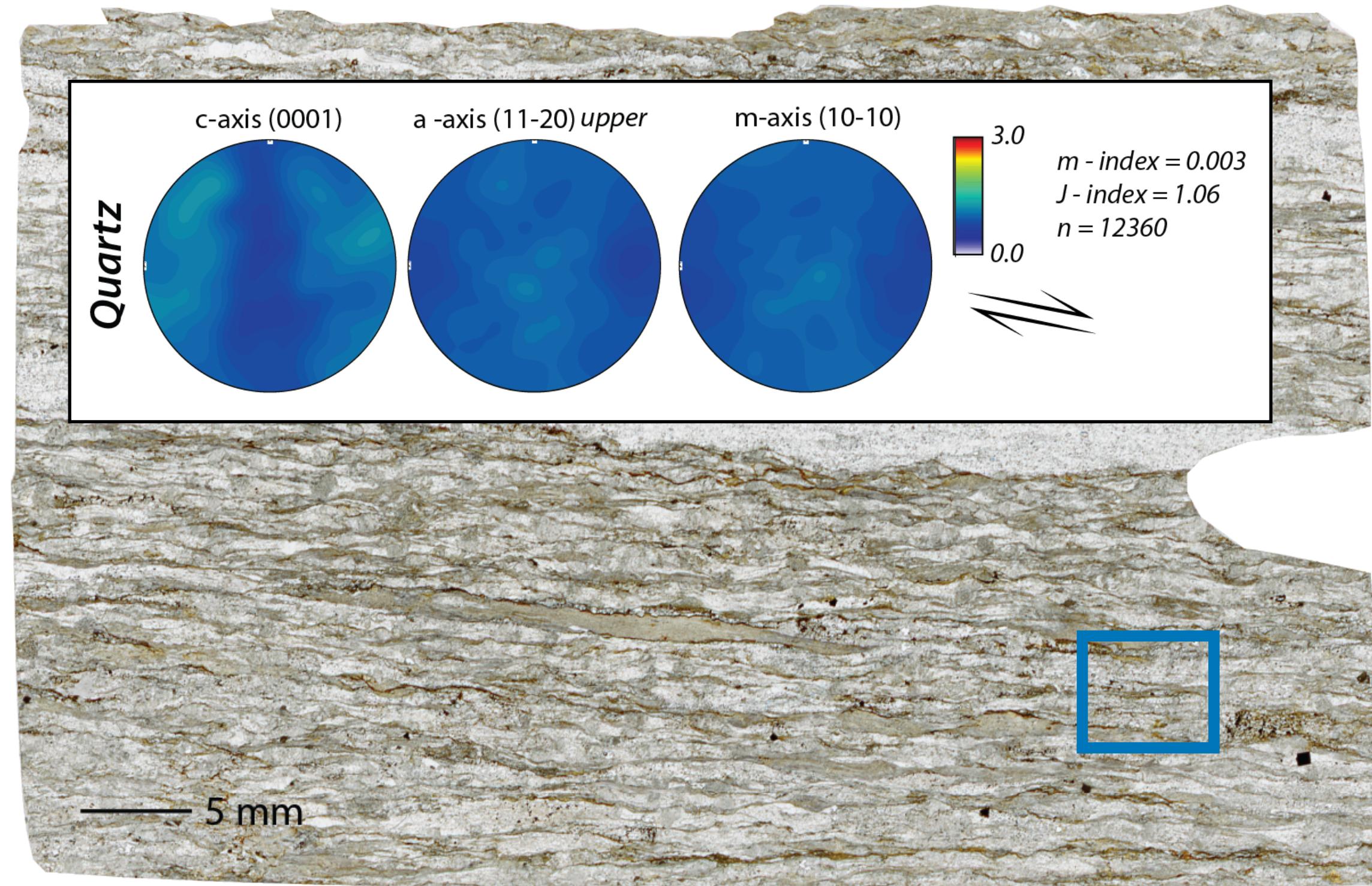


Rootless fold noses = strong blocks
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Matrix fabric - pressure-solution creep
 microstructures - grain size $\sim 15 \mu\text{m}$ in both Ab and Qtz - Fluid mediated process



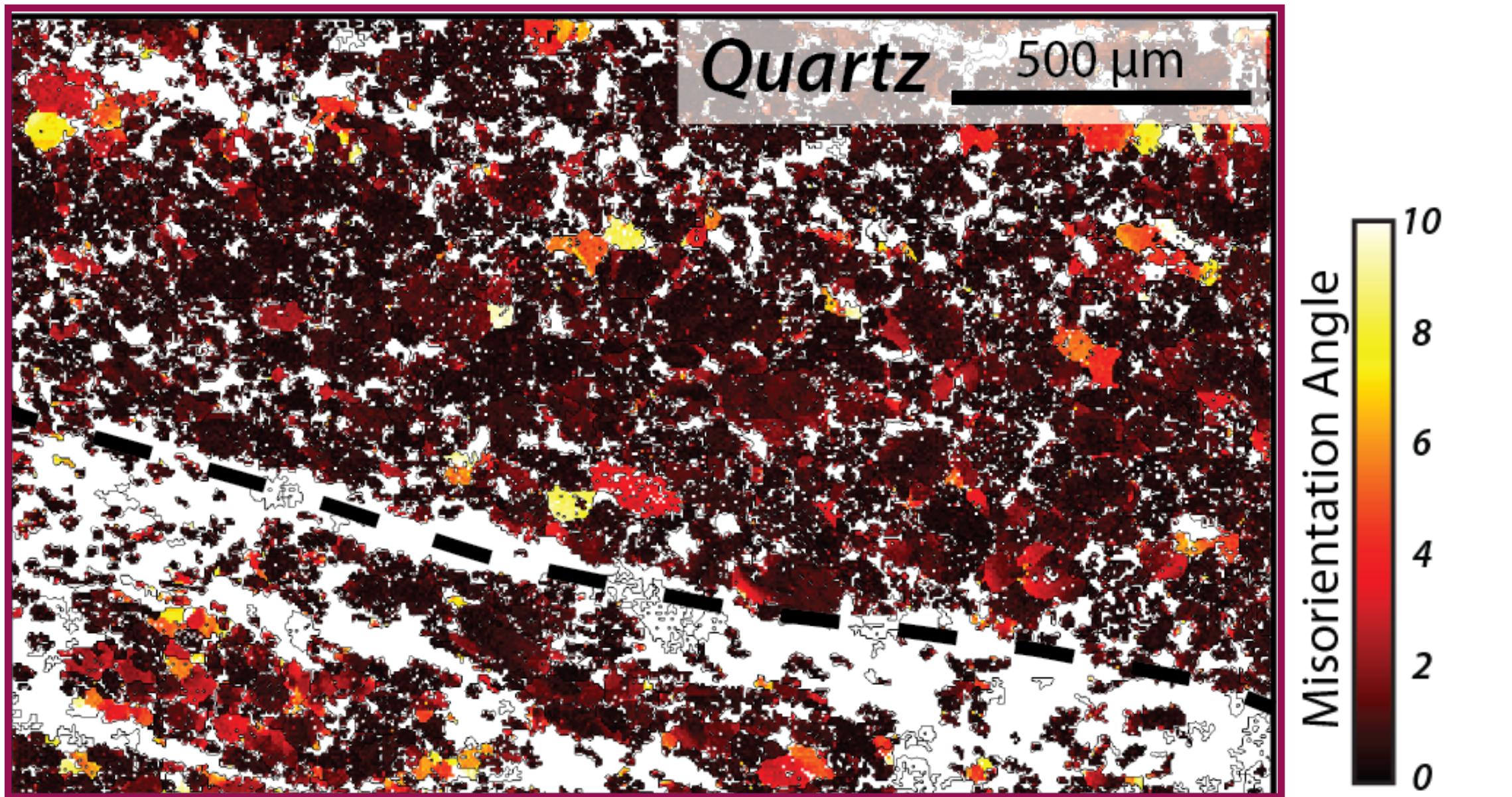
Fold Limb/Matrix



Misorientation Angle

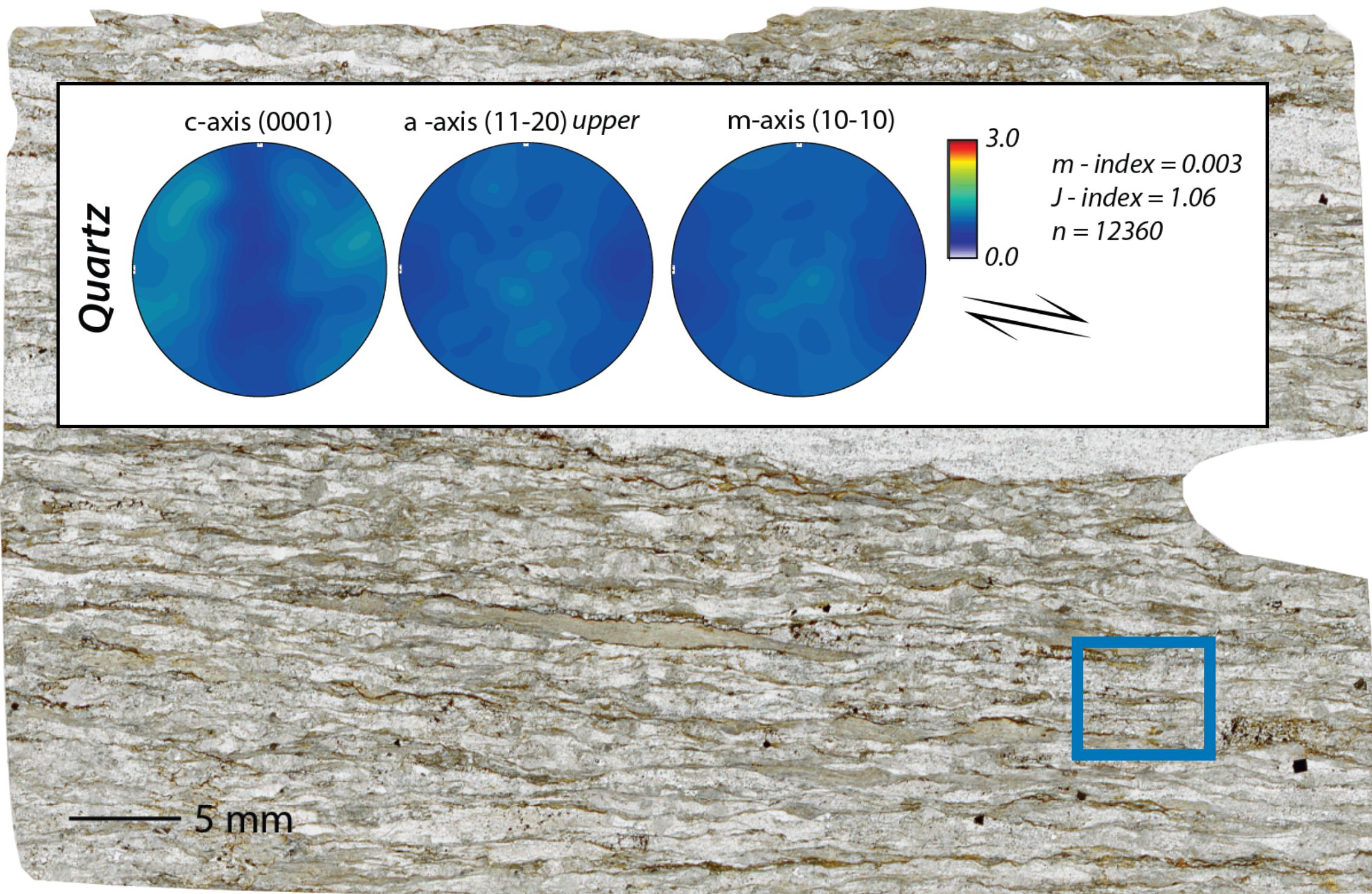
Quartz Mis2Mean EBSD map of matrix fabric

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Quartz Mis2Mean EBSD map of matrix fabric

Fold Limb/Matrix



Matrix fabric - pressure-solution creep
 microstructures - grain size $\sim 15 \mu\text{m}$ in both Ab and Qtz - Fluid mediated process

What slip behaviors do these rocks host?

Thin film model for pressure
solution creep of quartz

$$\dot{\varepsilon} = -\frac{AV_m c D_{gb} w \sigma \rho_f}{RT d^3 \rho_s}$$

Grain size dependent
Stress exponent = 1
Rutter, 1976

*Key area for improved
understanding*

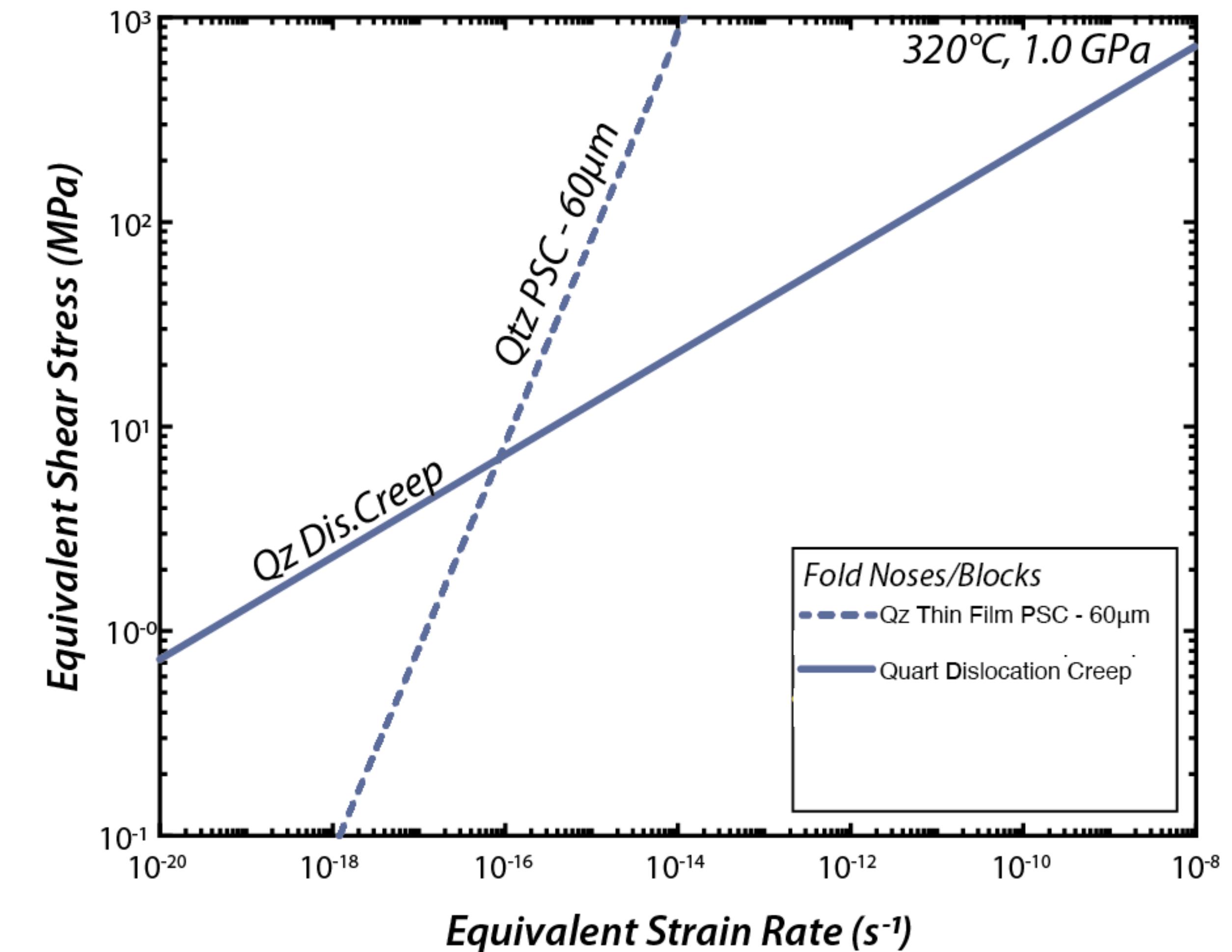
Dislocation creep of quartz

Grain size independent
Stress exponent = 4

$$\dot{\varepsilon} = A f_{H_2O} \sigma^4 \exp\left(-Q/RT\right)$$

Hirth et al., 2001;
Tokle et al., 2019

Rheology of the metasediments



Thin film model for pressure solution creep of quartz

$$\dot{\varepsilon} = \frac{AV_m c D_{gb} w \sigma \rho_f}{RT d^3 \rho_s}$$

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Key area for improved understanding

Dislocation creep of quartz

Grain size independent

Stress exponent = 4

$$\dot{\varepsilon} = A f_{H_2O} \sigma^4 \exp\left(-Q/RT\right)$$

Hirth et al., 2001;
Tokle et al., 2019

Diffusion Creep of albite

$$\dot{\varepsilon} = A \sigma d^{-3} \exp\left(\frac{-Q}{RT}\right)$$

Grain size dependent
Stress exponent = 1

Offerhaus, 2001

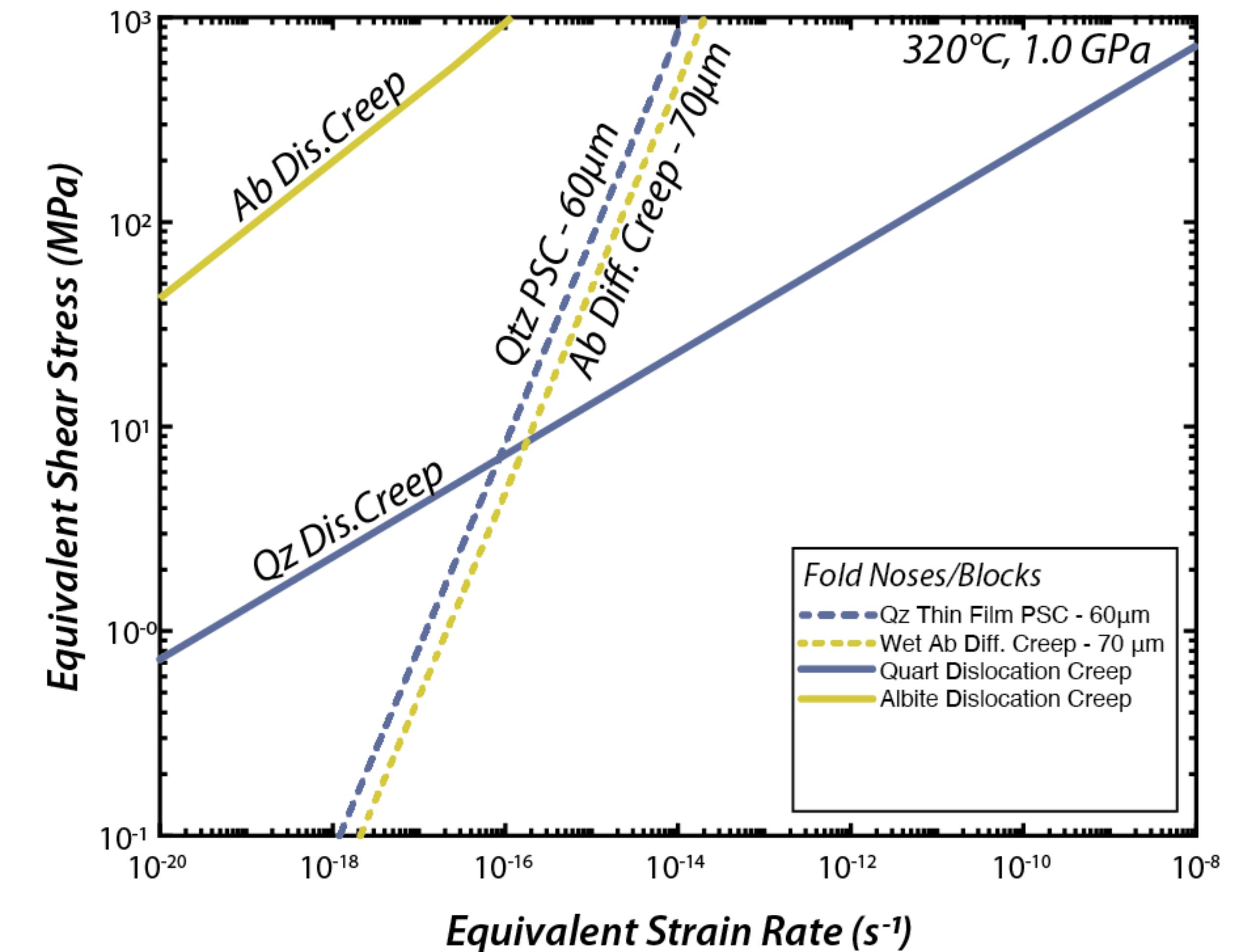
Key area for improved understanding

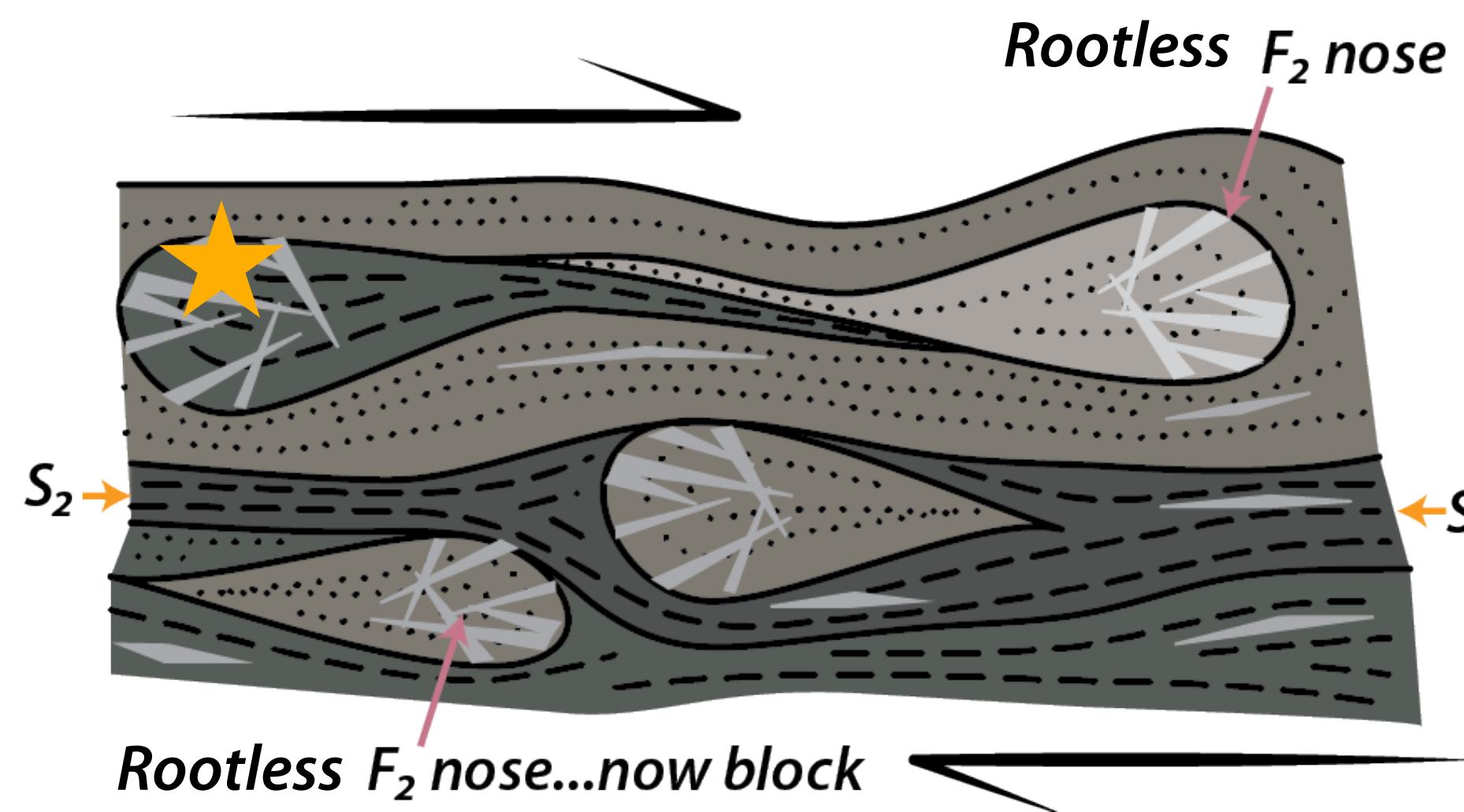
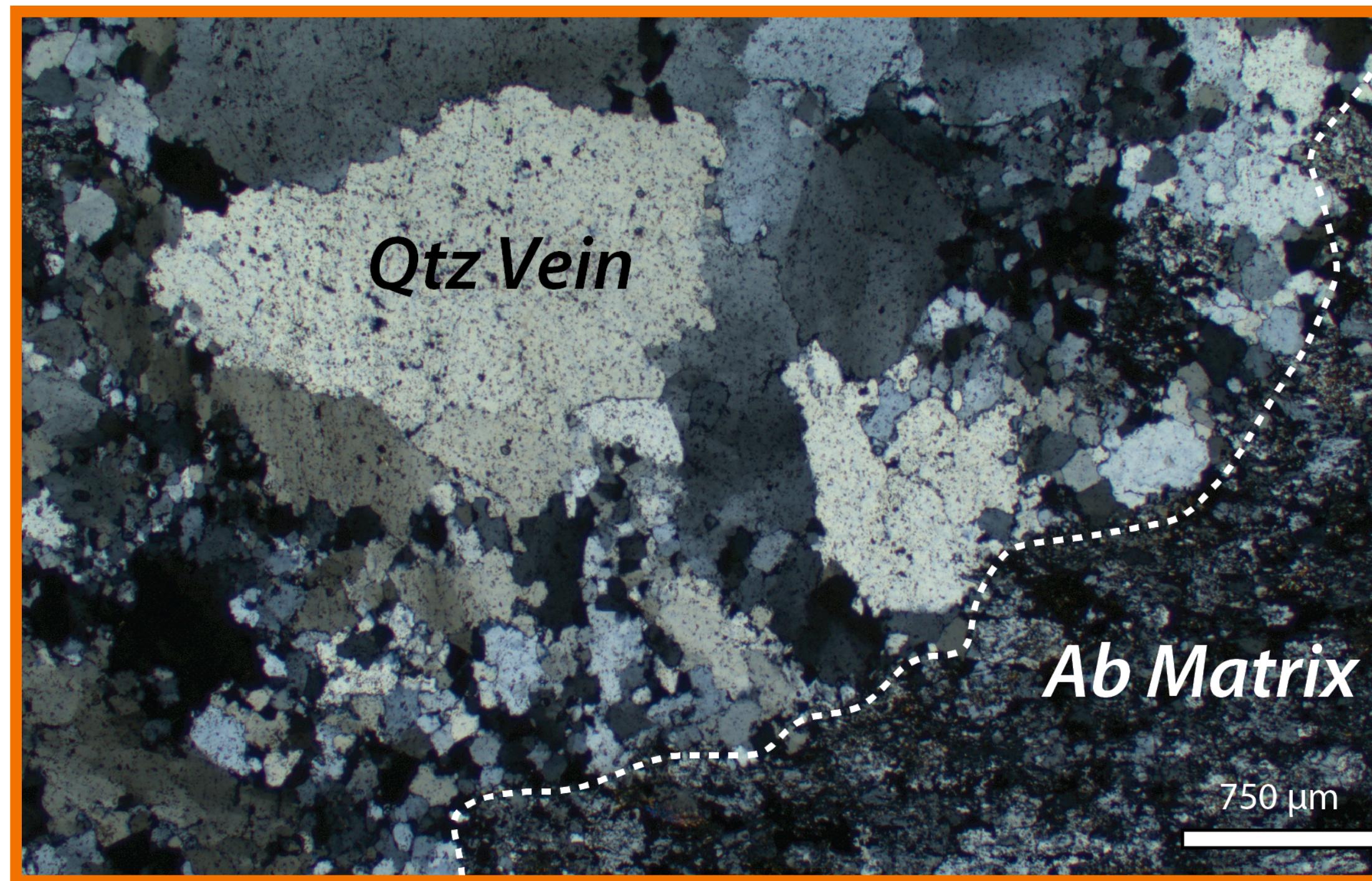
Dislocation Creep of feldspar

Grain size independent
Stress exponent = 3

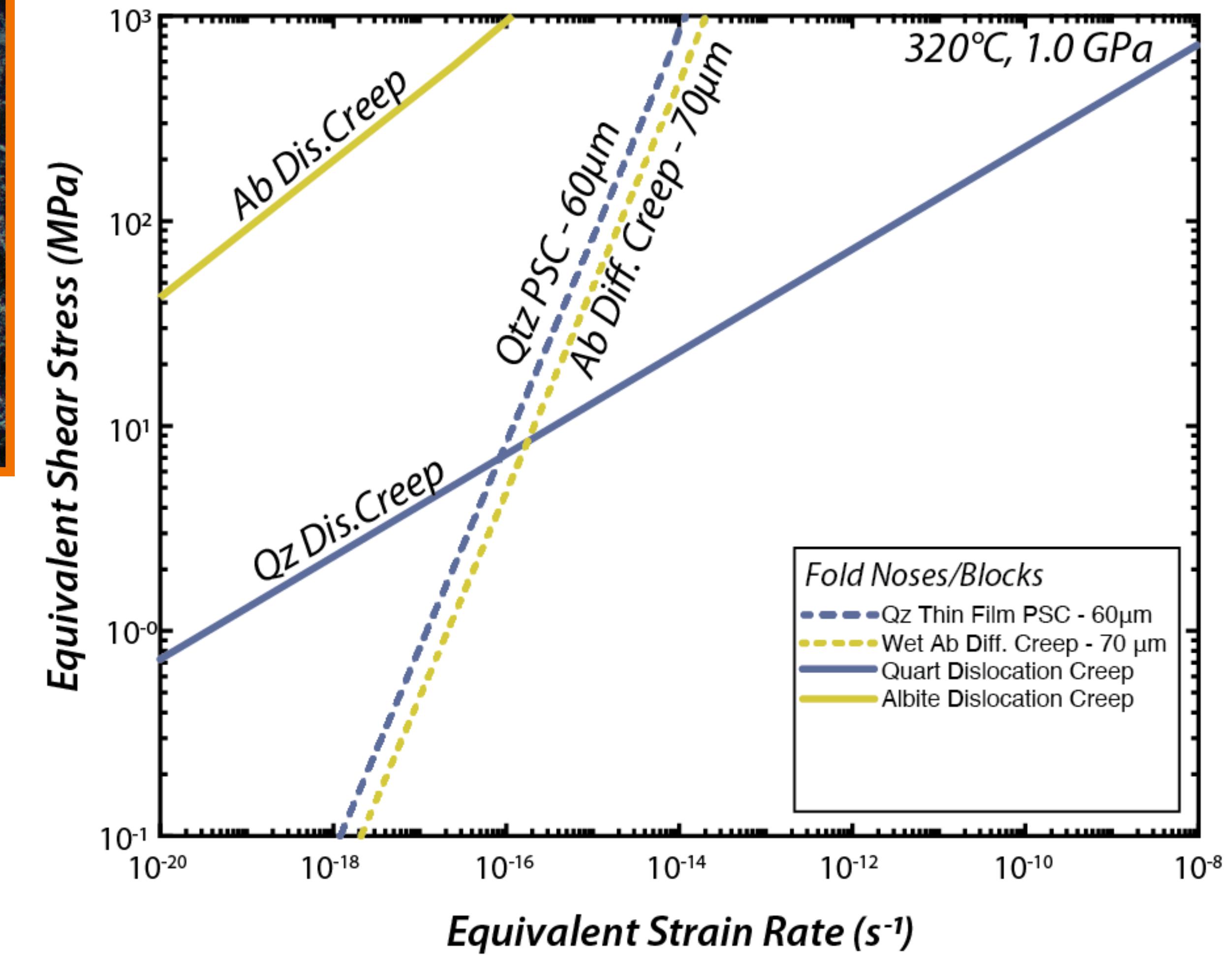
Rybacki & Dresen, 2000

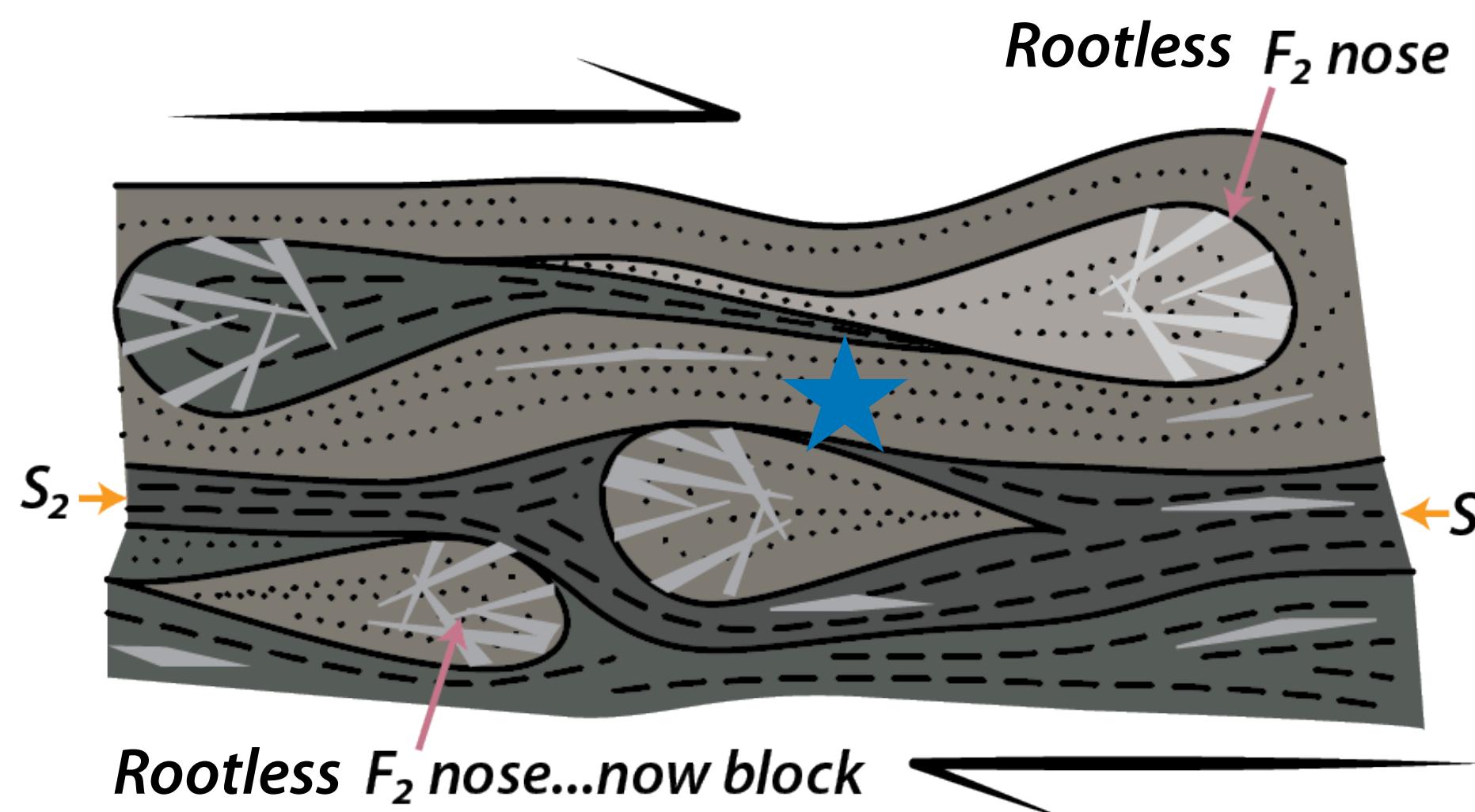
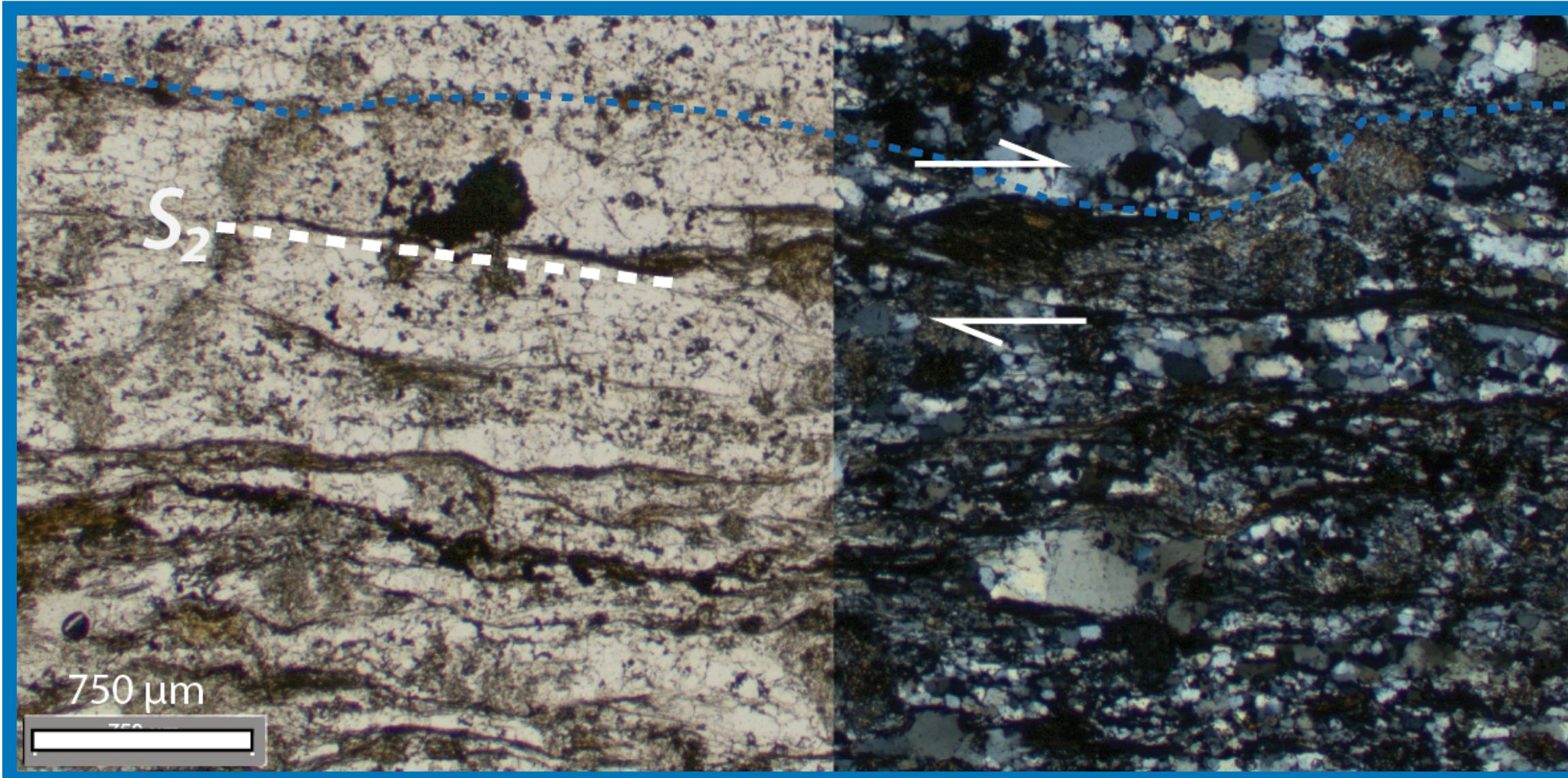
Rheology of the metasediments



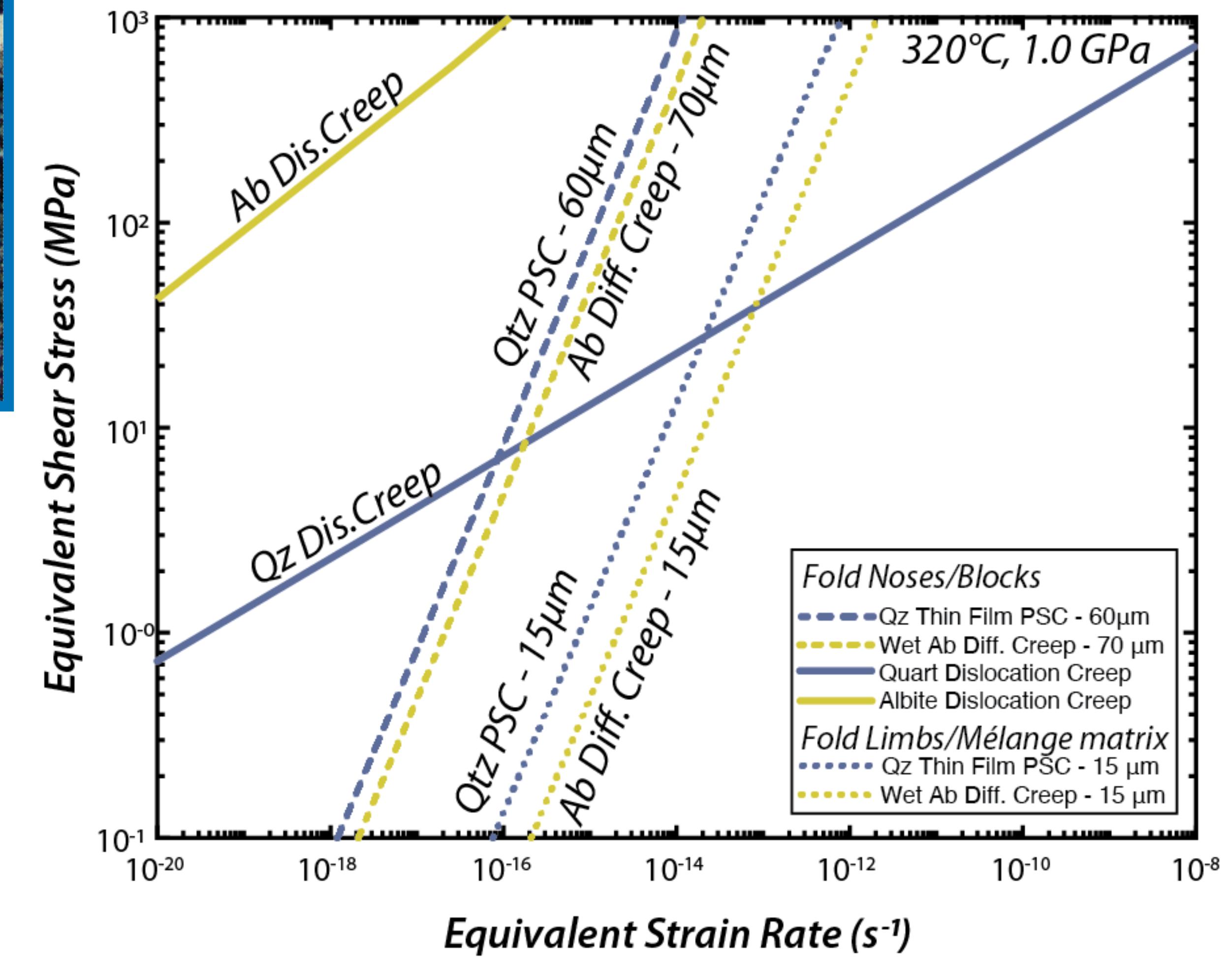


Rheology of the metasediments



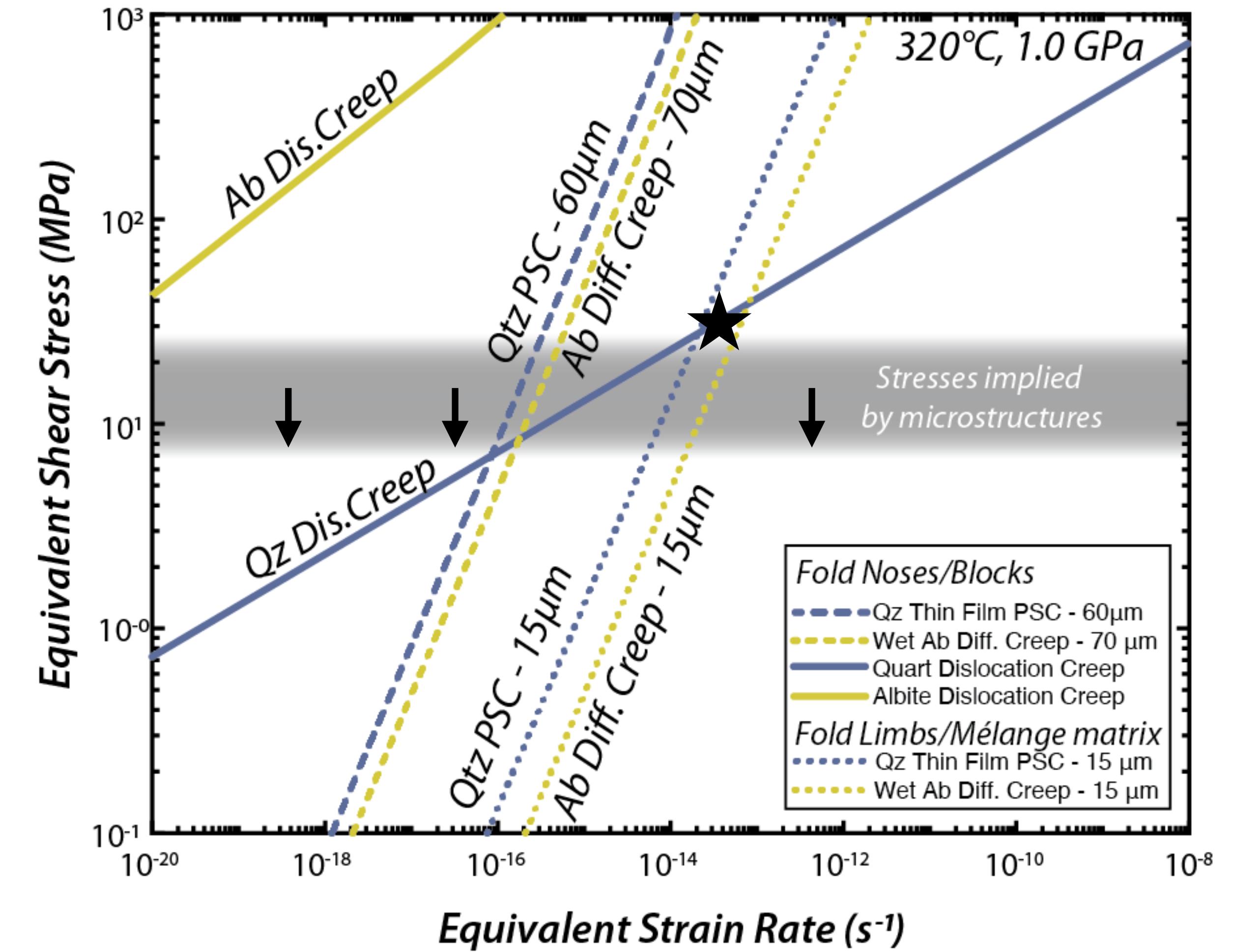
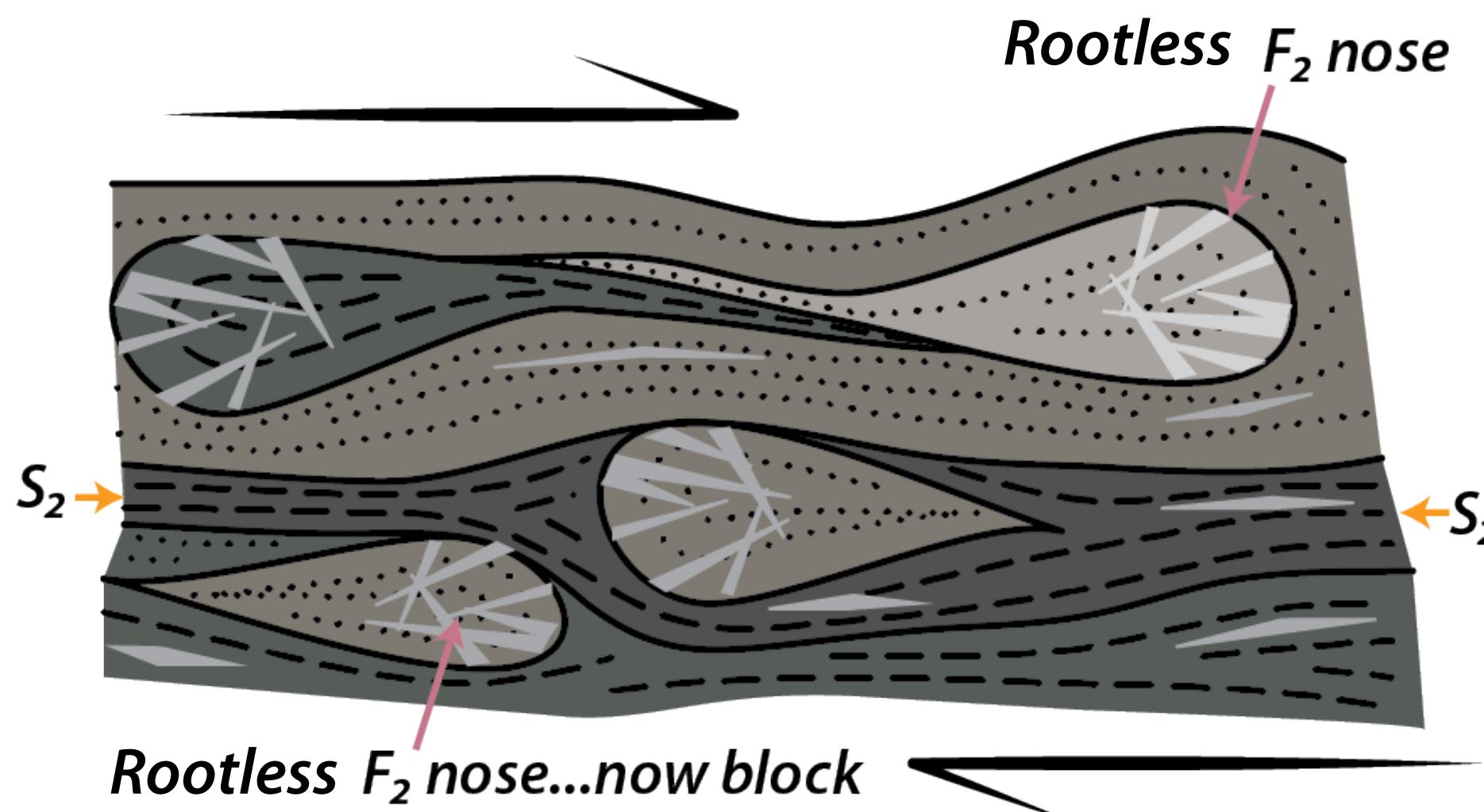


Rheology of the metasediments



Requires $\sigma < \sim 25$ MPa for PSC of matrix to be weaker than dislocation creep of the quartz veins

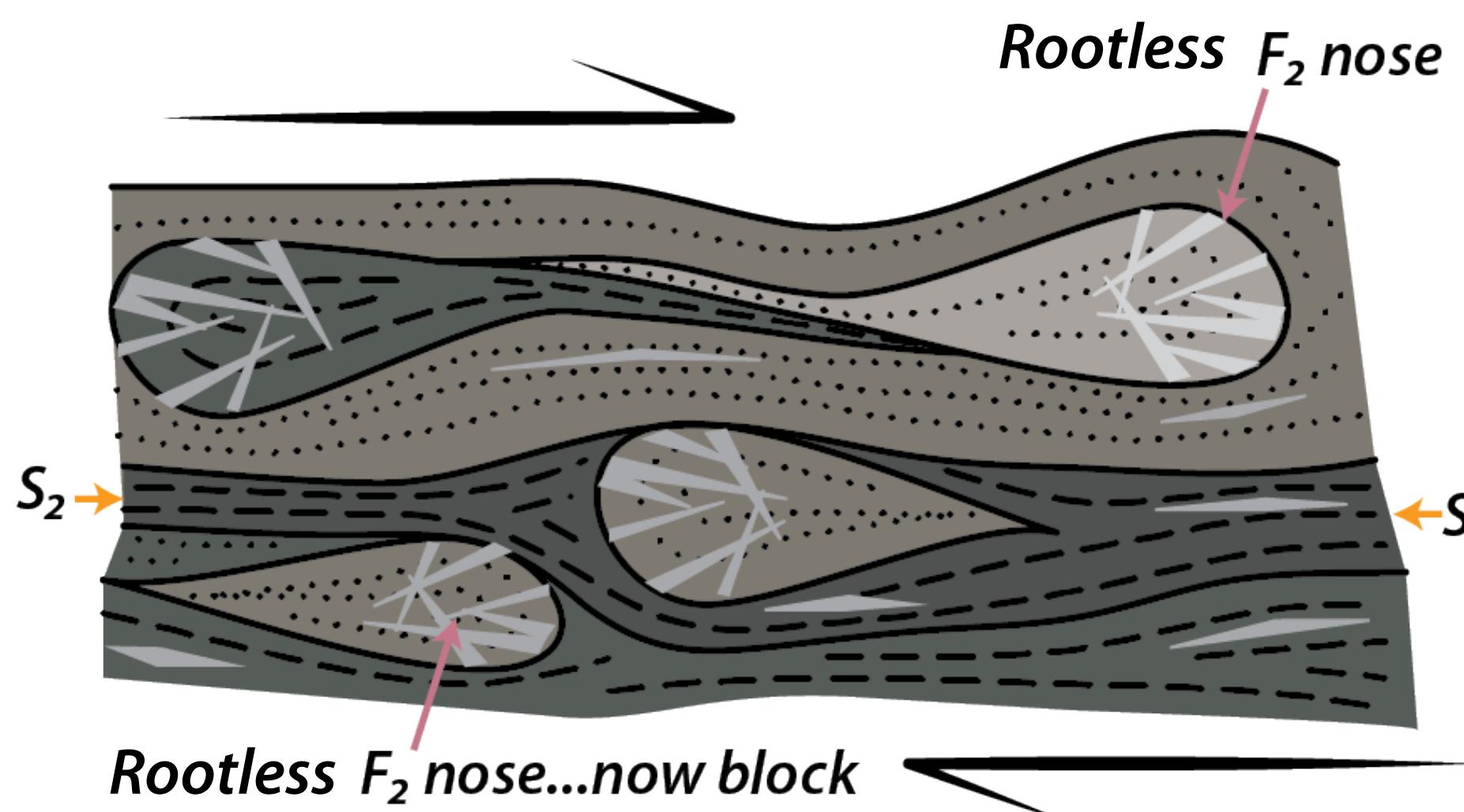
Rheology of the metasediments



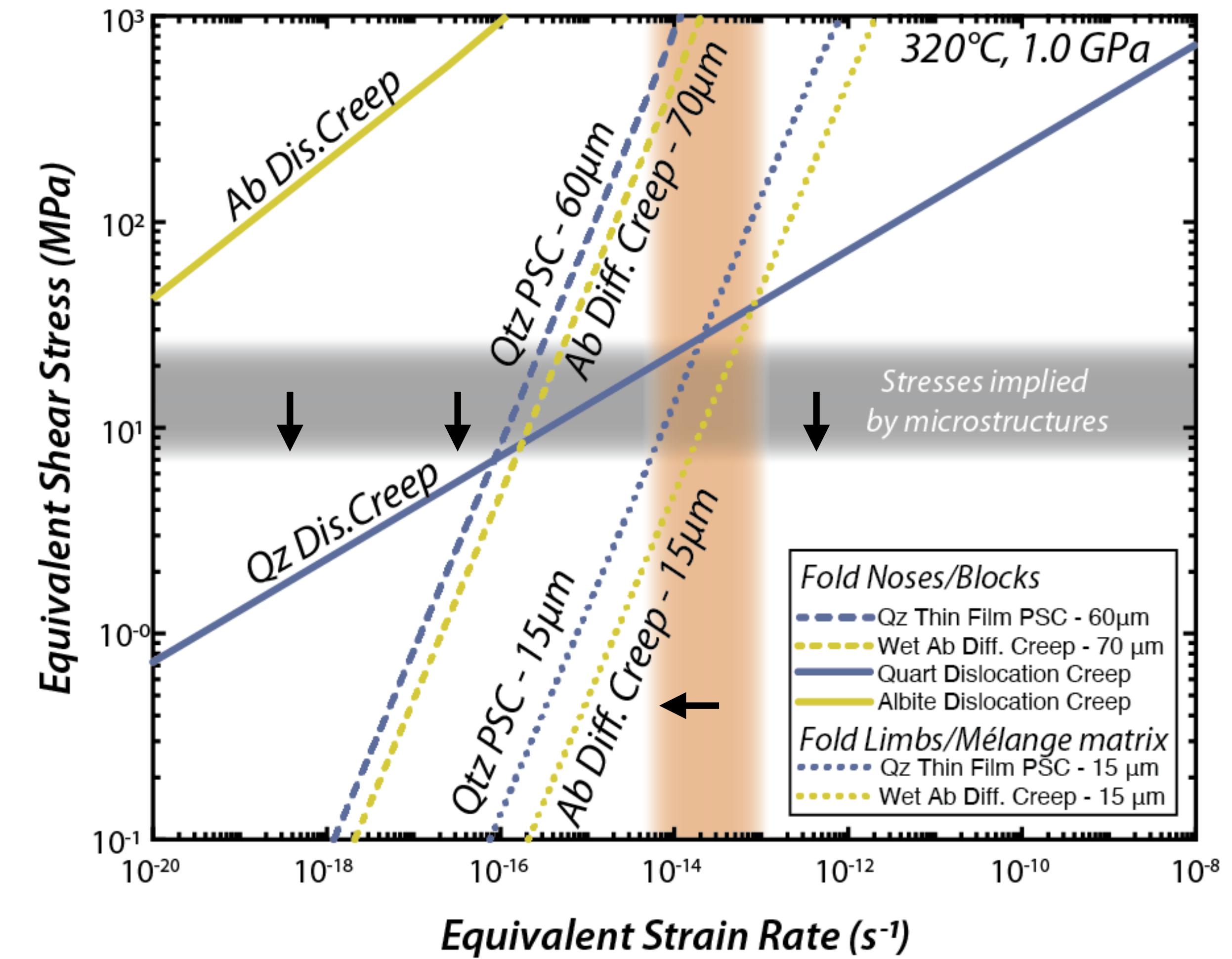
Requires $\sigma < \sim 25$ MPa for PSC of matrix to be weaker than dislocation creep of the quartz veins

Results in strain rates 10^{-13} s $^{-1}$ or slower

With reasonable thicknesses (200-2000m): tectonic strain rates



Rheology of the metasediments

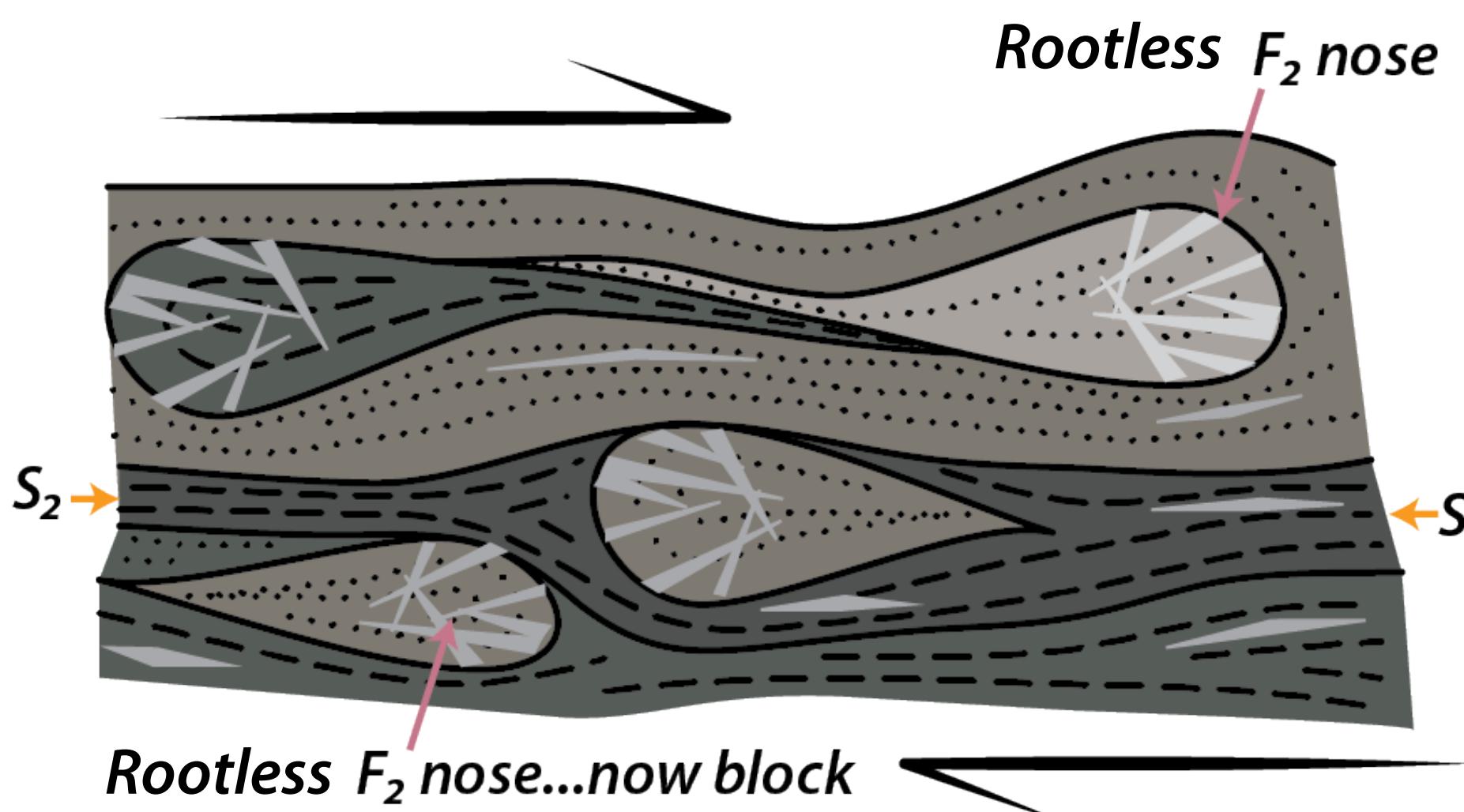


Requires $\sigma > \sim 25$ MPa for PSC of matrix to be weaker than dislocation creep of the quartz veins

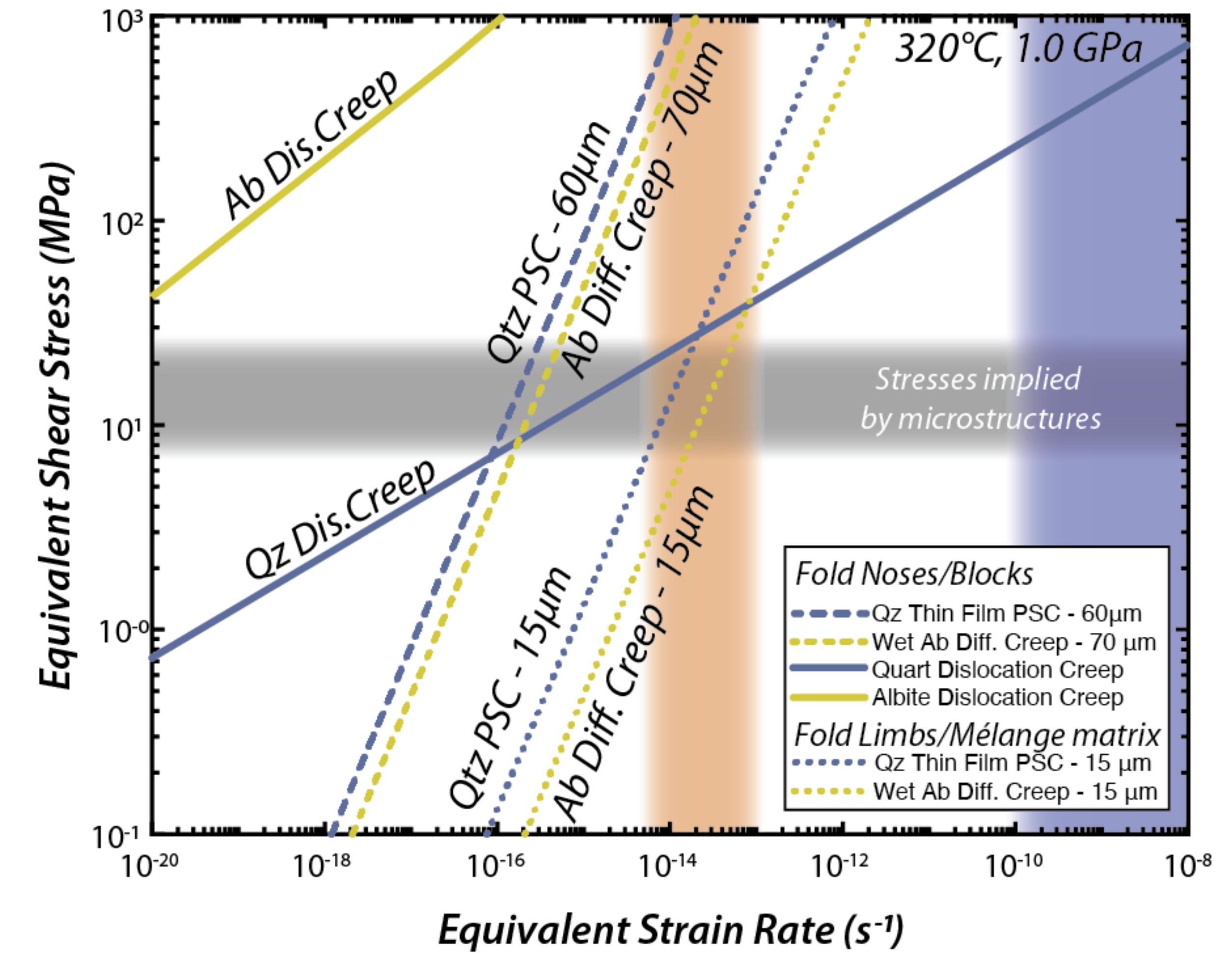
Results in strain rates 10^{-12} s^{-1} or slower

With reasonable thicknesses (200-2000m): tectonic strain rates

Slow slip rates of 10^{-10} s^{-1} or faster



Rheology of the metasediments



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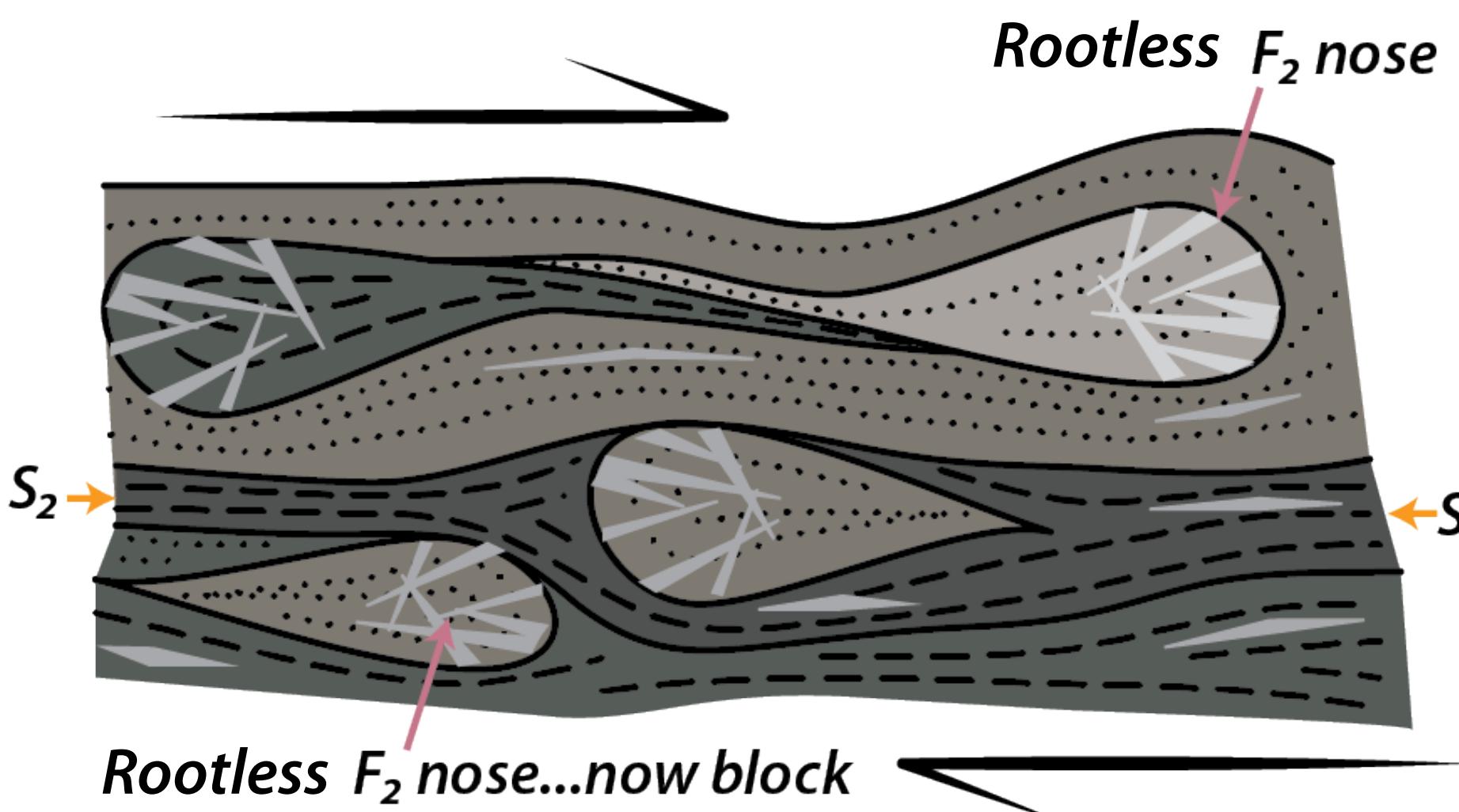
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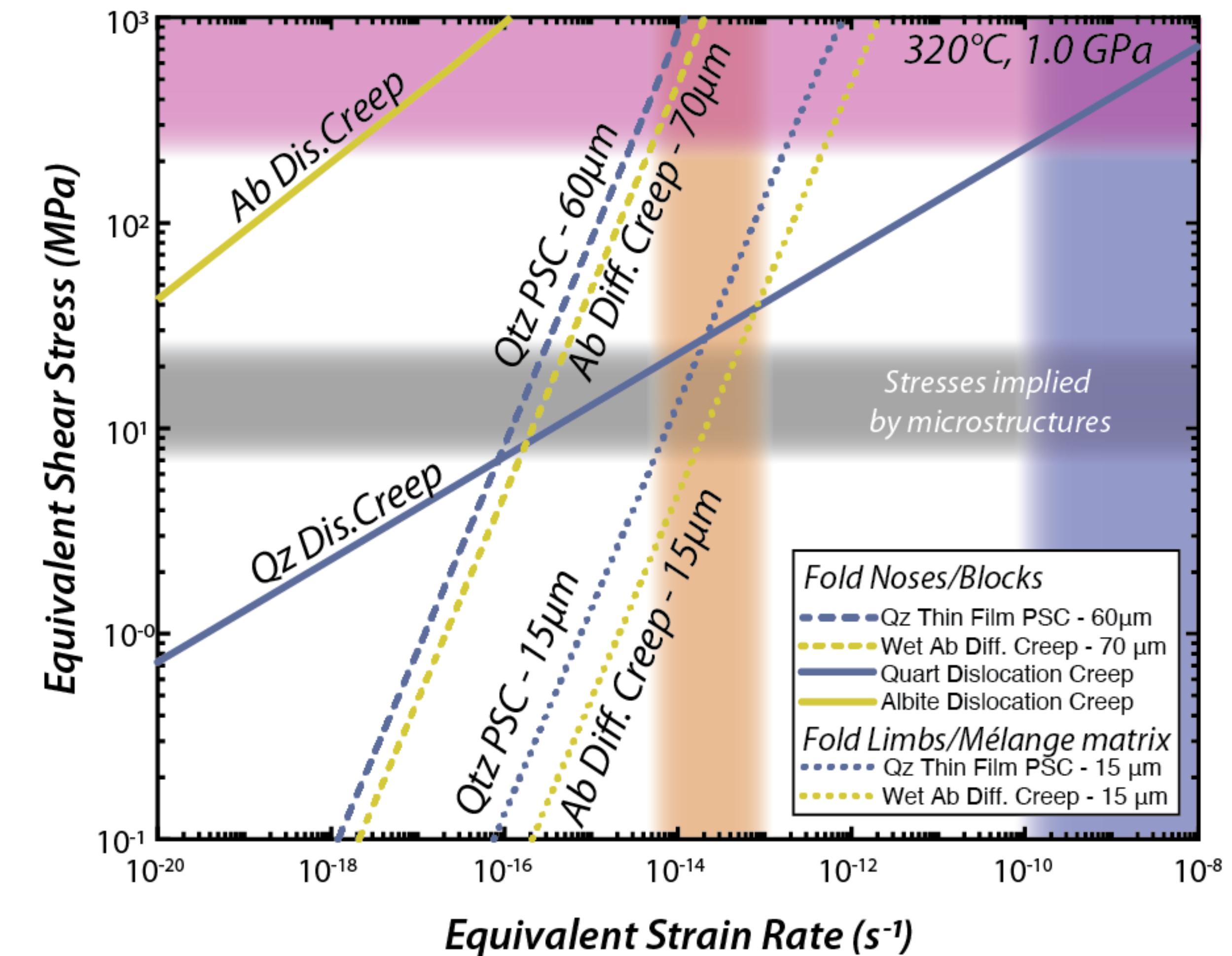
Stresses needed for slow slip strain rates:

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>1000 MPa for thin film pressure solution creep



Rheology of the metasediments



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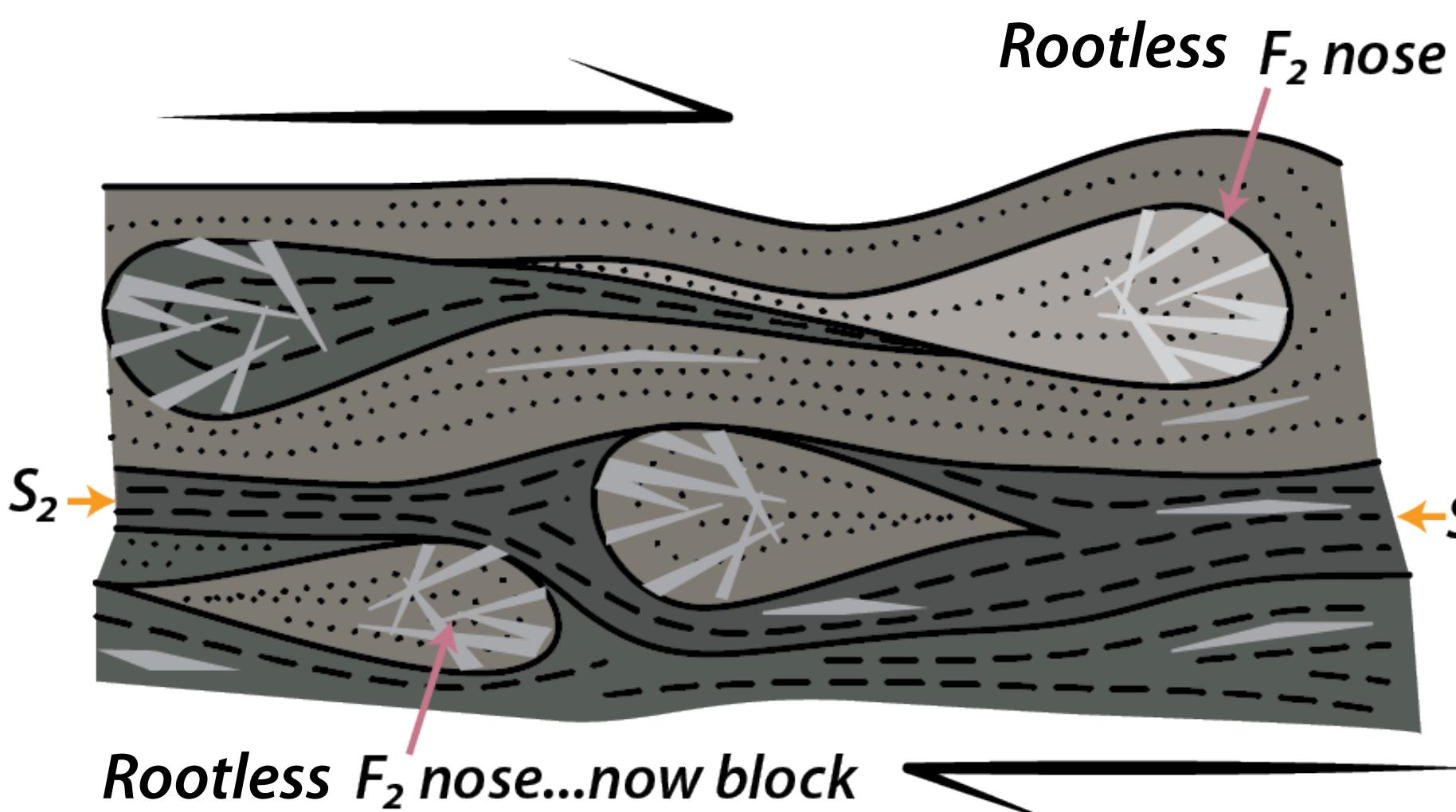
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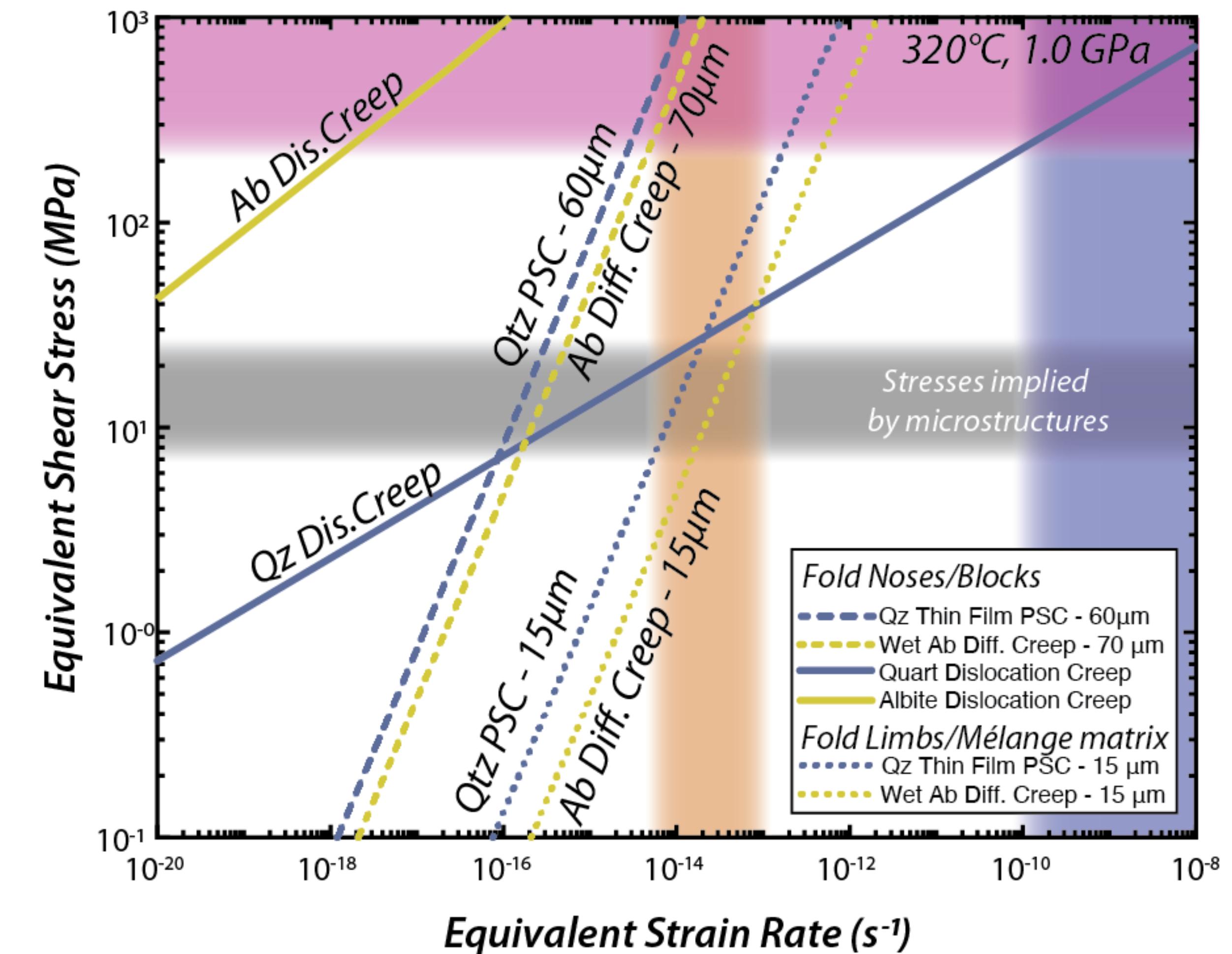
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Rheology of the metasediments



Host tectonic creeping; not slow slip

Observations from the exhumed rock record from the base of the subduction seismogenic zone

What are the phenomena we see in rocks that we need improved:

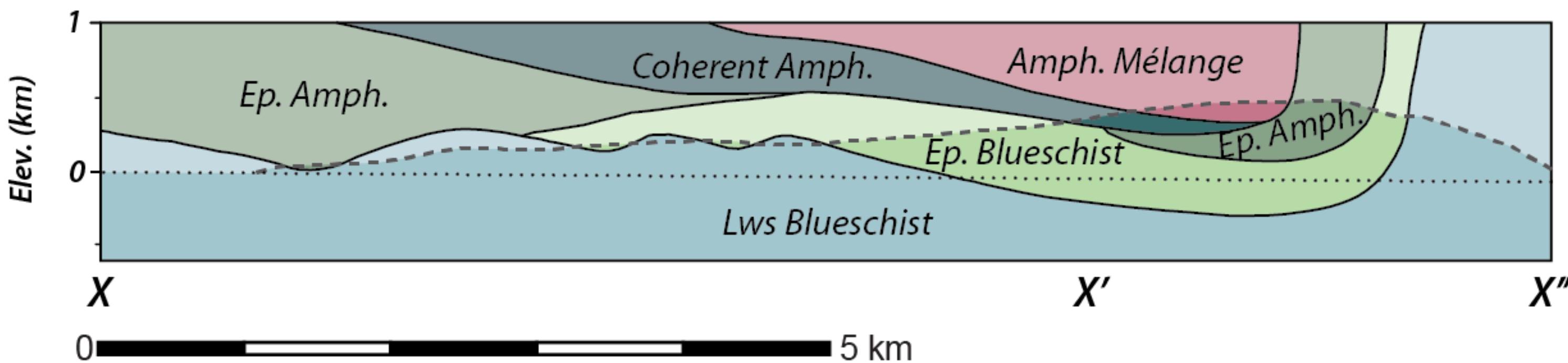
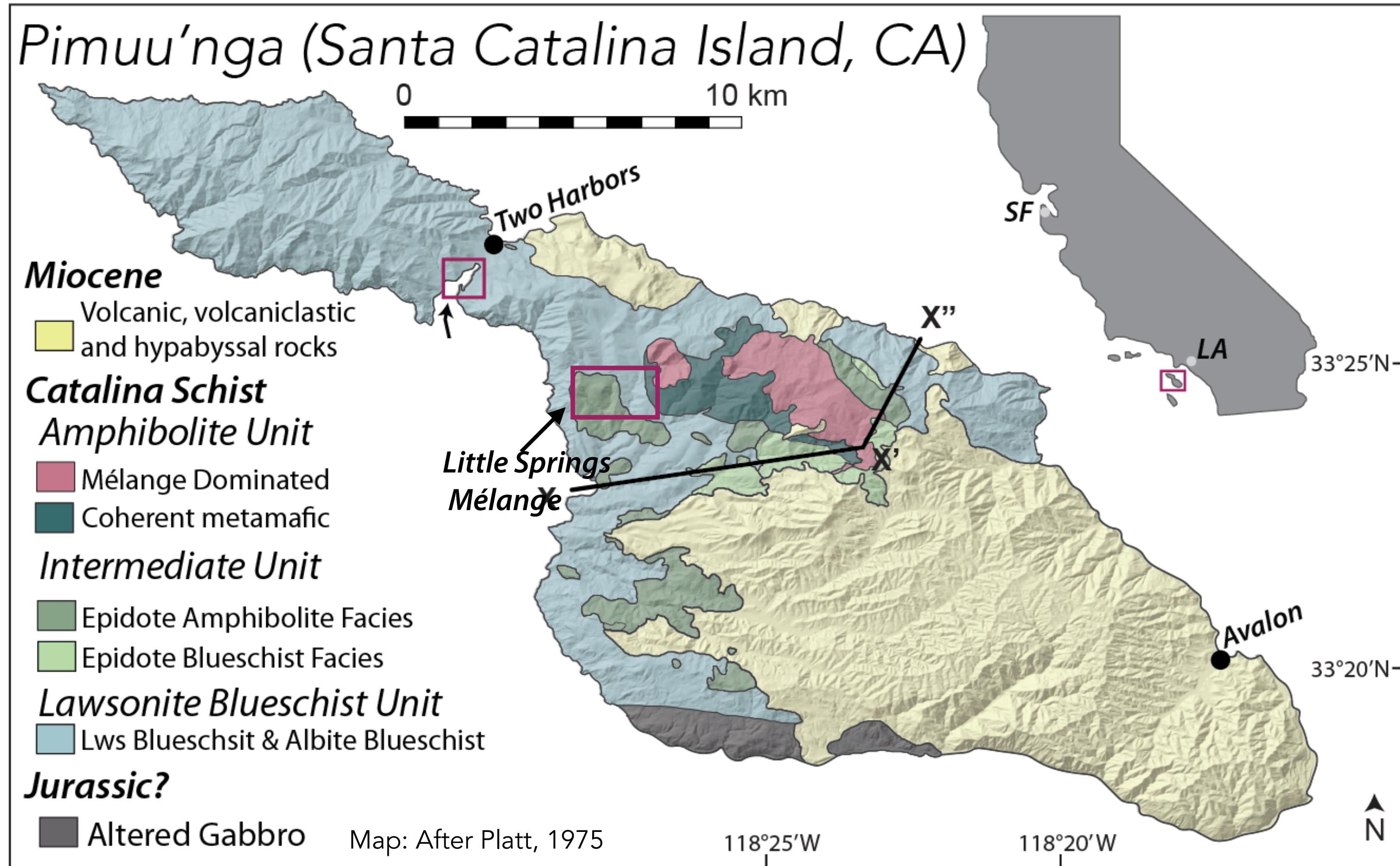
...Constitutive relations for?

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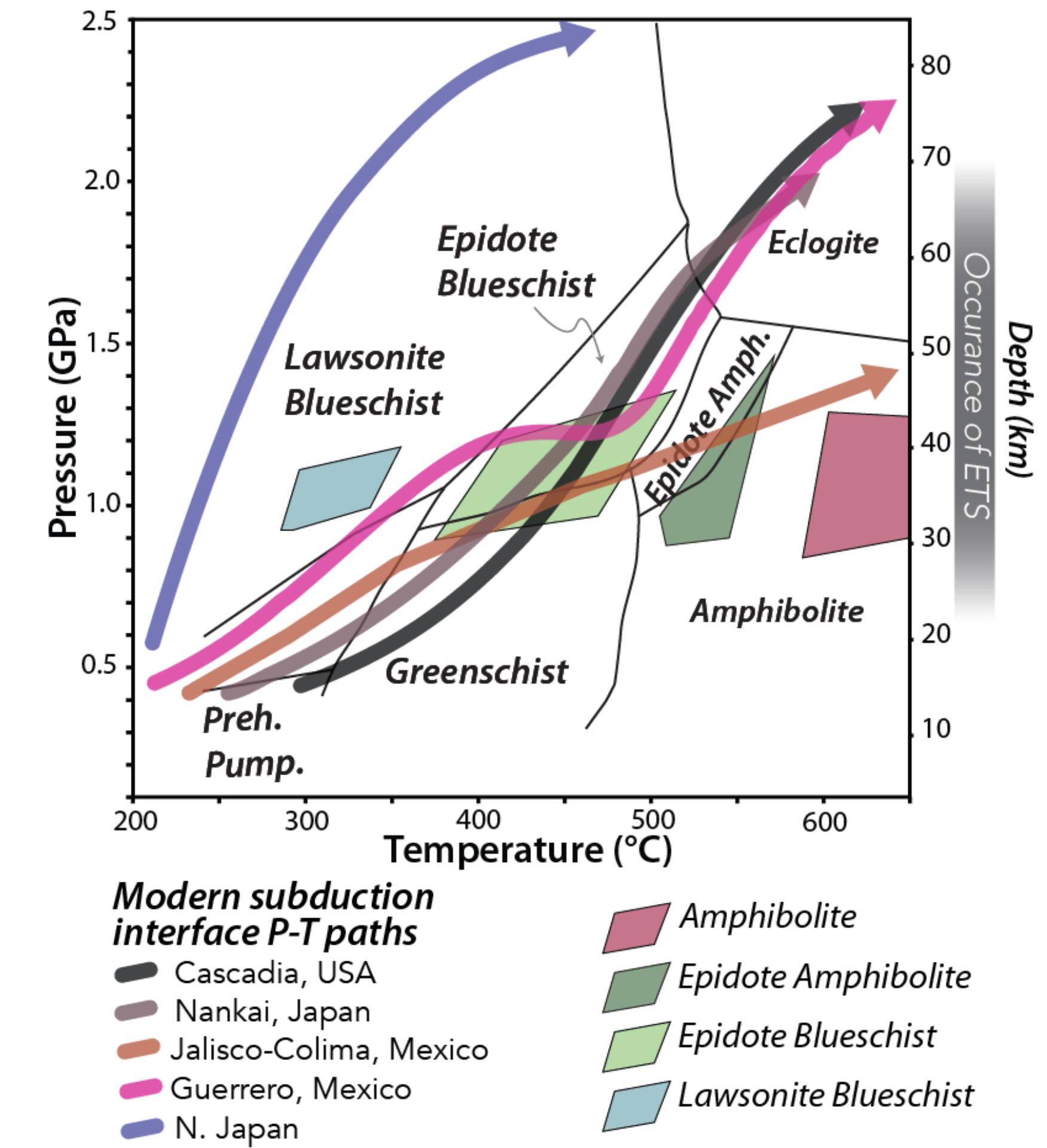
Pressure Solution Creep along the plate interface is not slow slip

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Evidence for brittle deformation (**seismic slip?**) at these deep depths intimately related to fluid process



Cross-section: Platt and Schmidt, 2024

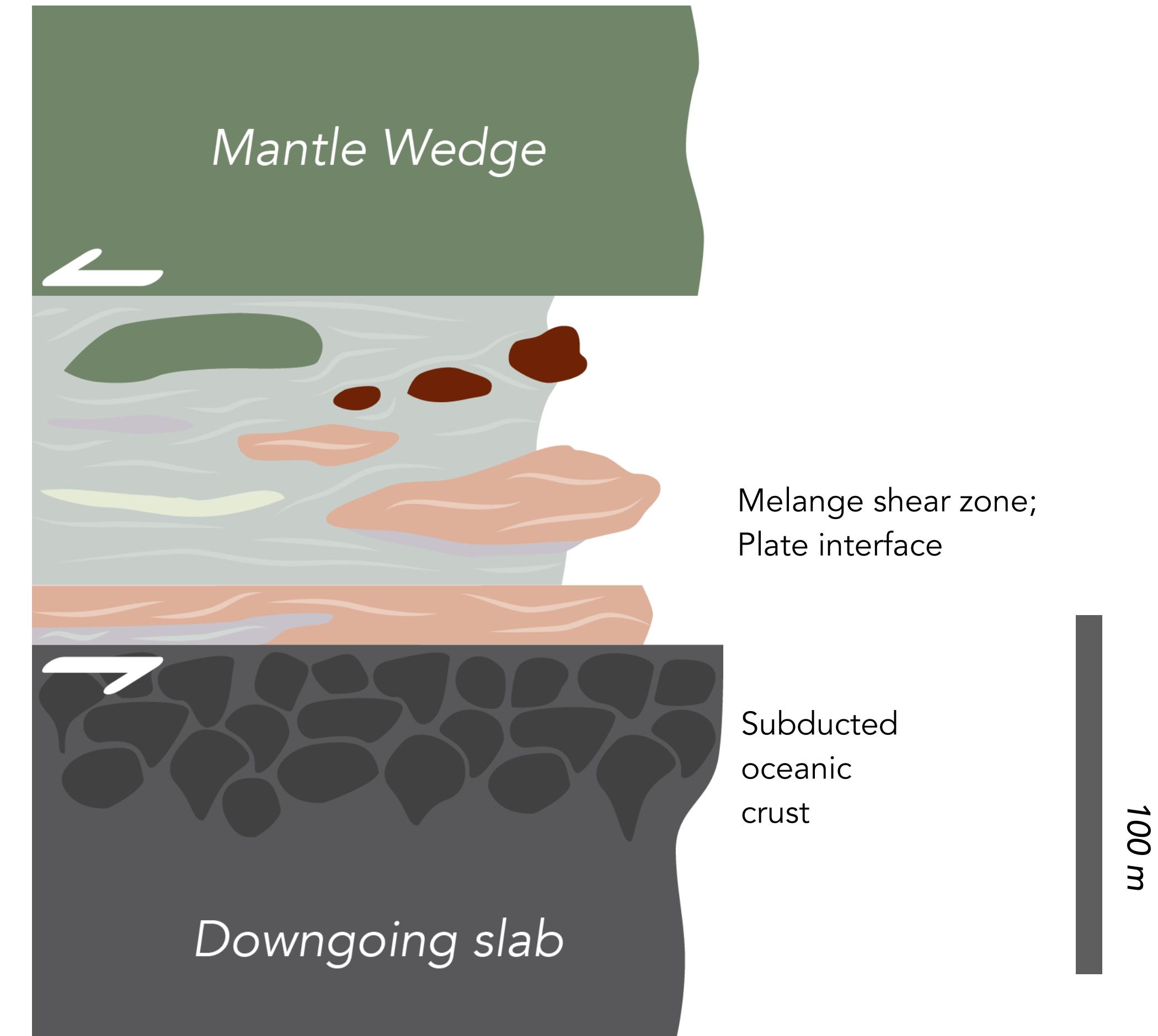
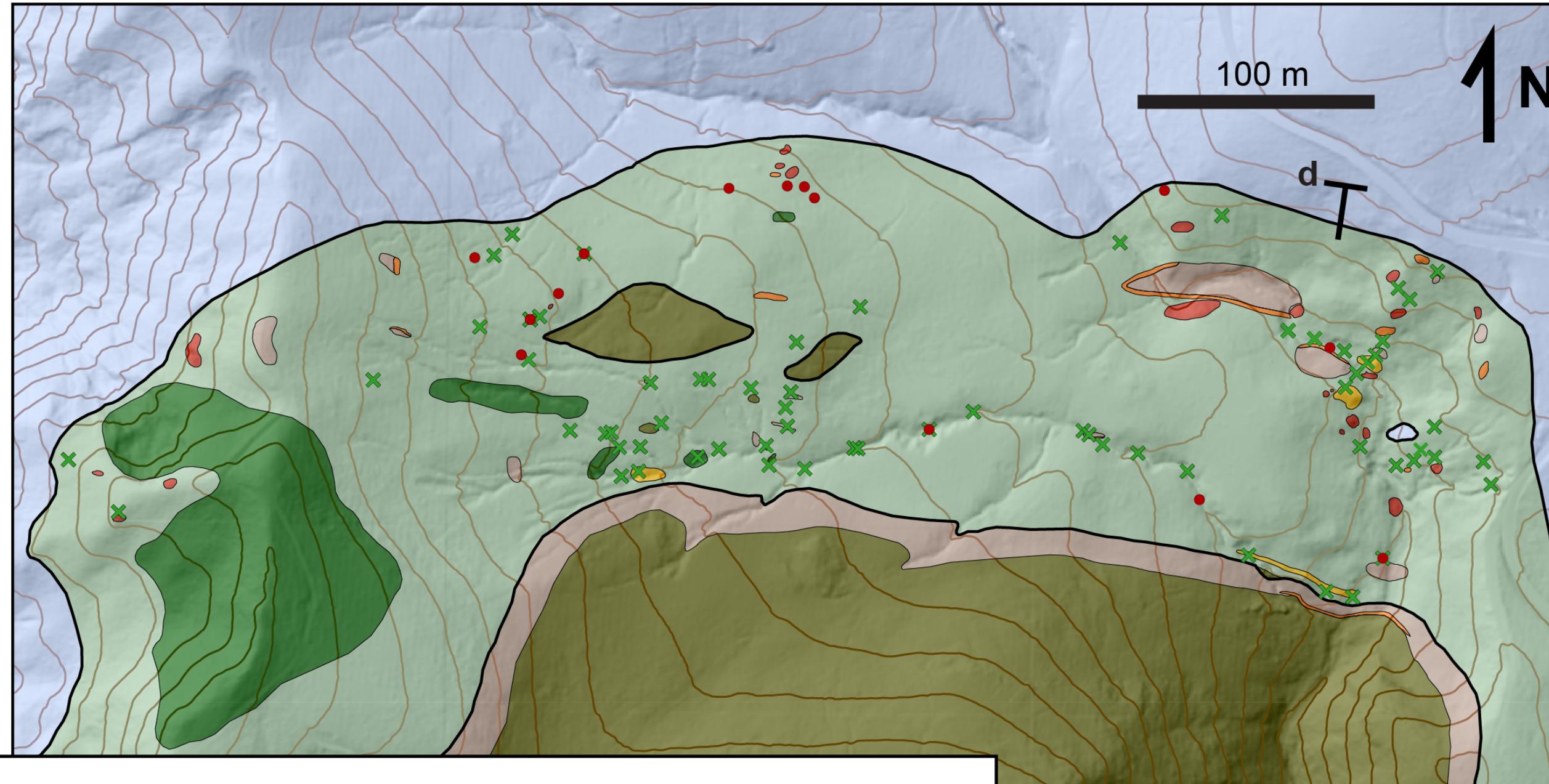


Epidote Amphibolite Unit - ~1.2 GPa (35-45) and 450-550 °C

P-T conditions: Hoover, Condit et al., in Revision G3

Metasomatic mélange, talc, chlorite, and actinolite schists, Serpentinized mantle blocks, metasedimentary rocks
Block-in-matrix structure underlain by lithostratigraphically intact metacherts and metabasalts from the slab

Epidote Amphibolite Facies Plate interface: 450-550°C, 1.2 GPa

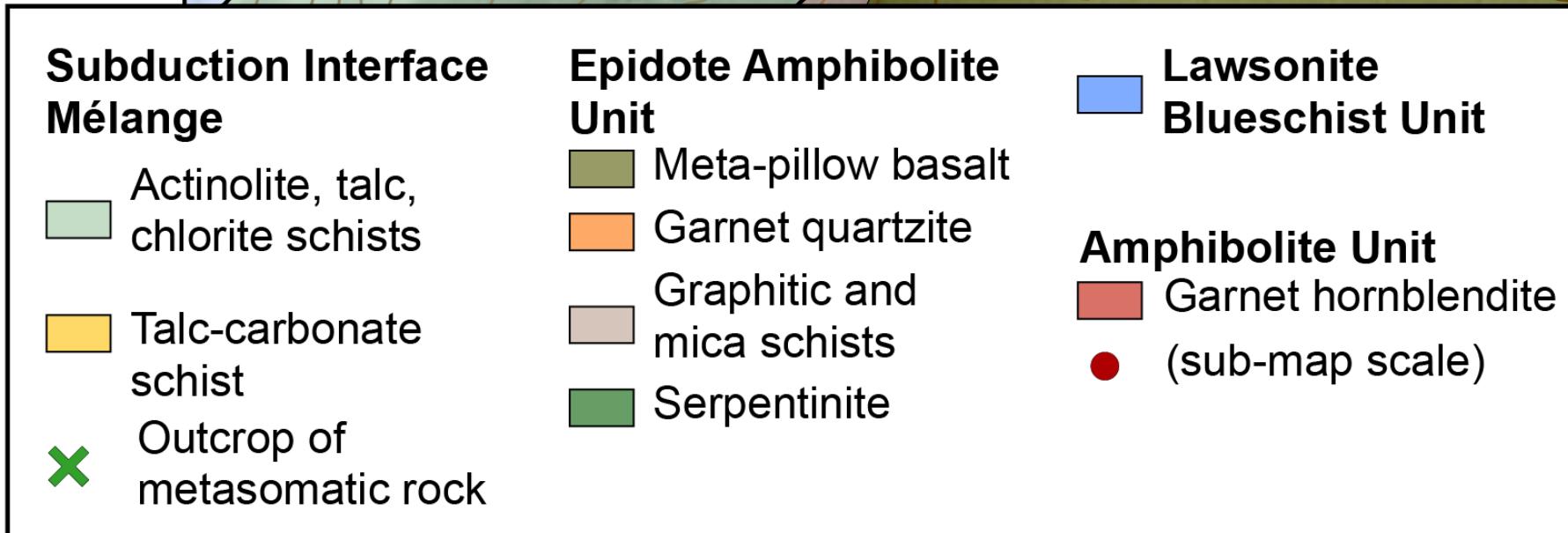
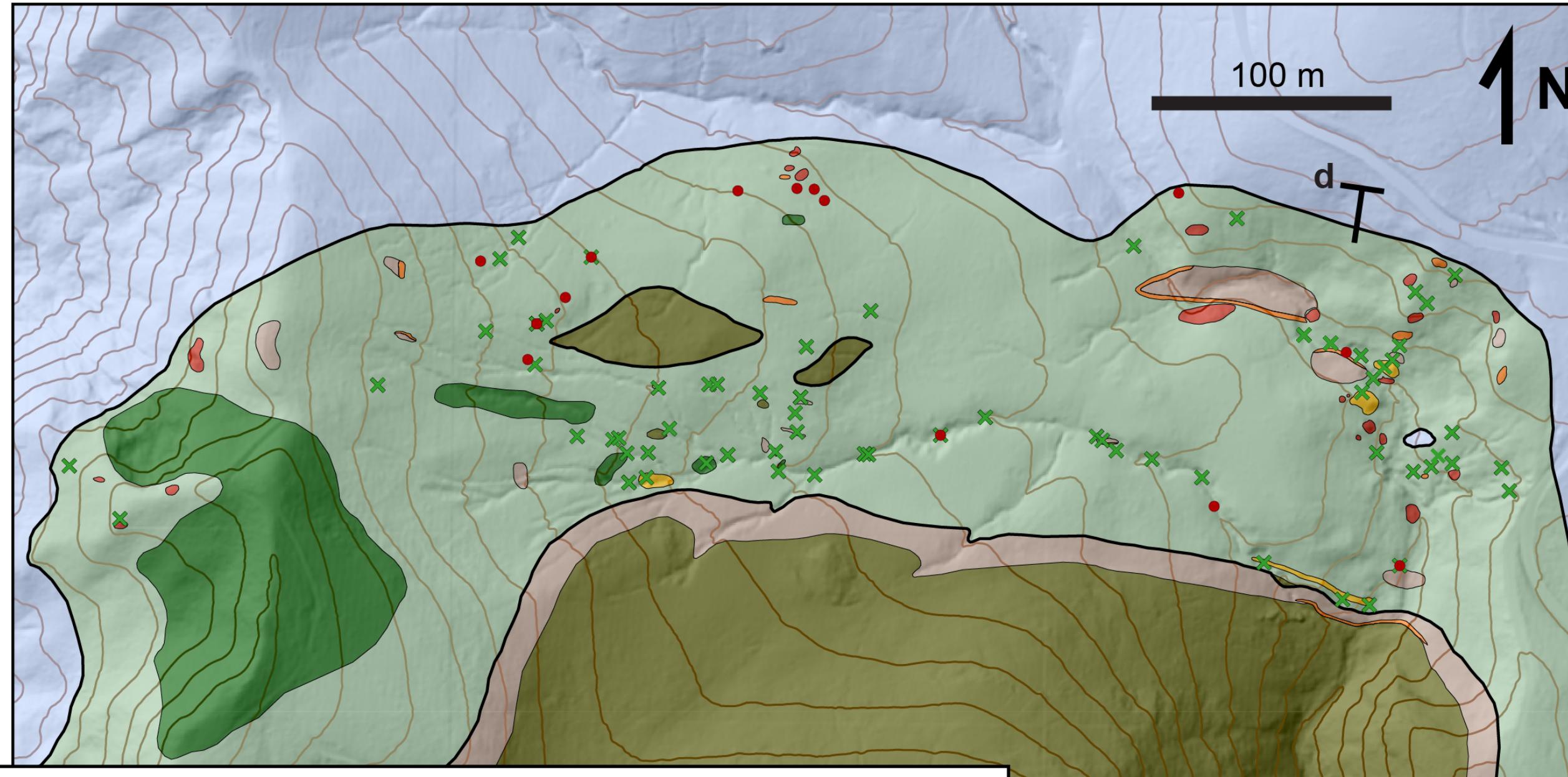


Hoover, Condit et al; 2022 GRL

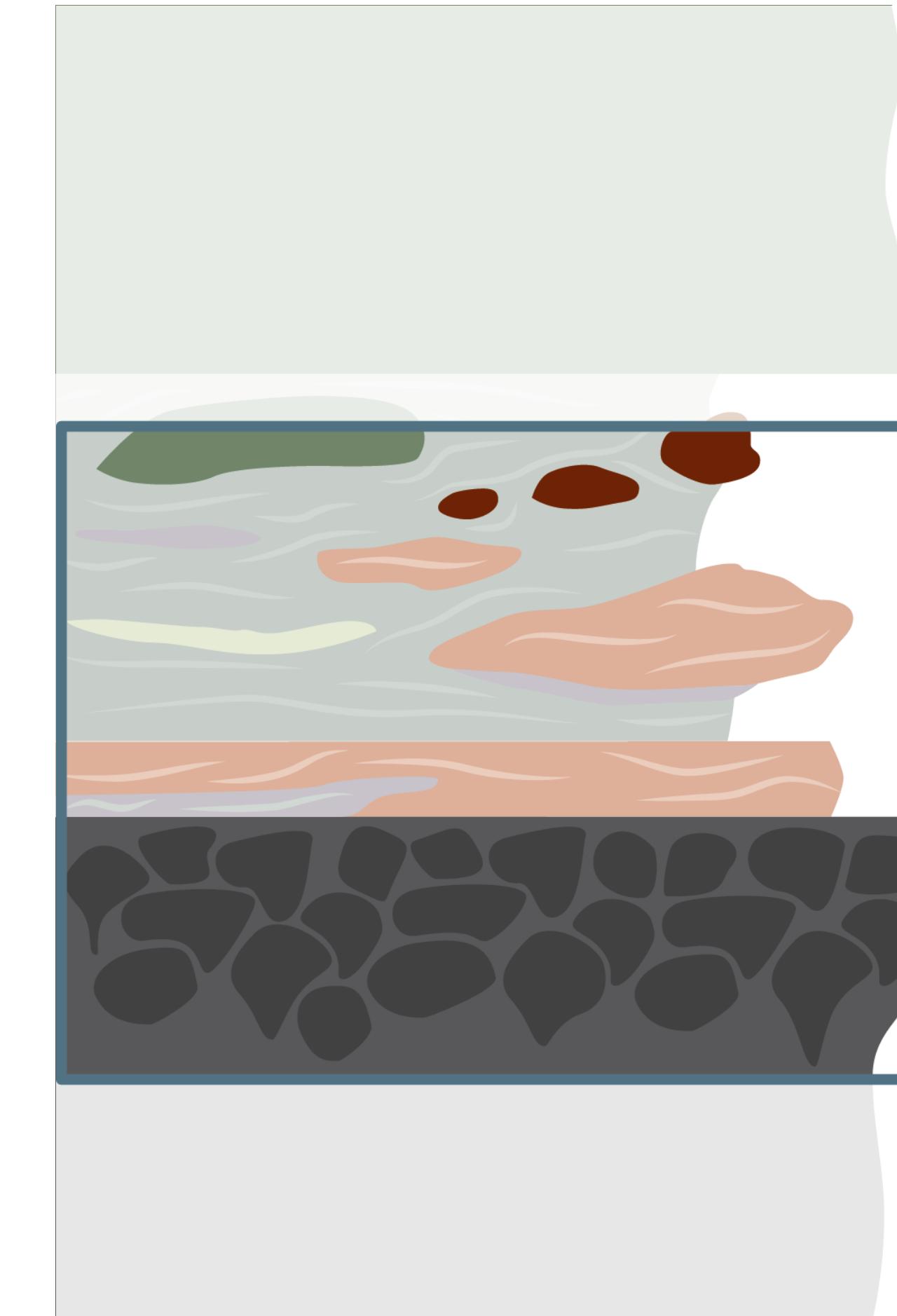
Adapted from Easthouse, Hoover et al., 2025, Geology
Lindquist, Condit et al., in prep

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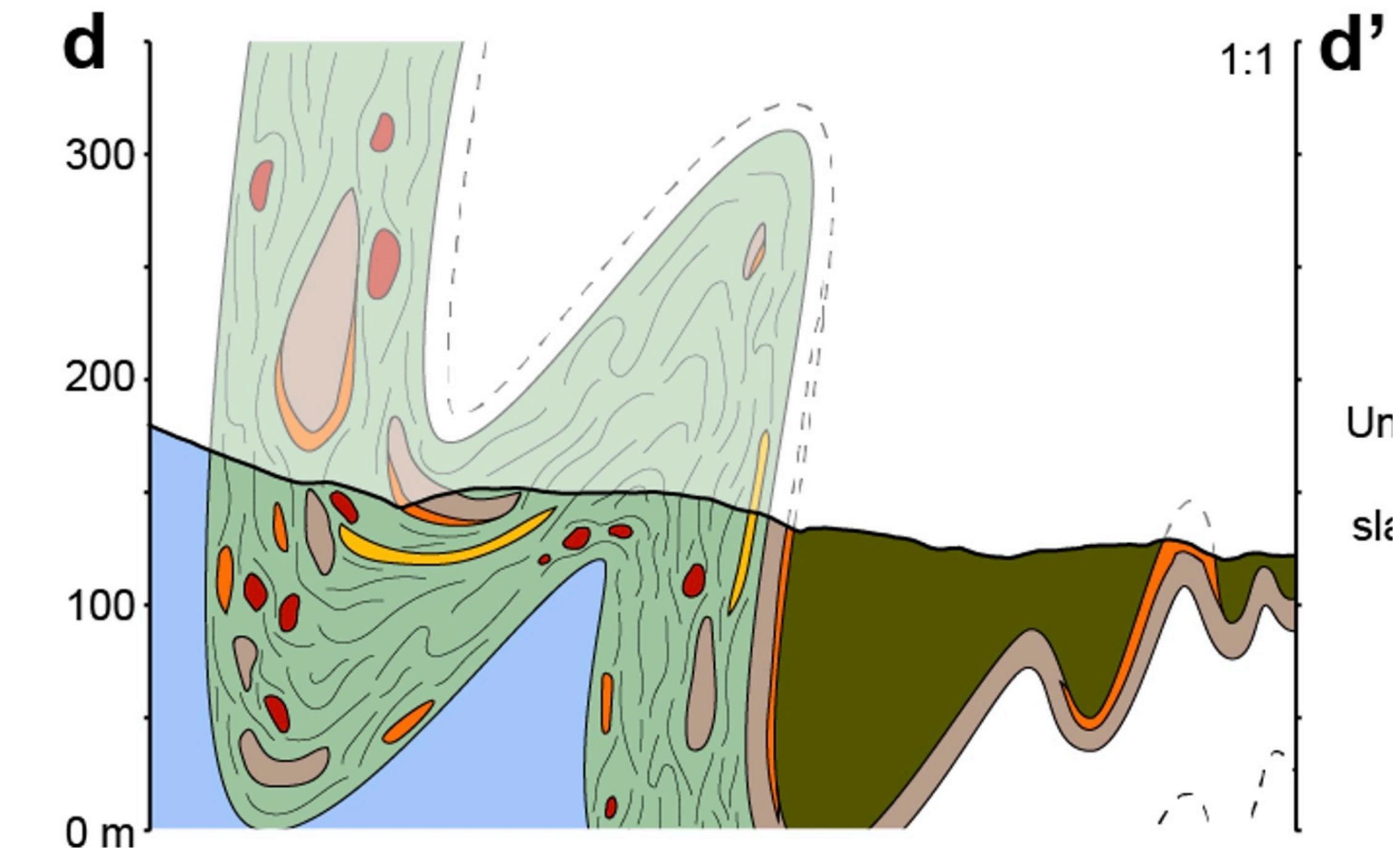
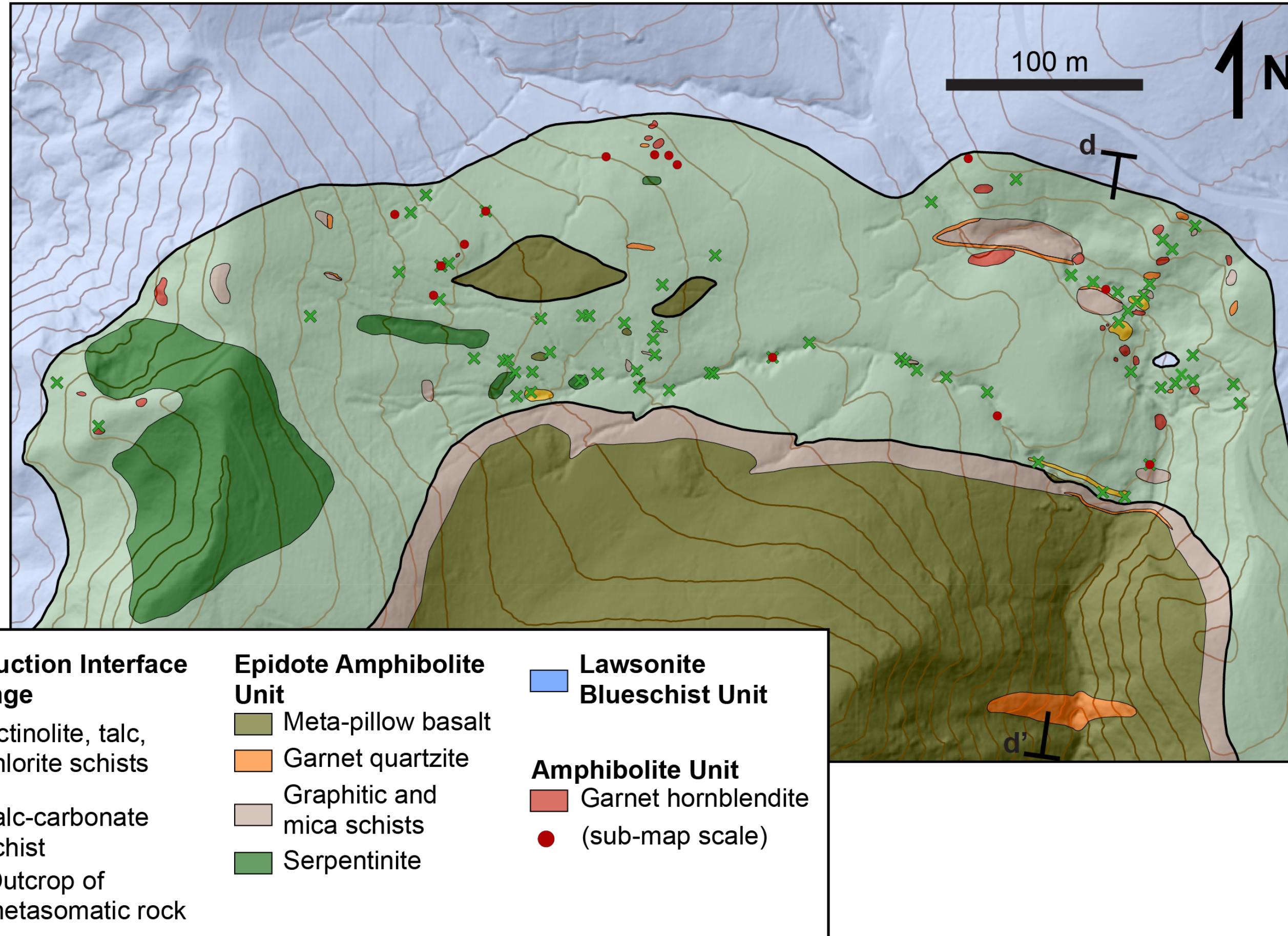
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Epidote Amphibolite Facies Plate interface: 550°C, 1.0 GPa (~35 km paleodepth)



Hoover, Condit et al., in Revision

Ultramafic-Metasedimentary melange matrix



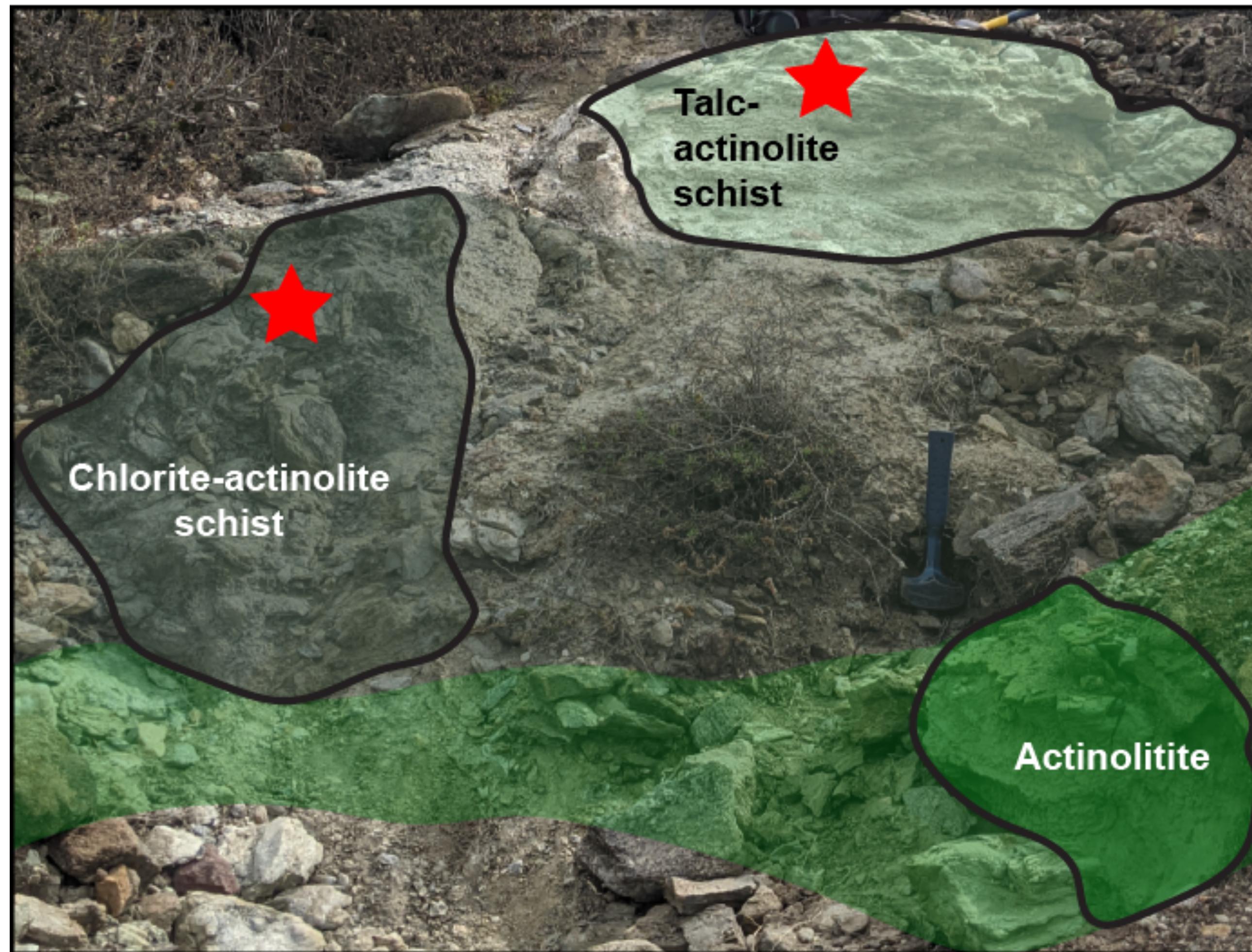
Ultramafic-pelitic melange

Mechanical and chemical mixing between

- ultramafic rocks from mantle wedge
- metasedimentary schists

Result: talc and chlorite schists

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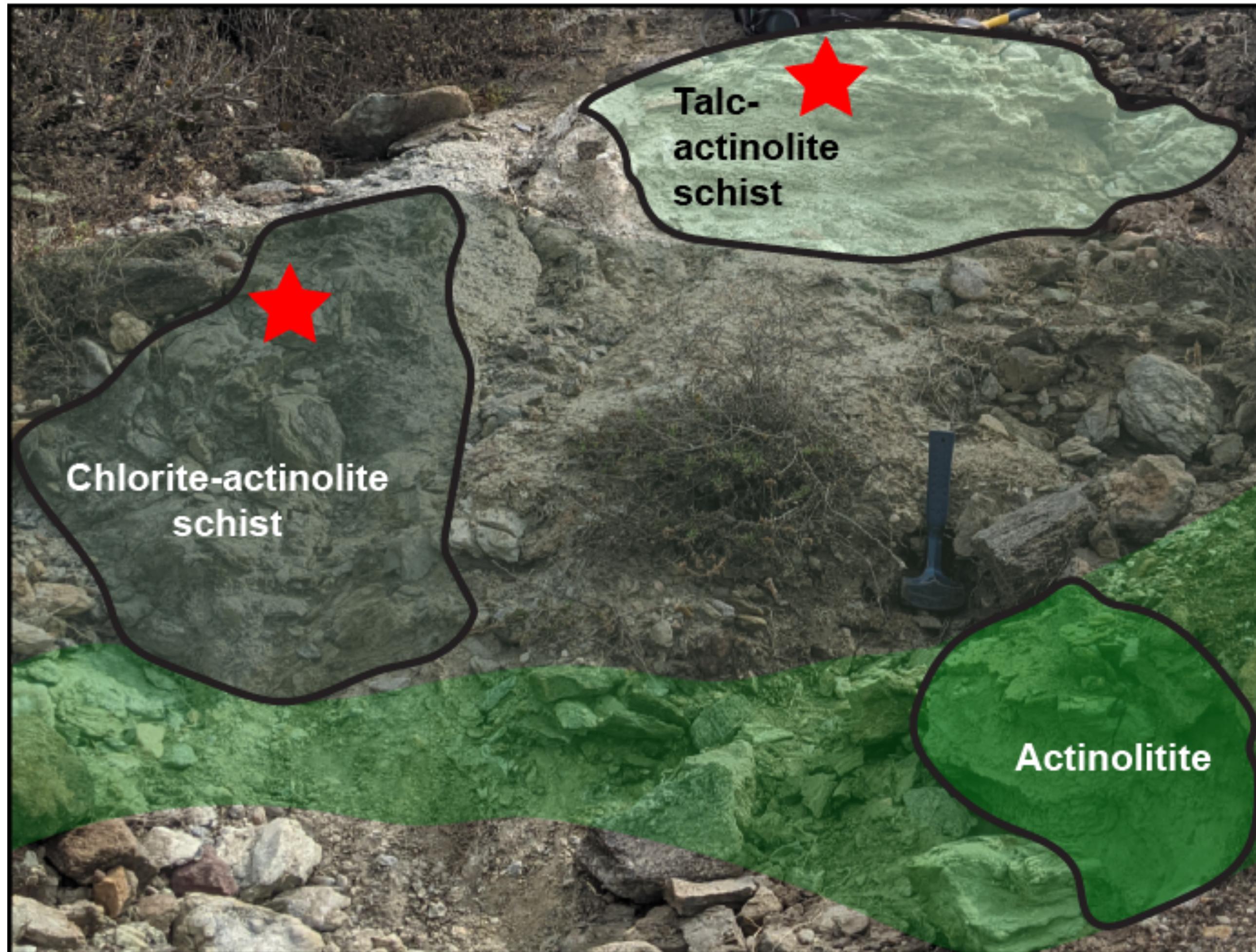
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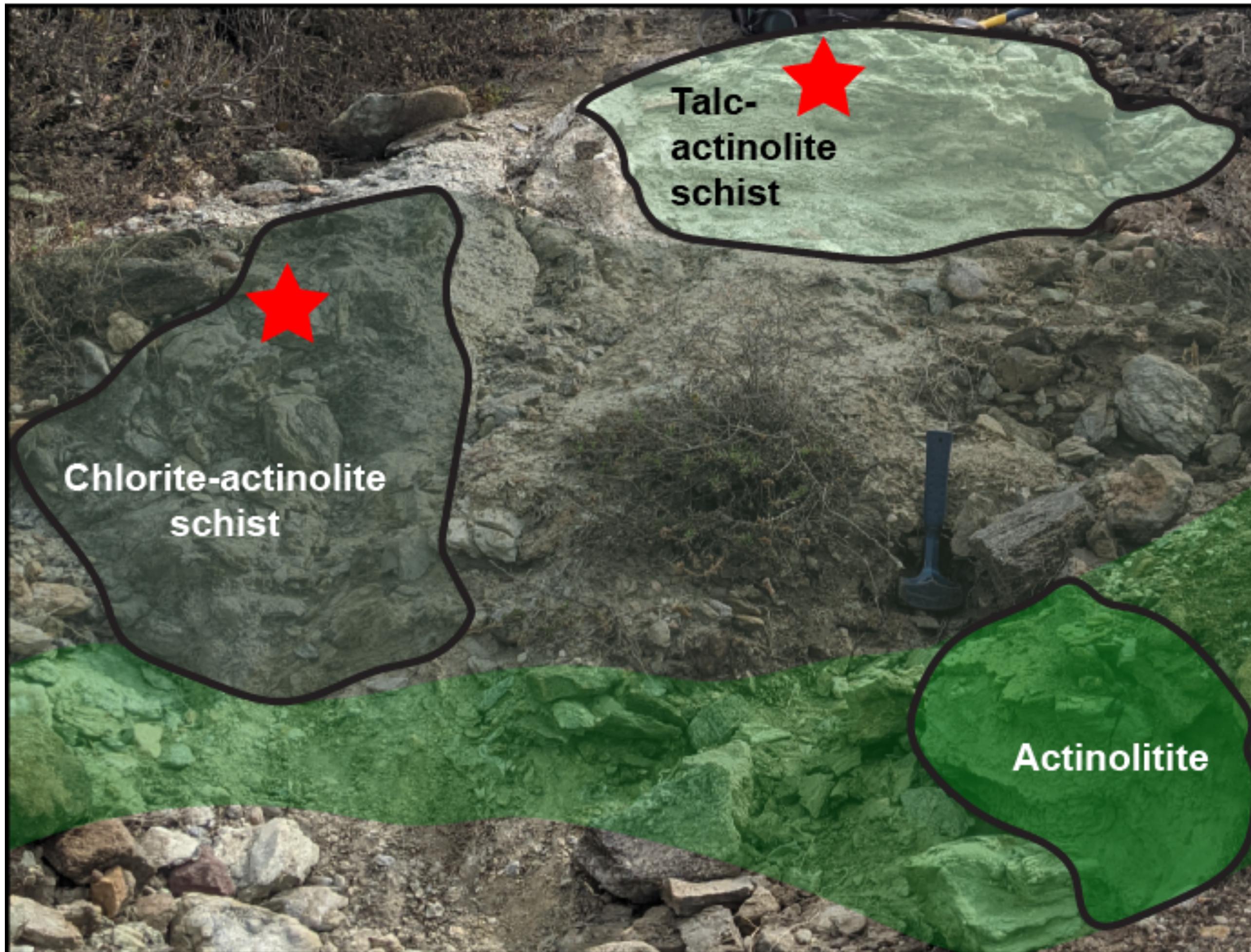
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*Recent work suggests
metasomatic rocks might host
slow slip;*

*Frictionally weak and ~rate
strengthening*

Ultramafic-Metasedimentary melange matrix



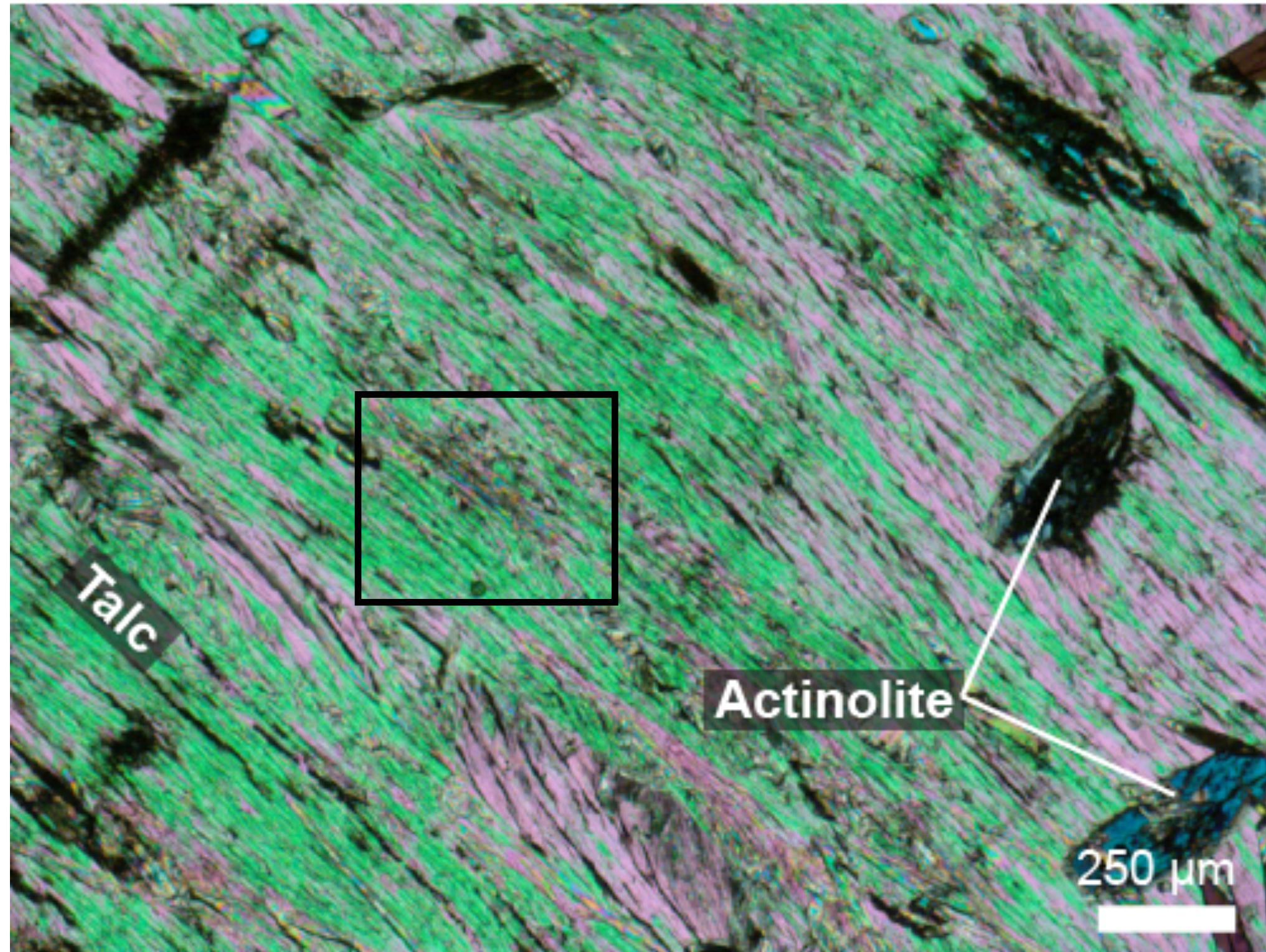
Hoover, Condit et al; 2022 GRL
Easthouse, Hoover, 2025 Geology

Evidence of fluids

Talc-Actinolite Schist - Metasomatic with an ultramafic protolith

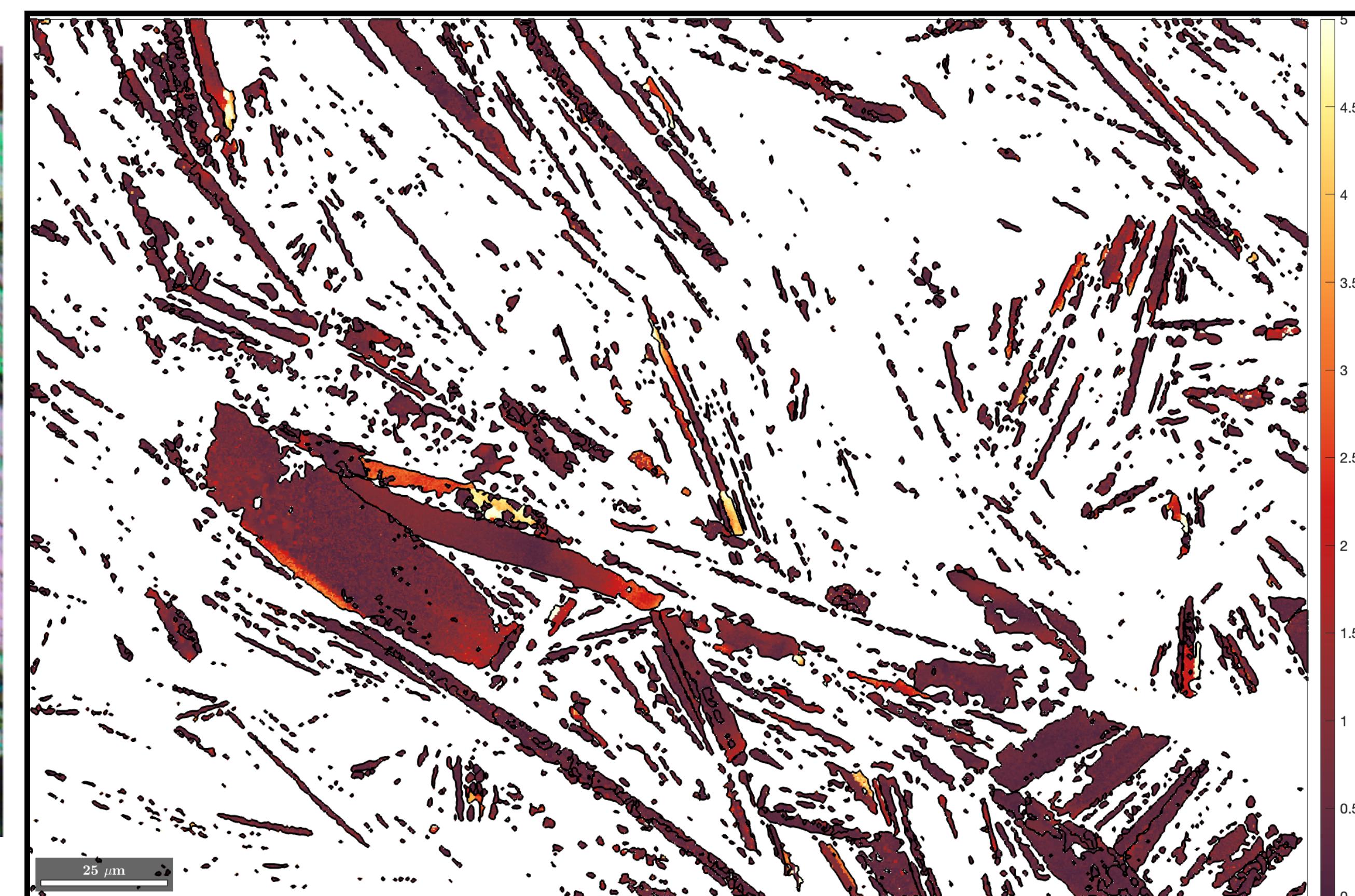
Talc Schist

Photomicrograph



Talc strongly aligned in some places;

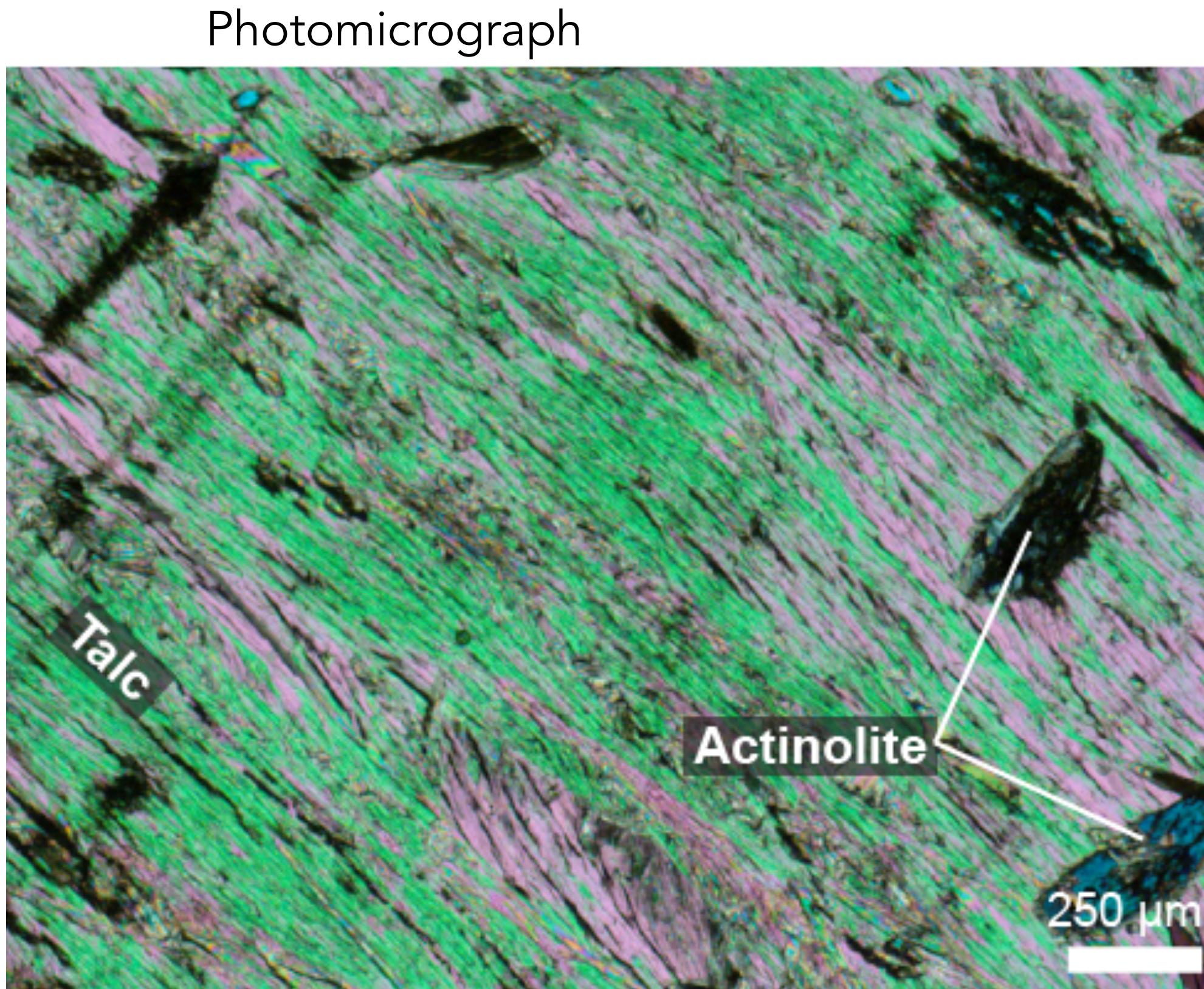
Talc Mis2Mean EBSD map



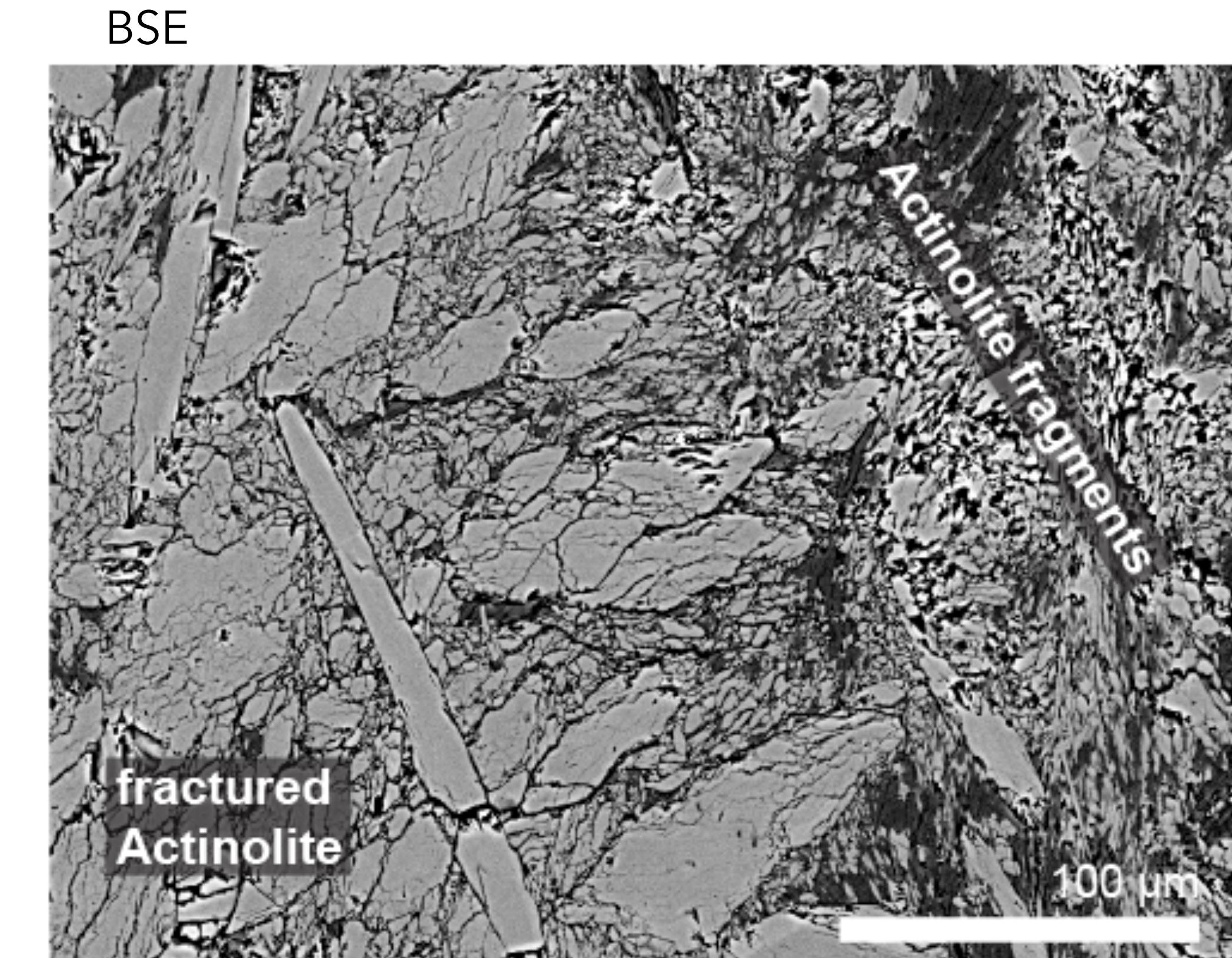
Hoover, Condit, Cross; In Prep

Talc-Actinolite Schist - Metasomatic with an ultramafic protolith

Talc Schist



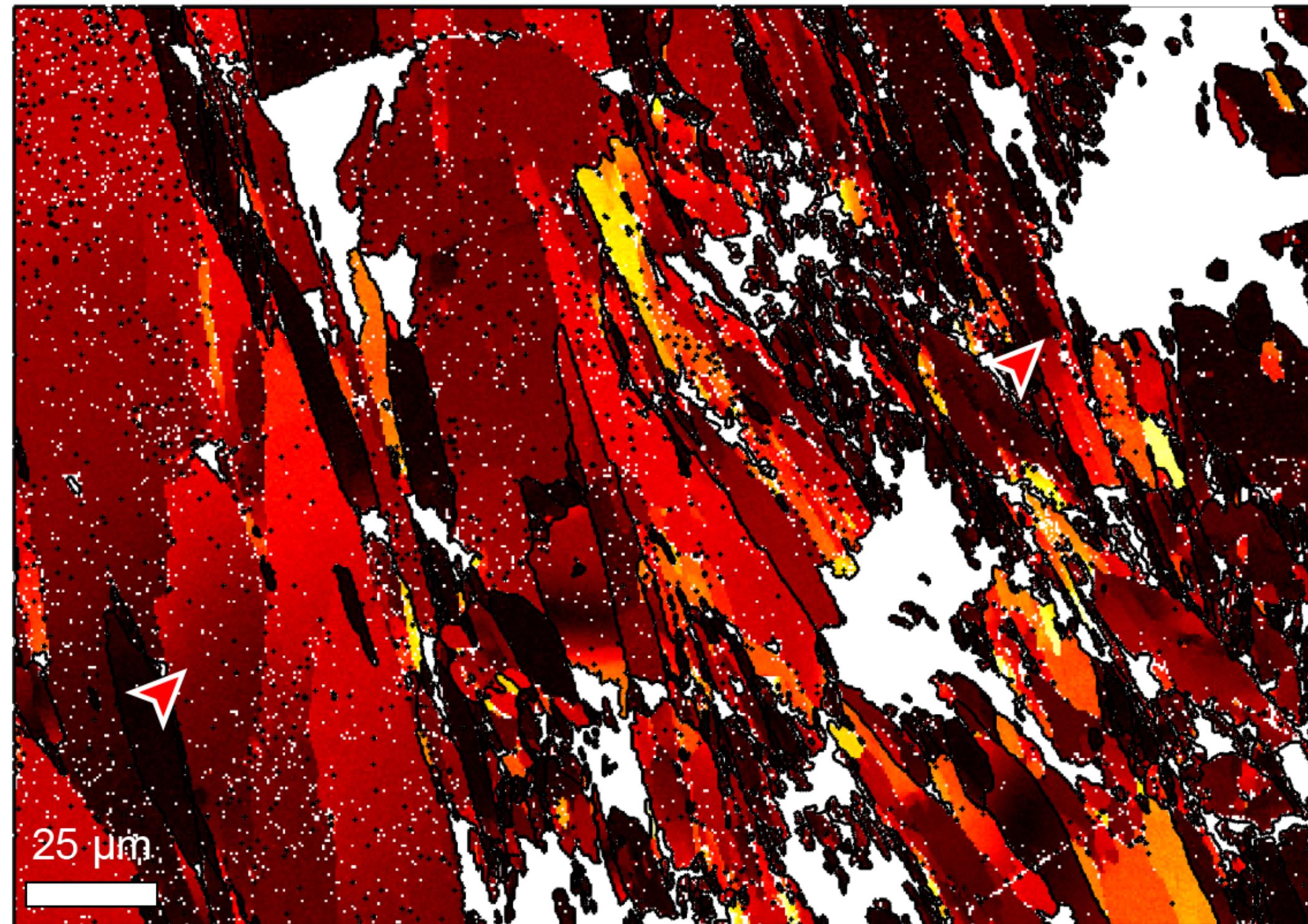
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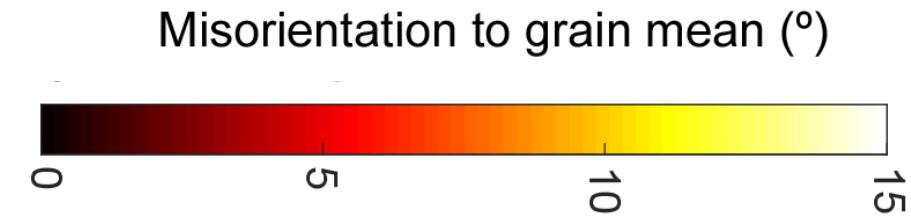
Actinolite has subgrain formation and cataclastic features

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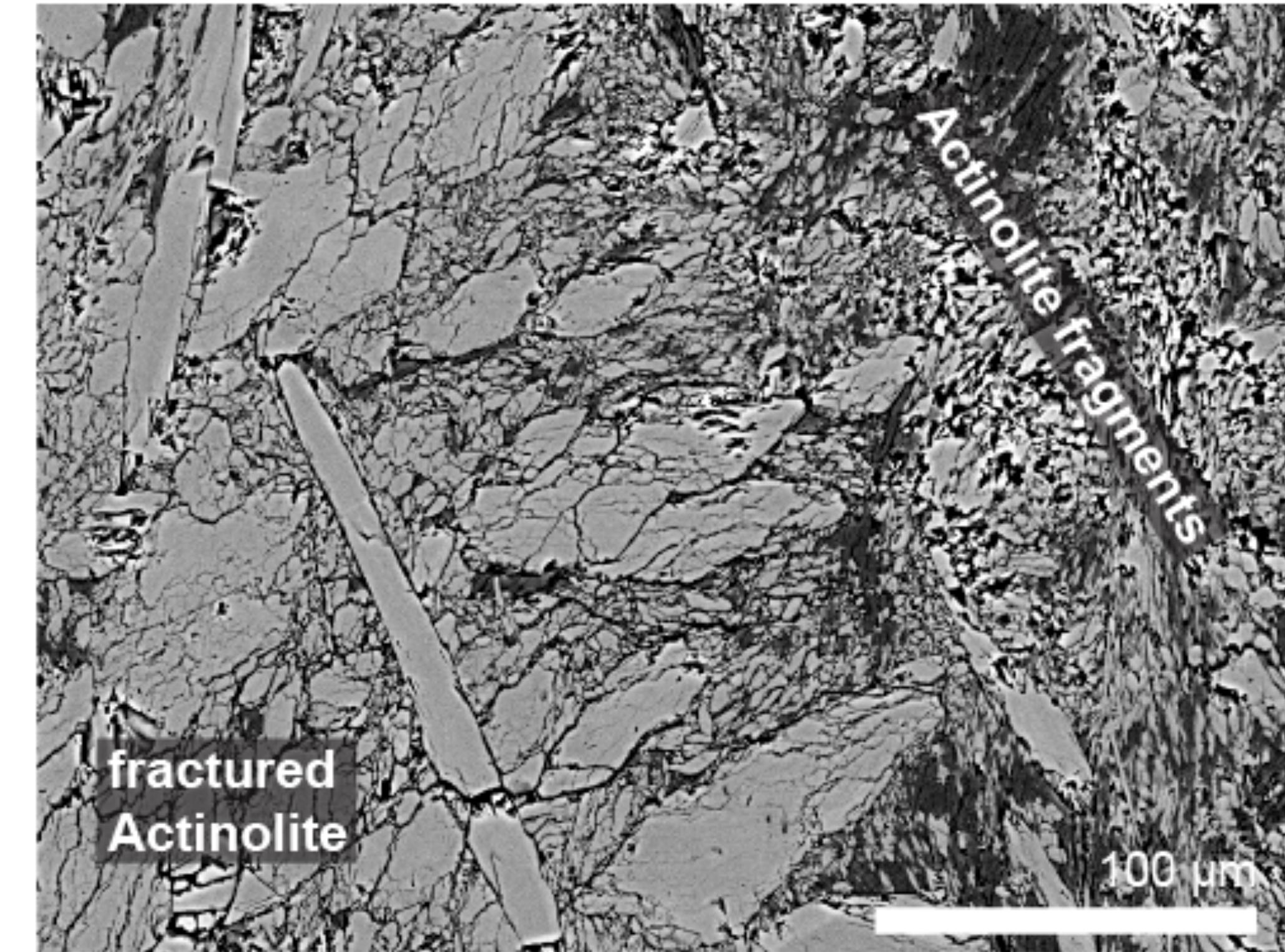
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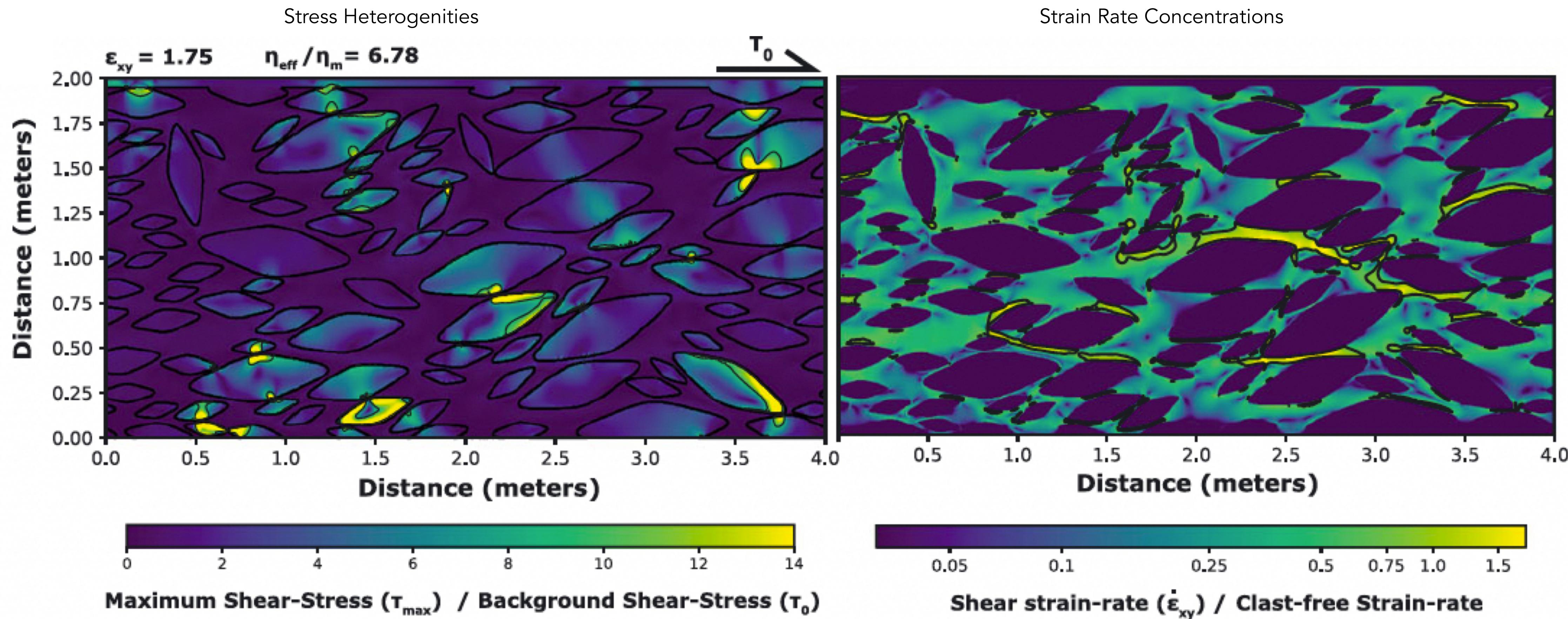
Hoover, Condit, et al; 2022 GRL



Actinolite has subgrain formation and cataclastic features

Actinolite deforming by ~higher stress mechanisms

Geodynamic models of stress amplifications in block and matrix structures

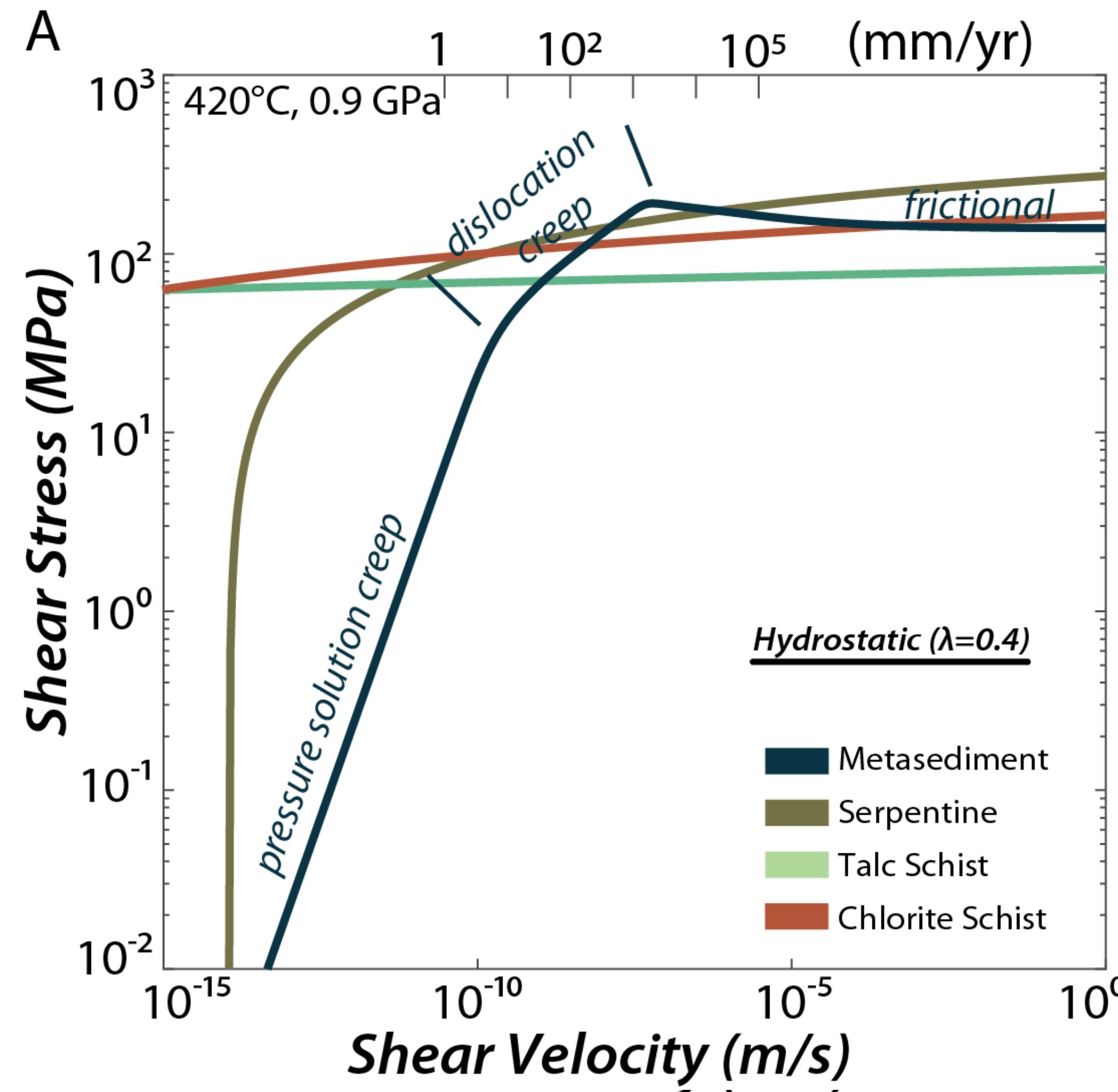


Stress amplifications in melange matrix:

Beall et al., 2019, GRL

- Weak matrix (in our case Talc) can localize strain and deform quickly
- Create force chains + high stress zones in stronger blocks (in our case actinolite)

How weak might the talc be?



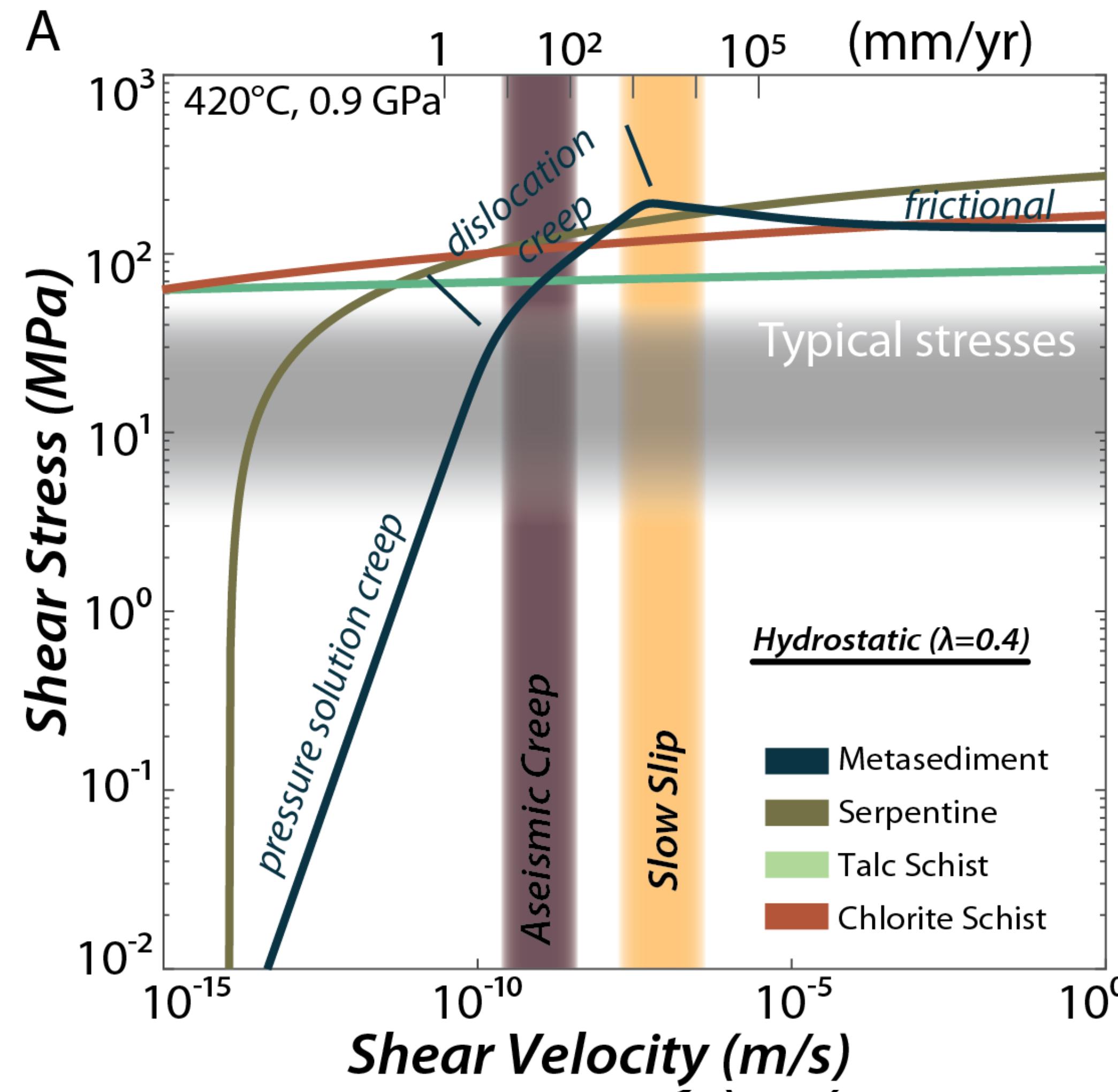
Experimental studies: talc is very weak, frictional deformation is important at a range of P-T conditions

Horn & Skemer, 2023, Boneh et al., 2023

After French & Condit, 2019

Hoover, Condit, and Cross in prep

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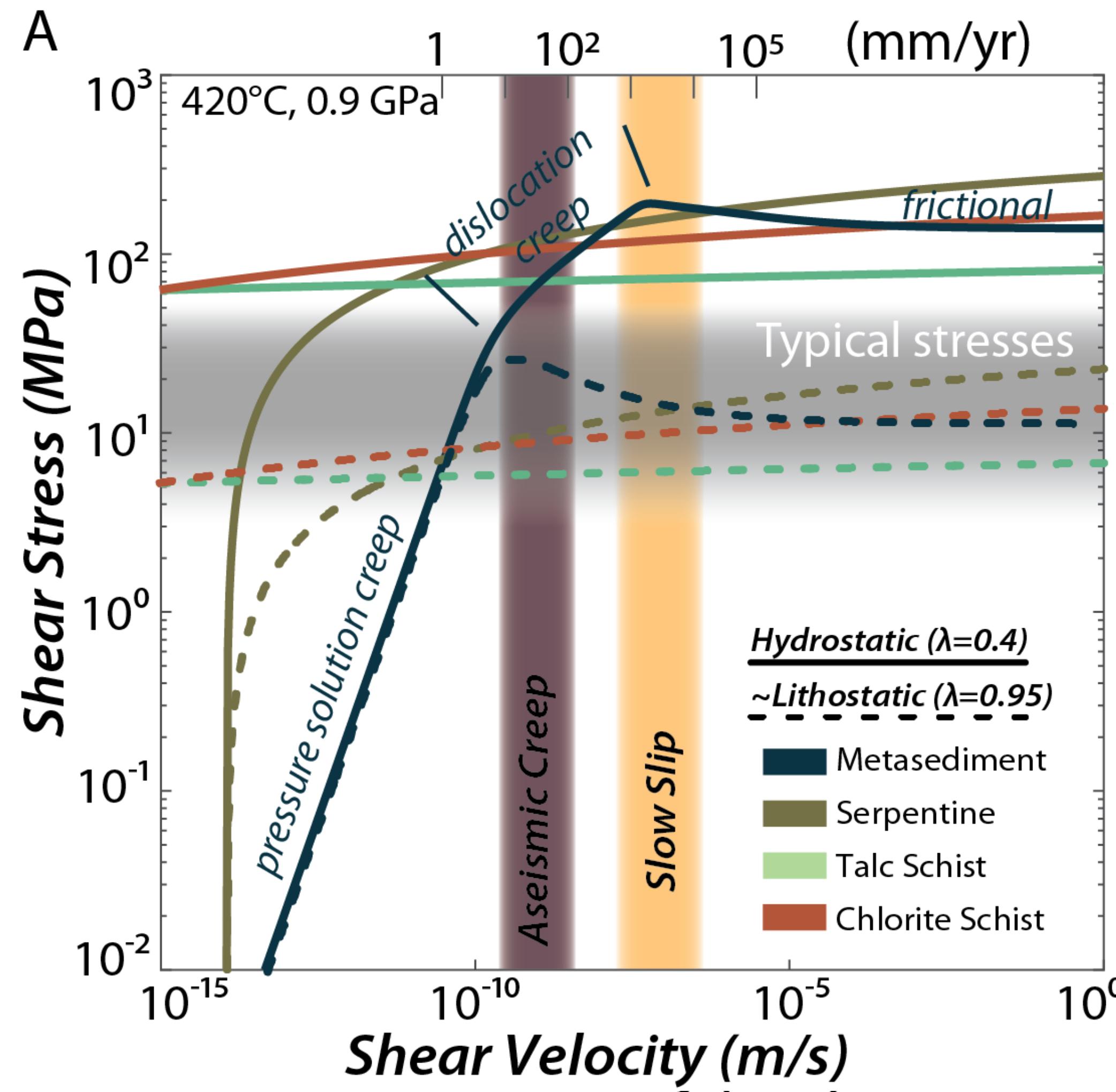
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Aseismic creep and **slow slip** velocities

Hydrostatic P_f - high stresses

How weak might the talc be? Can it host slow earthquakes with high P_f ?



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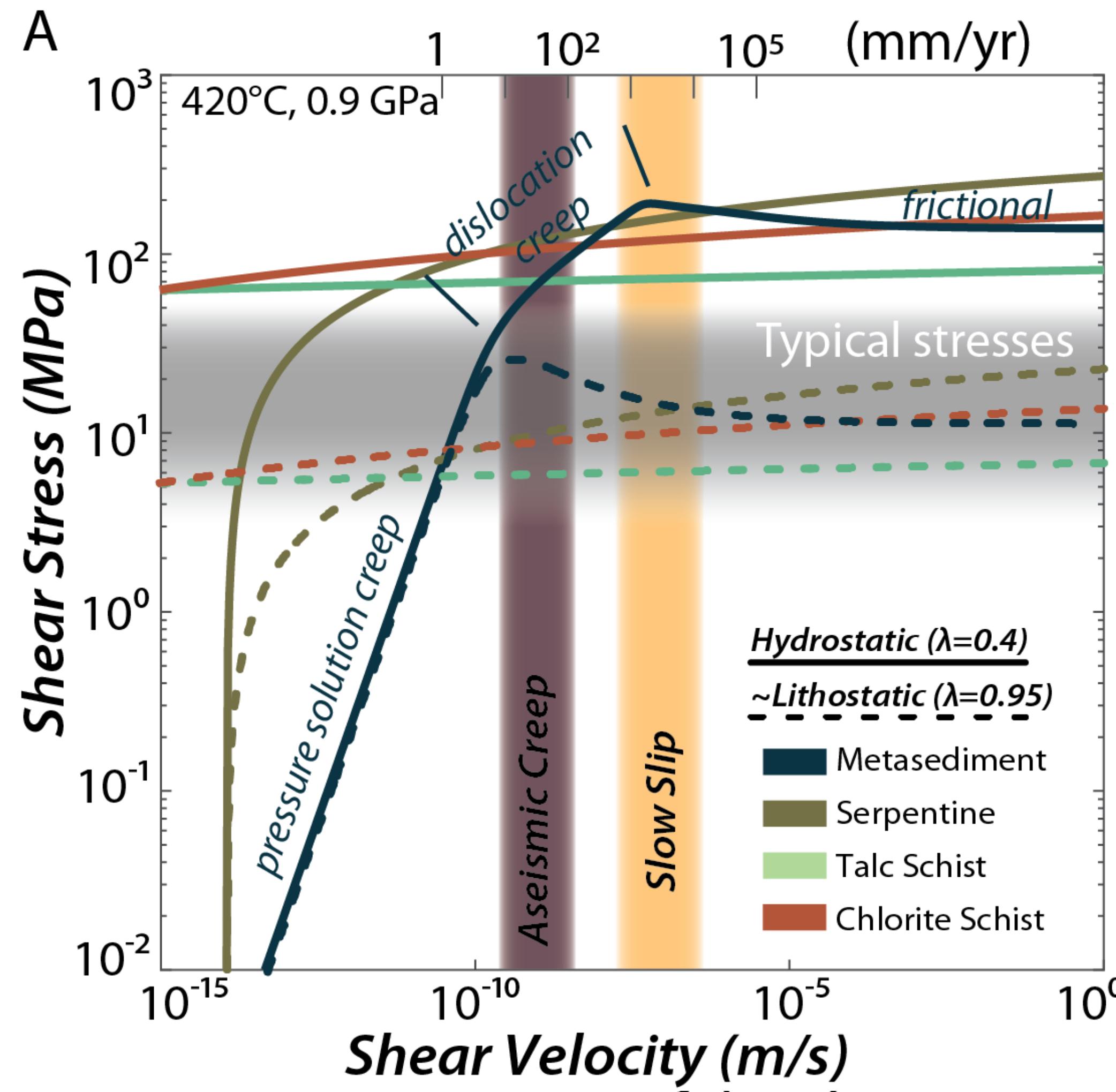
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Aseismic creep and slow slip velocities

Hydrostatic P_f - high stresses

Near lithostatic P_f - subduction zone stresses accommodate slow slip with frictional deformation

How weak might the talc be? Can it host slow earthquakes with high P_f ?



If P_f are ~lithostatic, talc-schists could host slow slip events

Near lithostatic P_f - subduction zone stresses accommodate slow slip with frictional deformation

Observations from the exhumed rock record from the base of the subduction seismogenic zone

What are the phenomena we see in rocks that we need improved:

...Constitutive relations for?

...Dynamic mechanistic understanding of chemical-mechanical feedbacks?

Pressure Solution Creep along the plate interface is not slow slip

Activation of frictional mechanisms with high Pf could be slow slip

Evidence for brittle deformation (**seismic slip?**) at these deep depths intimately related to fluid process

Ios Island, Greece
Cycladic Blueschist Unit

Mutually crosscutting
relationships between
early exhumation
subduction fabrics and
quartz veins \pm
pseudotachylite (?)

Fluids appear intimately
related to brittle failure \pm
seismic slip at ~ 400 °C,
0.8-1.0 GPa (~ 25 -35 km
depth)



Evidence for brittle deformation (seismic slip?) at these deep depths intimately related to fluid process

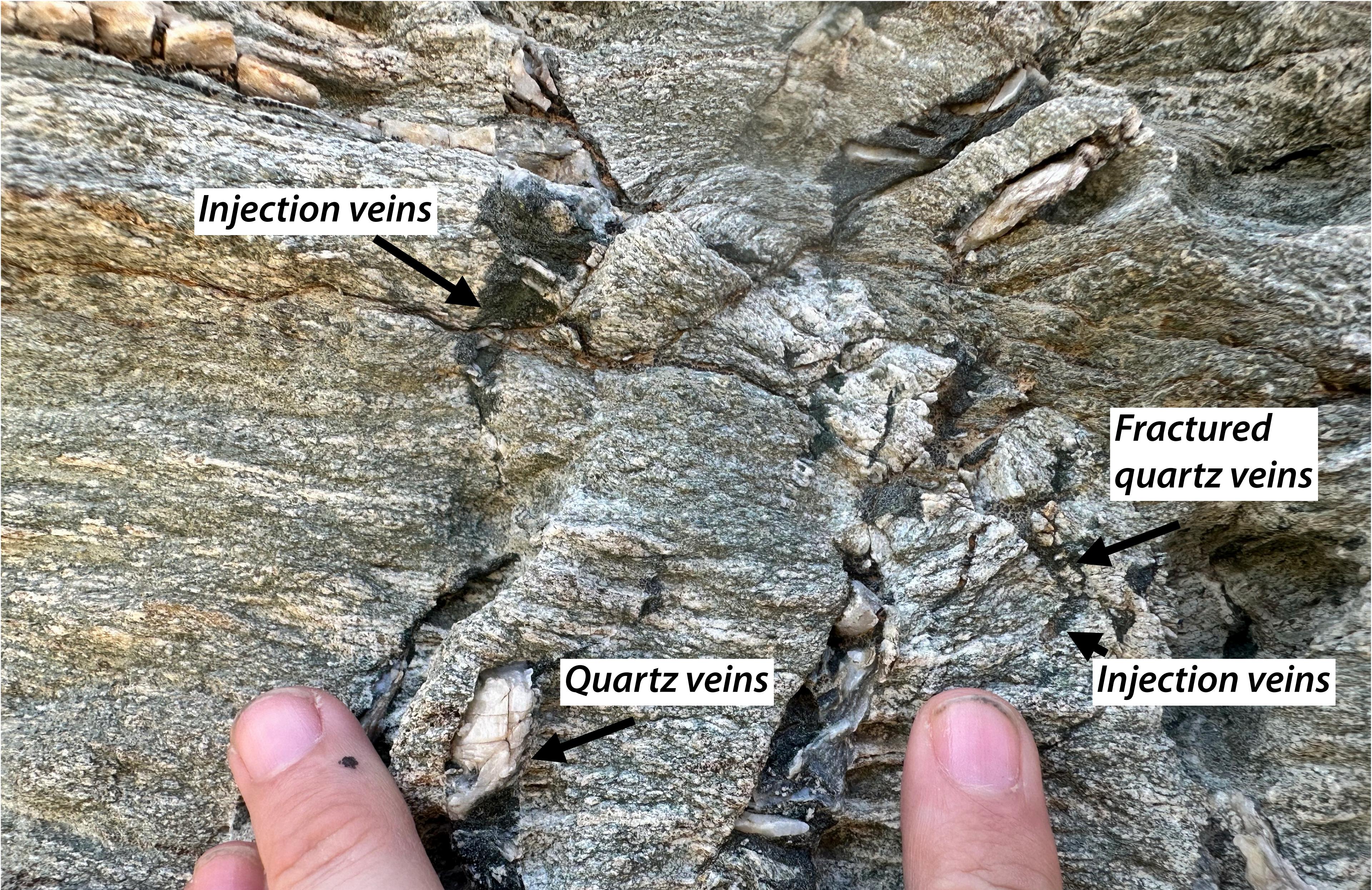
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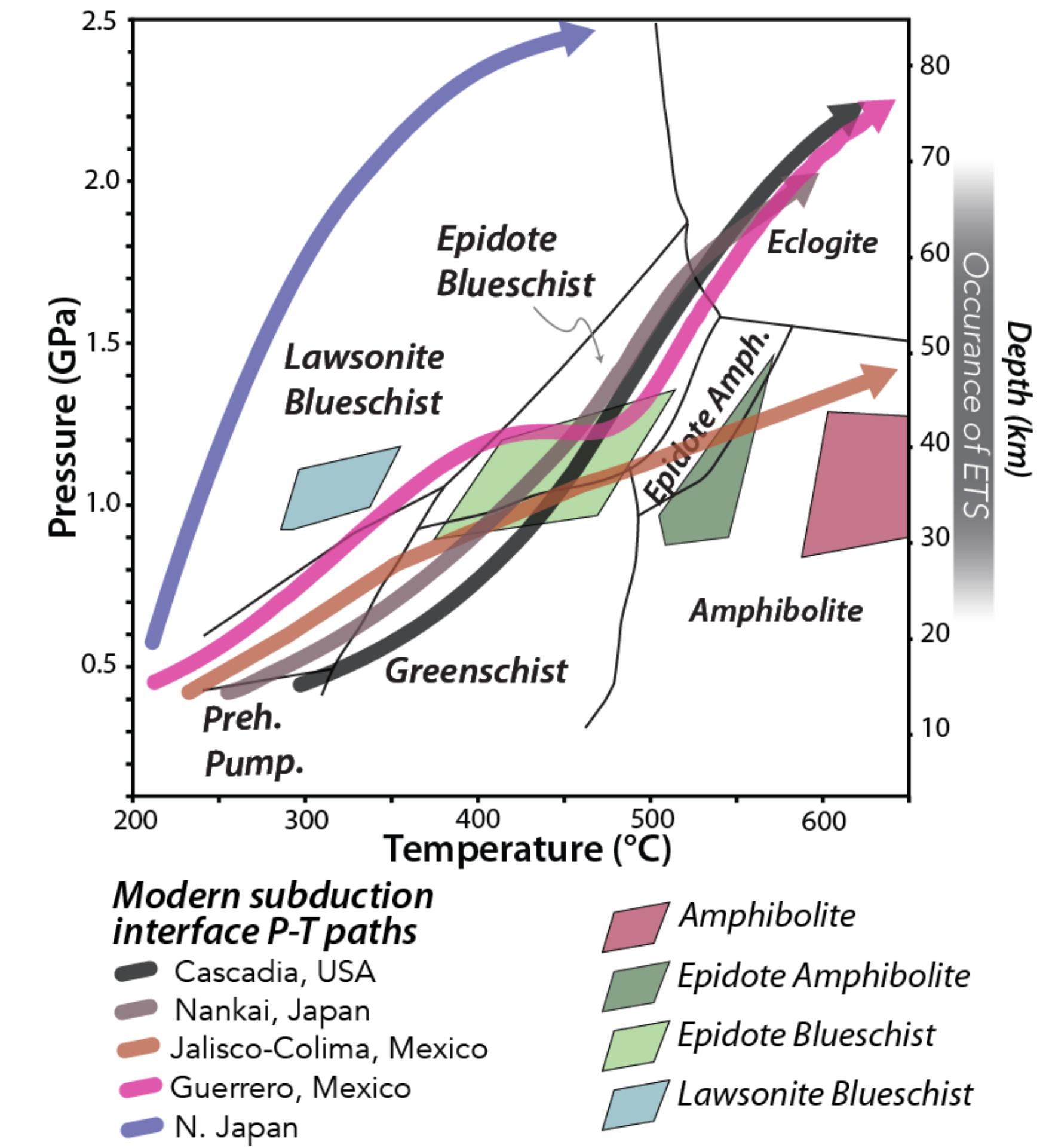
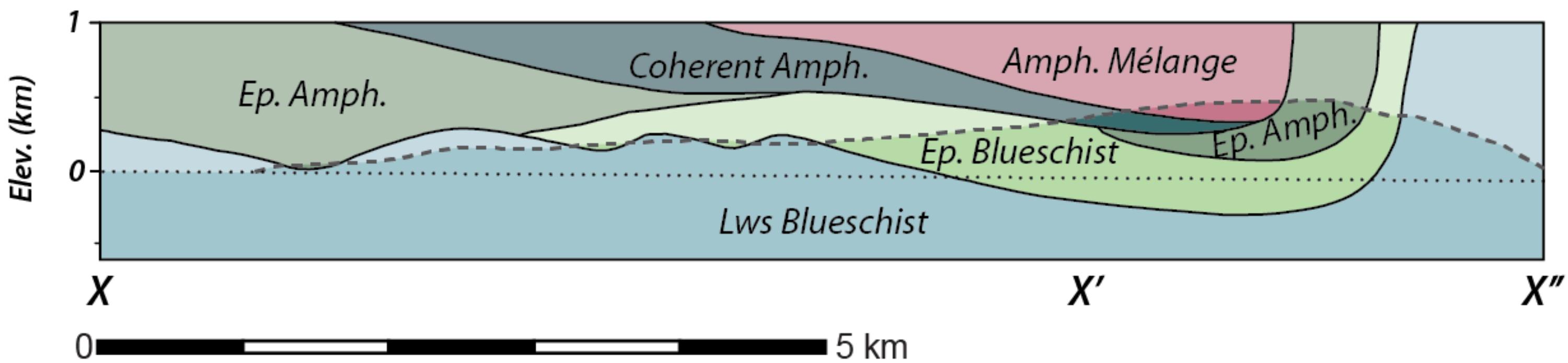
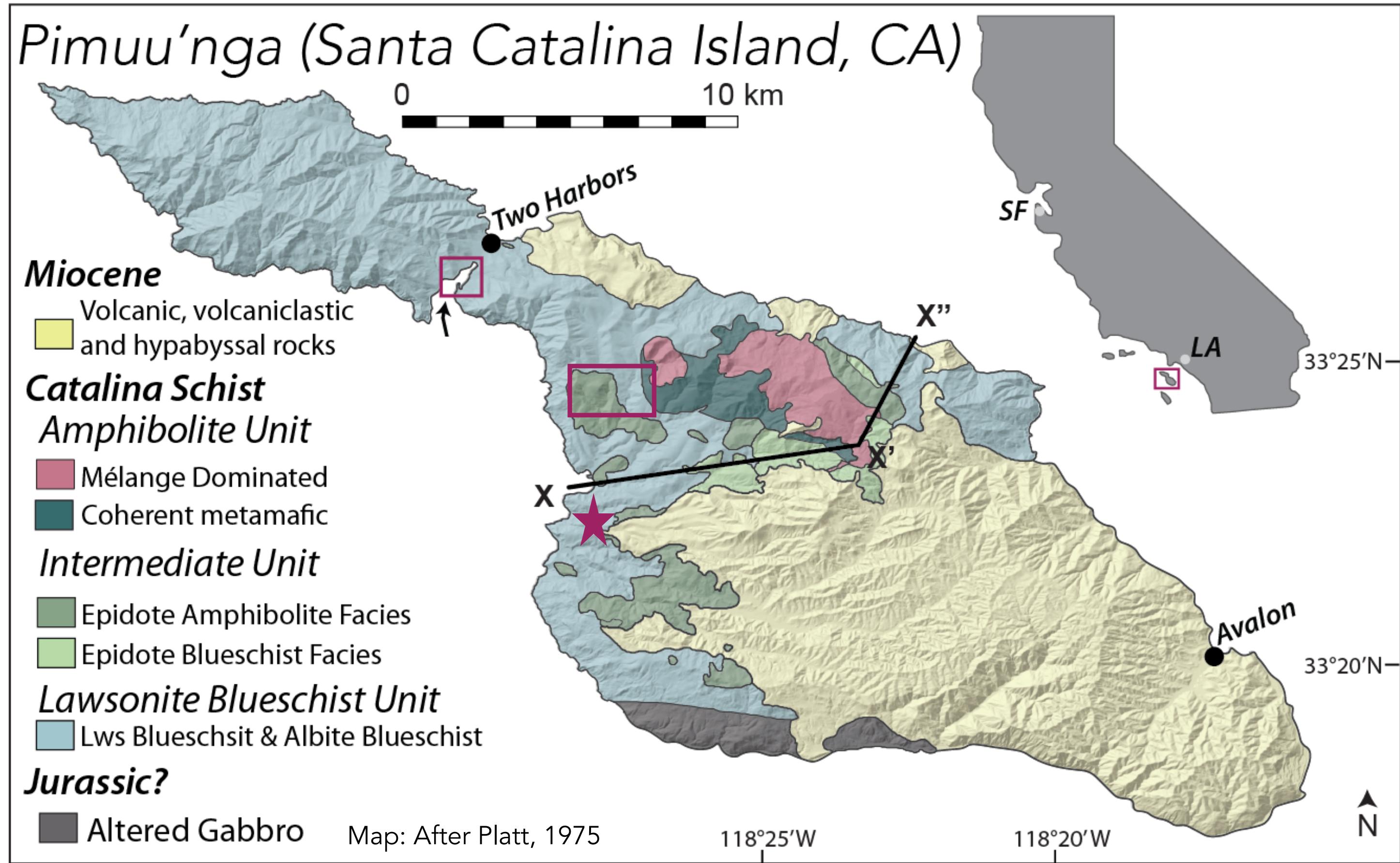
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Lawsonite blueschist unit - ~1.0 GPa (30-35 km) and ~320-330 °C

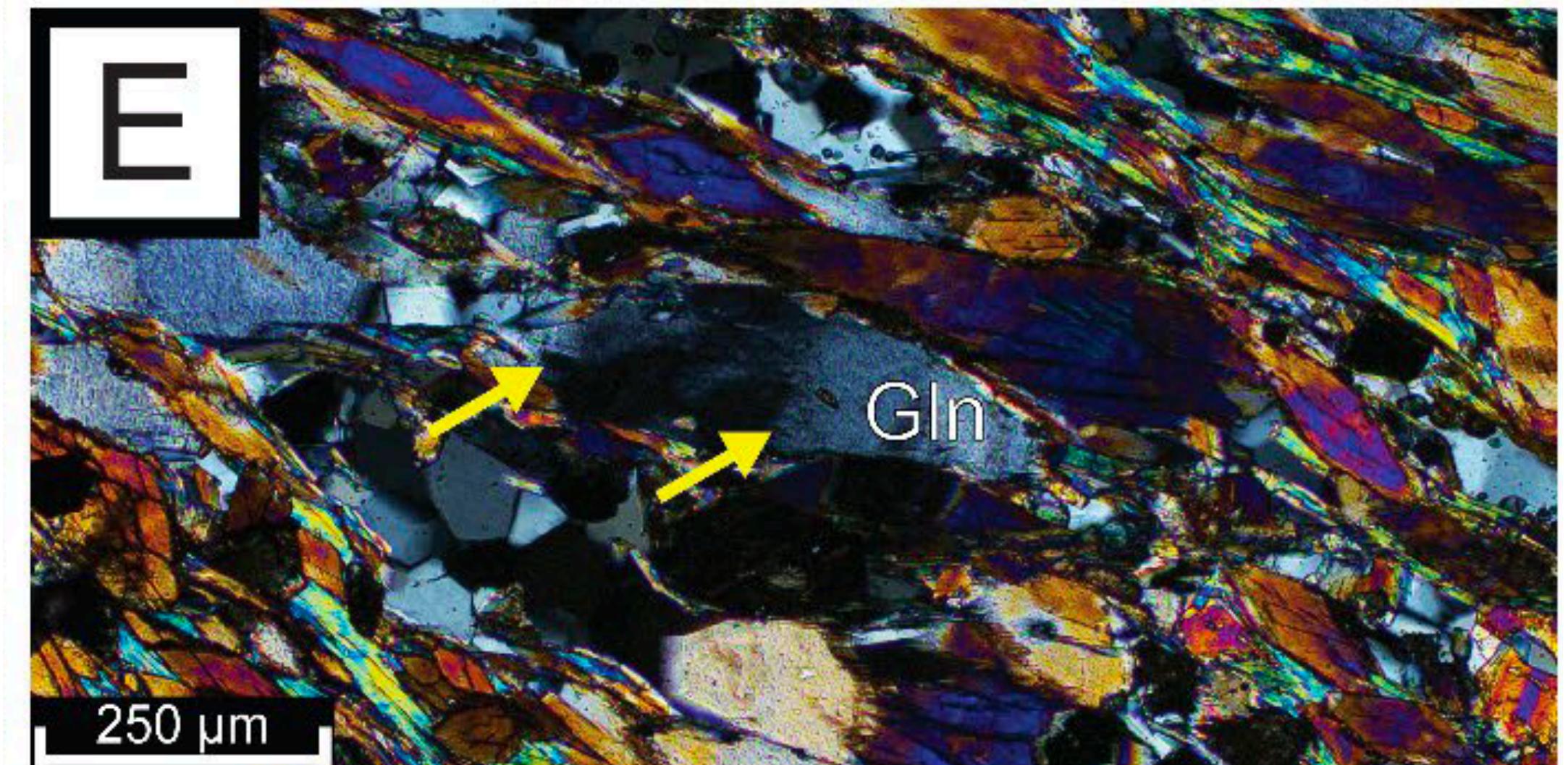
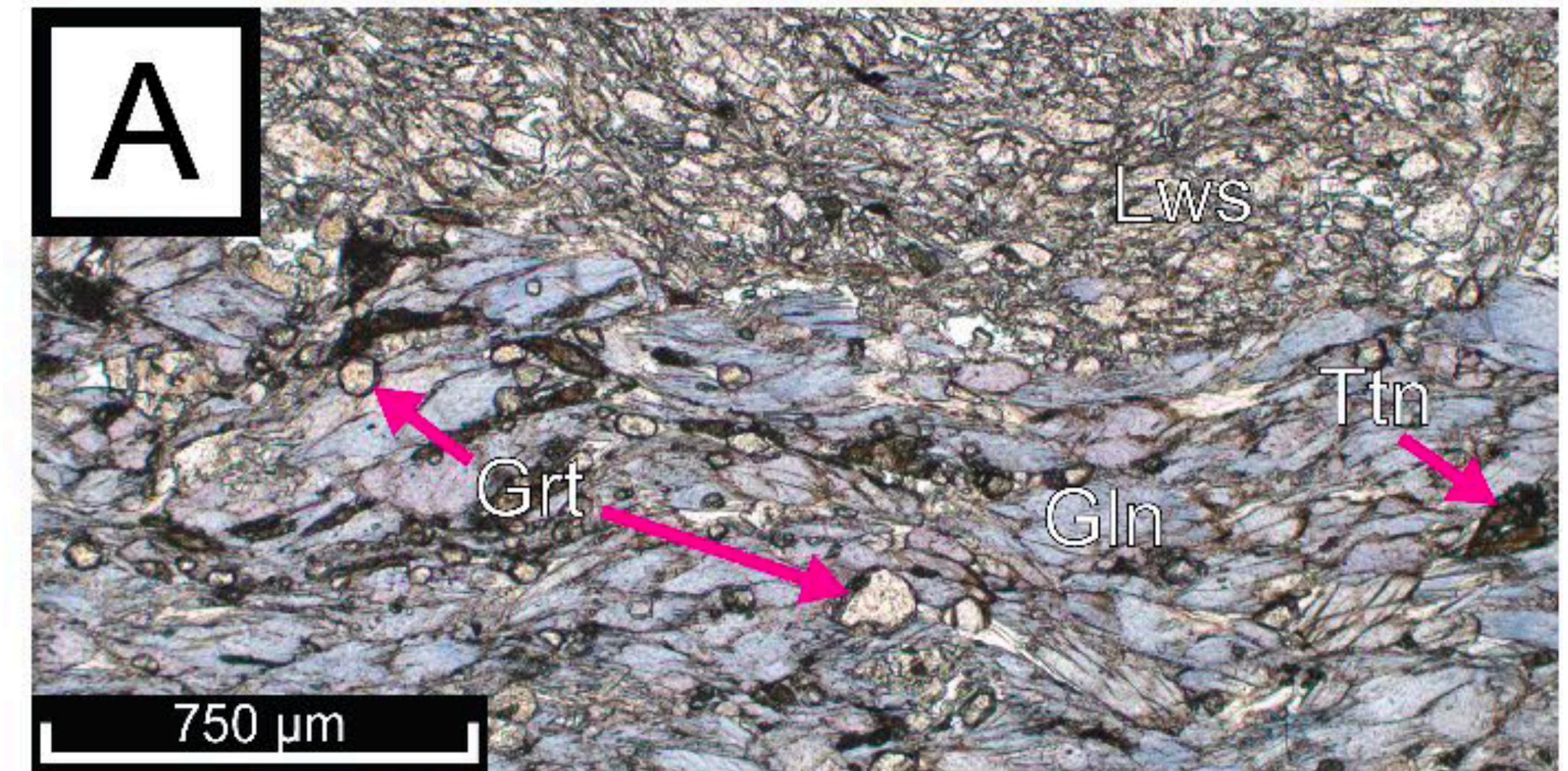
Temperatures: Platt and Schmidt, 2024

Dislocation to diffusive mechanism shift in glaucophane

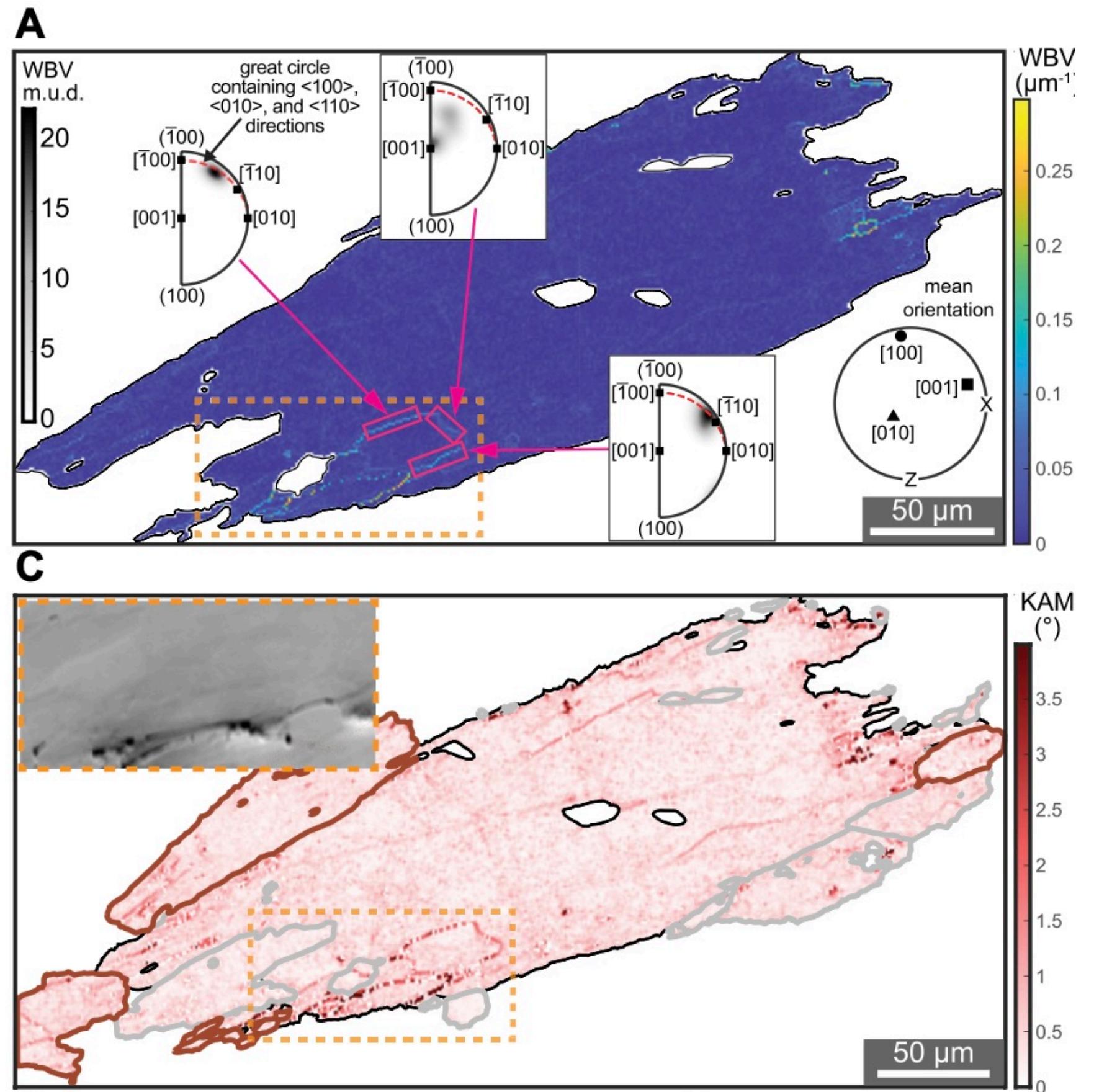


Mafic blueschists (the metamorphosed slab) accommodate a lot of deformation; Gluacophane (sodic amphibole) the main strain accommodating phase.

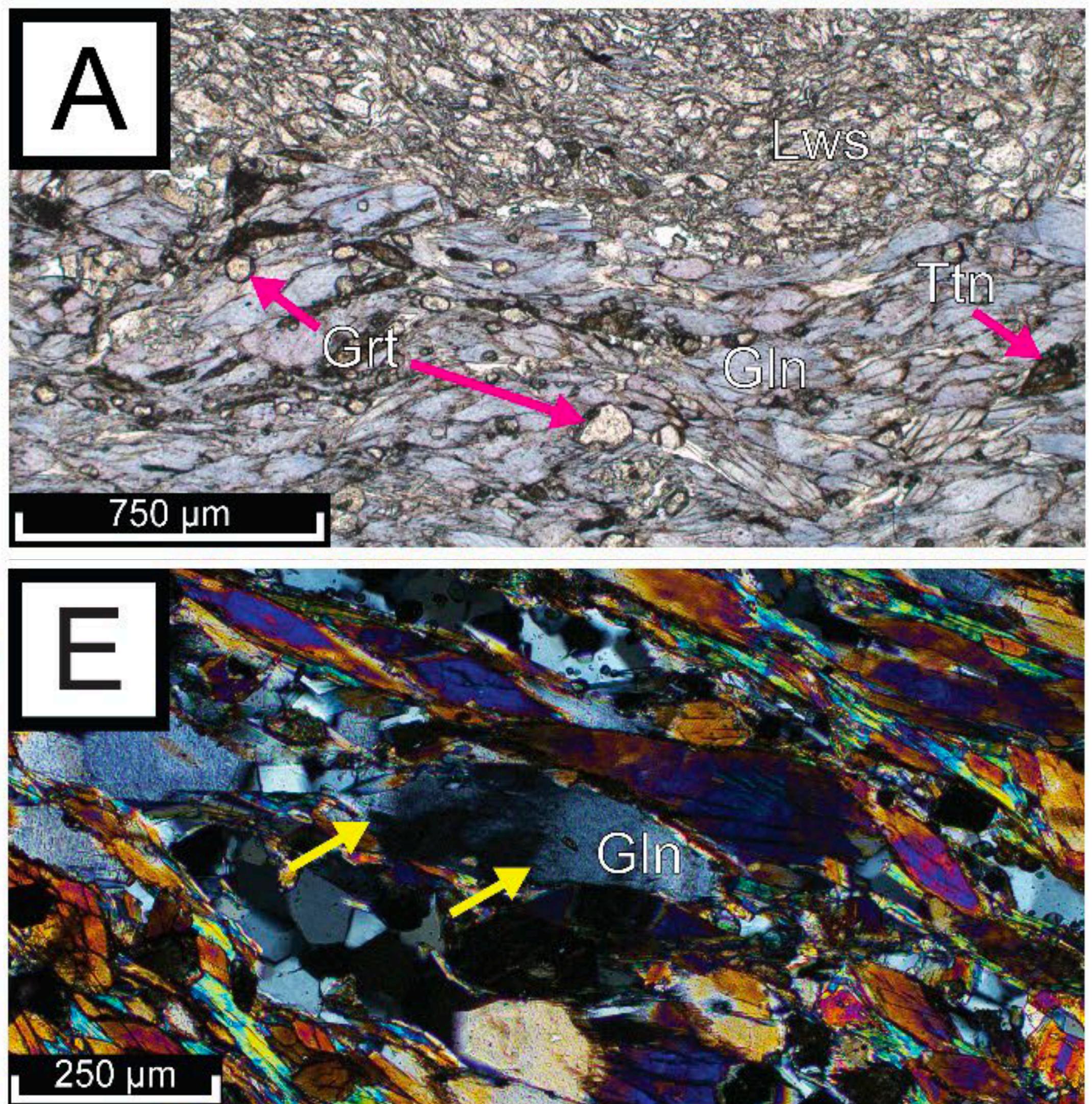
How does glaucophane deform?



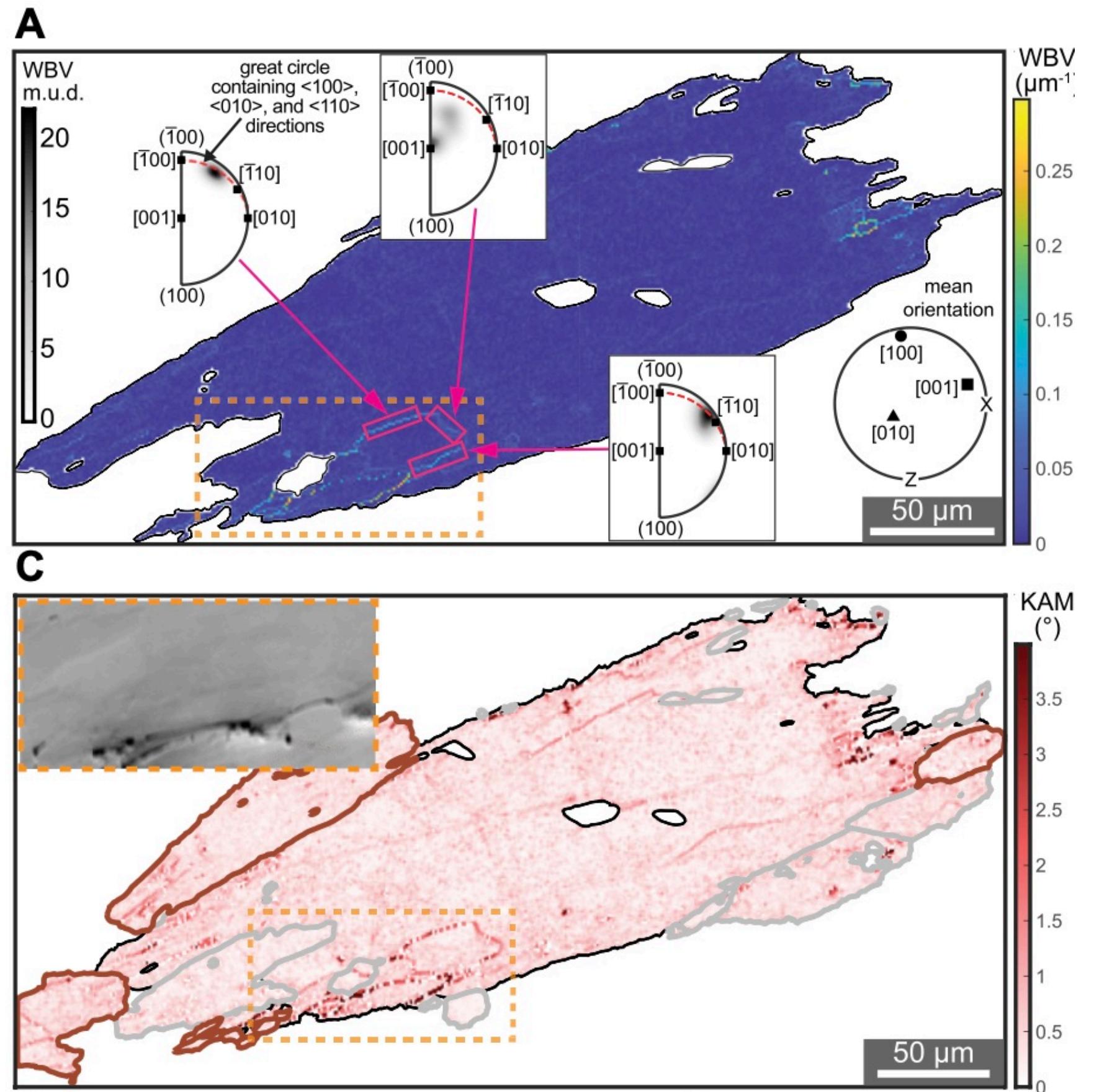
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Weighted Burgers vector analysis shows multiple slip systems active - Dislocation creep

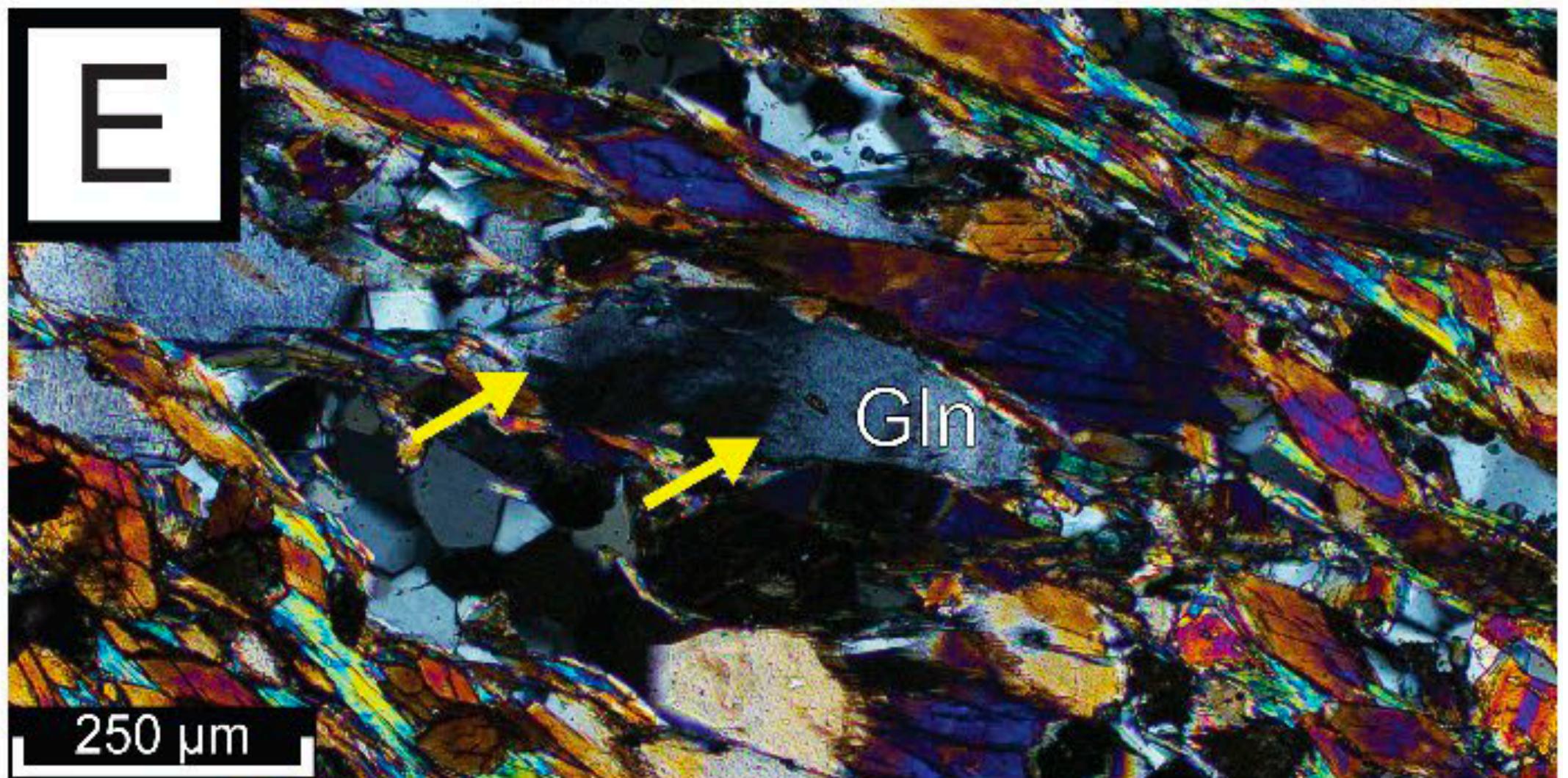
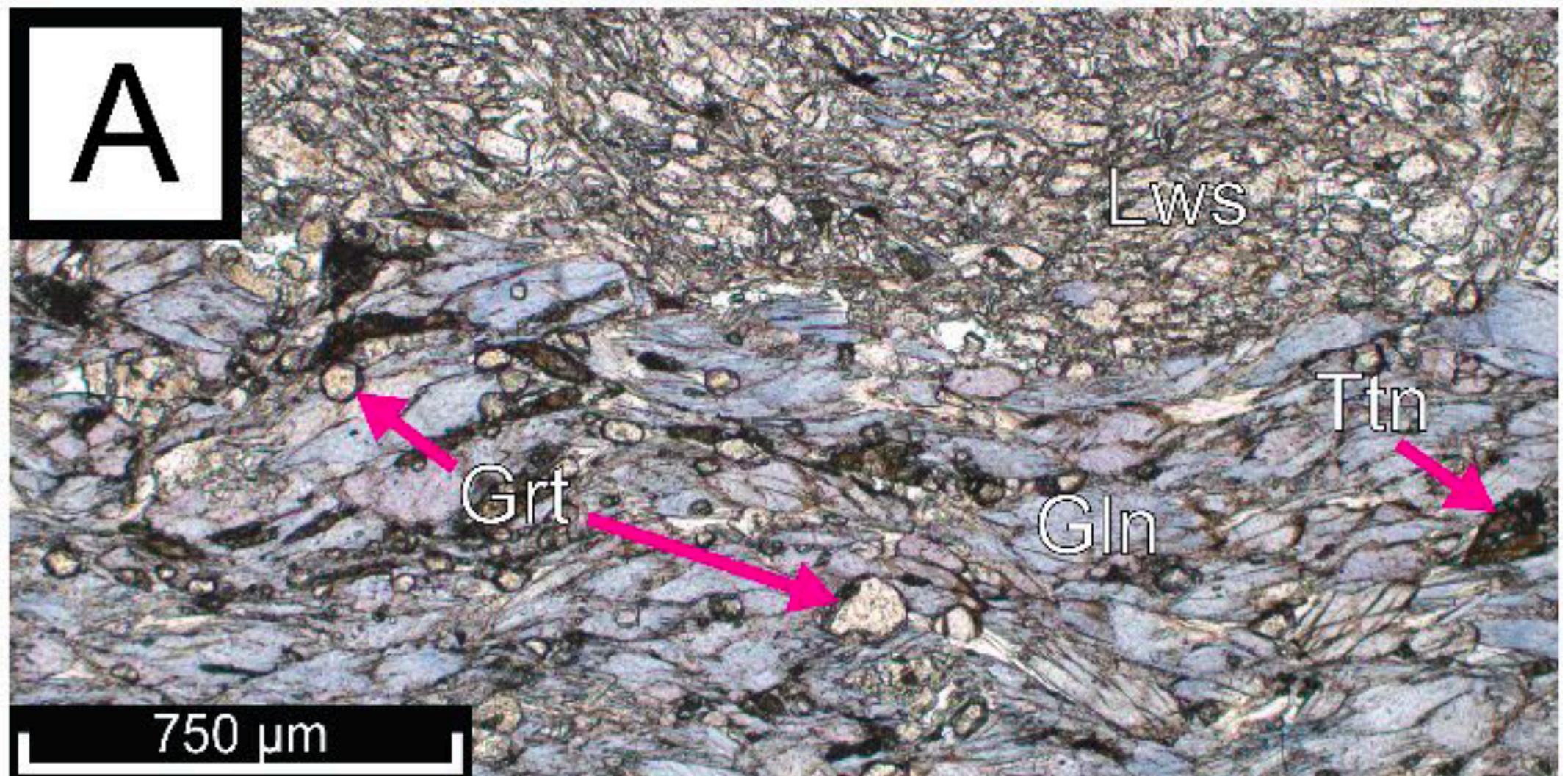


Dislocation to diffusive mechanism shift in glaucophane

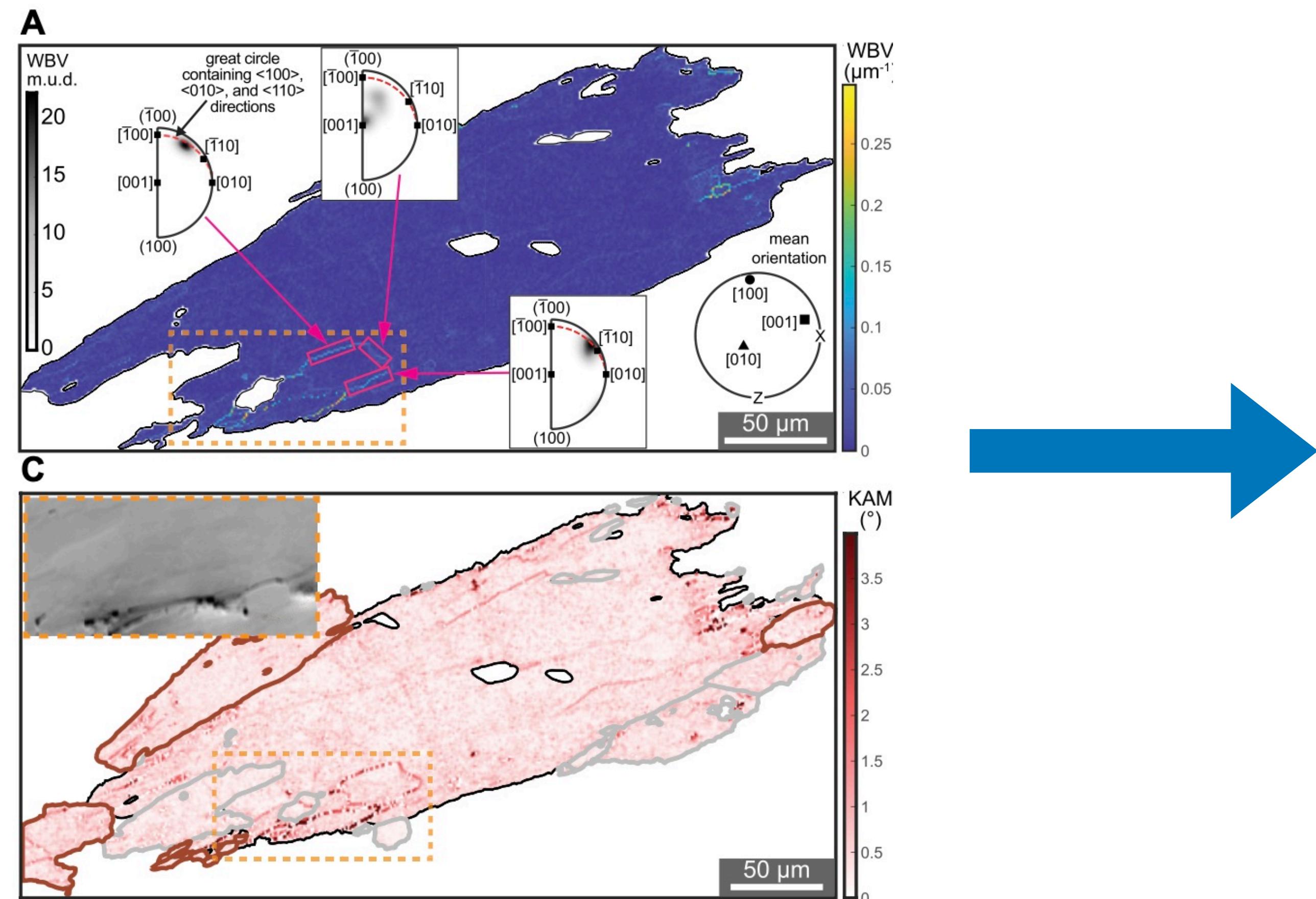


Weighted Burgers vector analysis shows multiple slip systems active - Dislocation creep

Griggs type experiments reveal $n = 5$; flow law in prep.
Work of Jason Ott (Ott, Condit, & Pec, *in prep*)

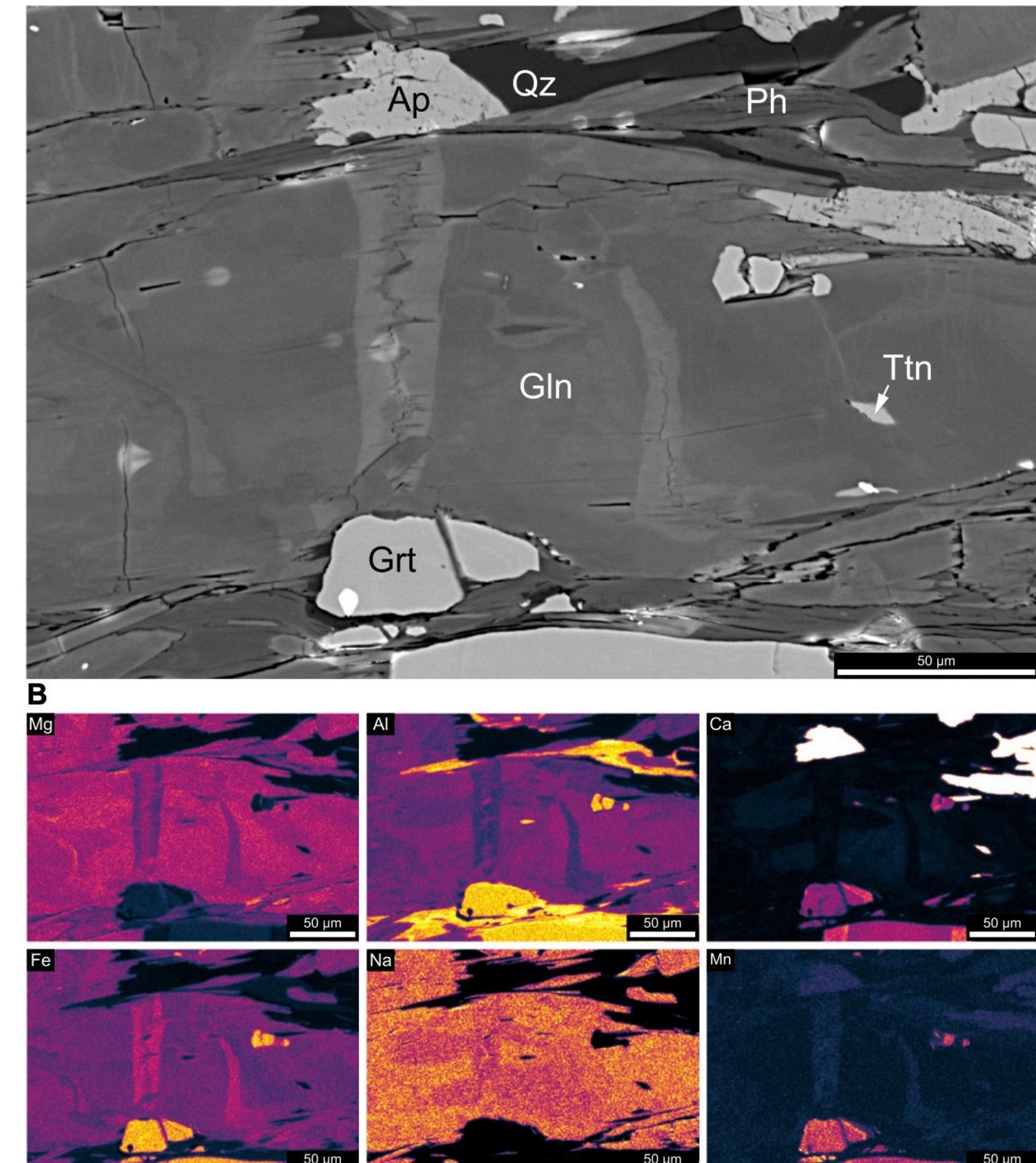


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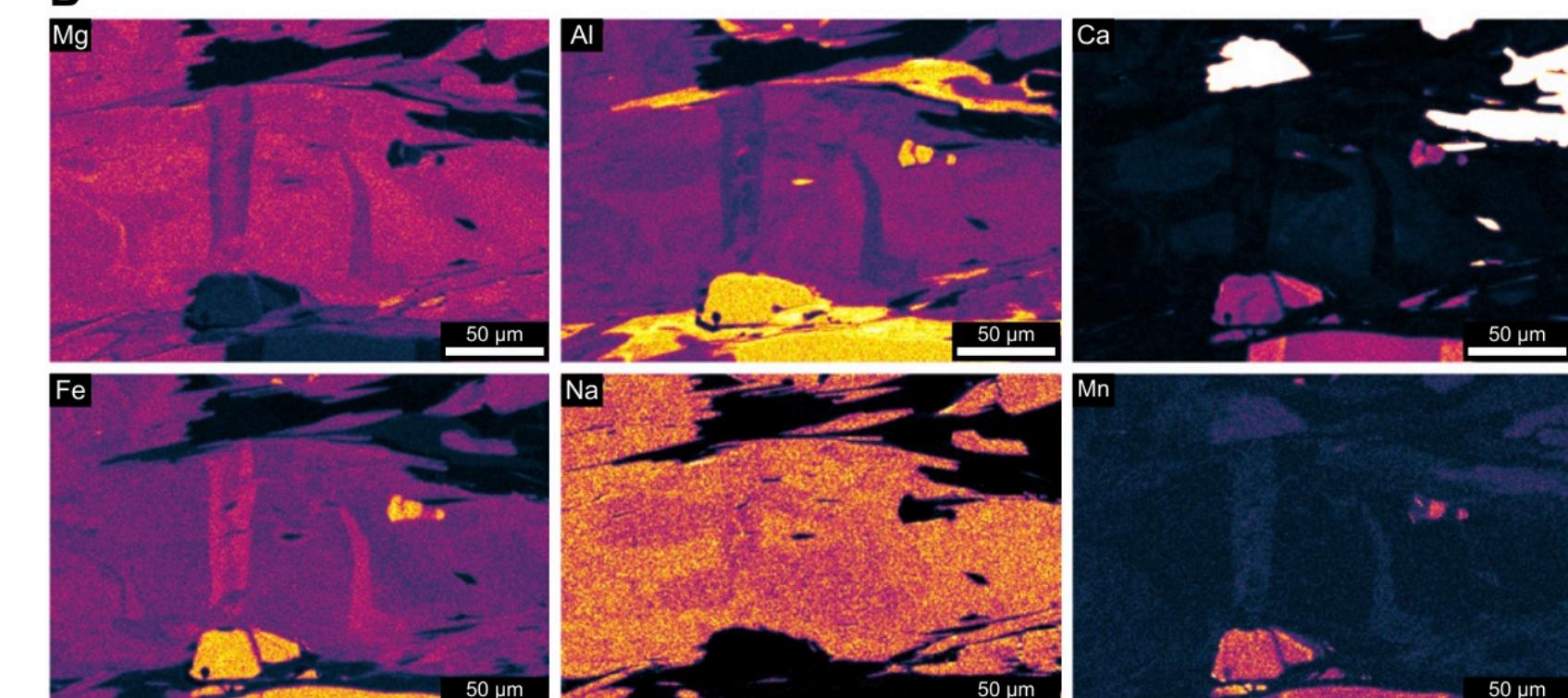
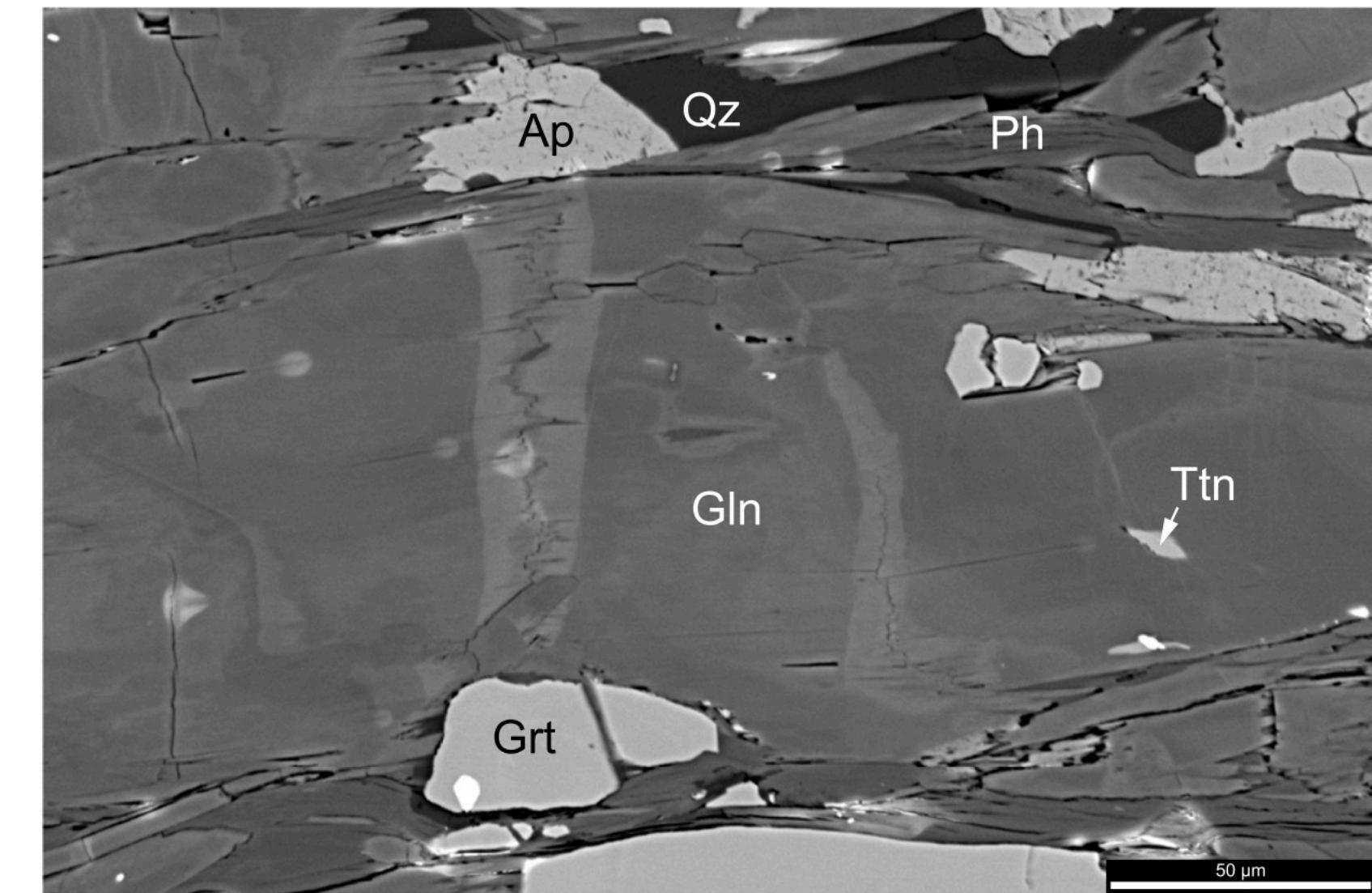
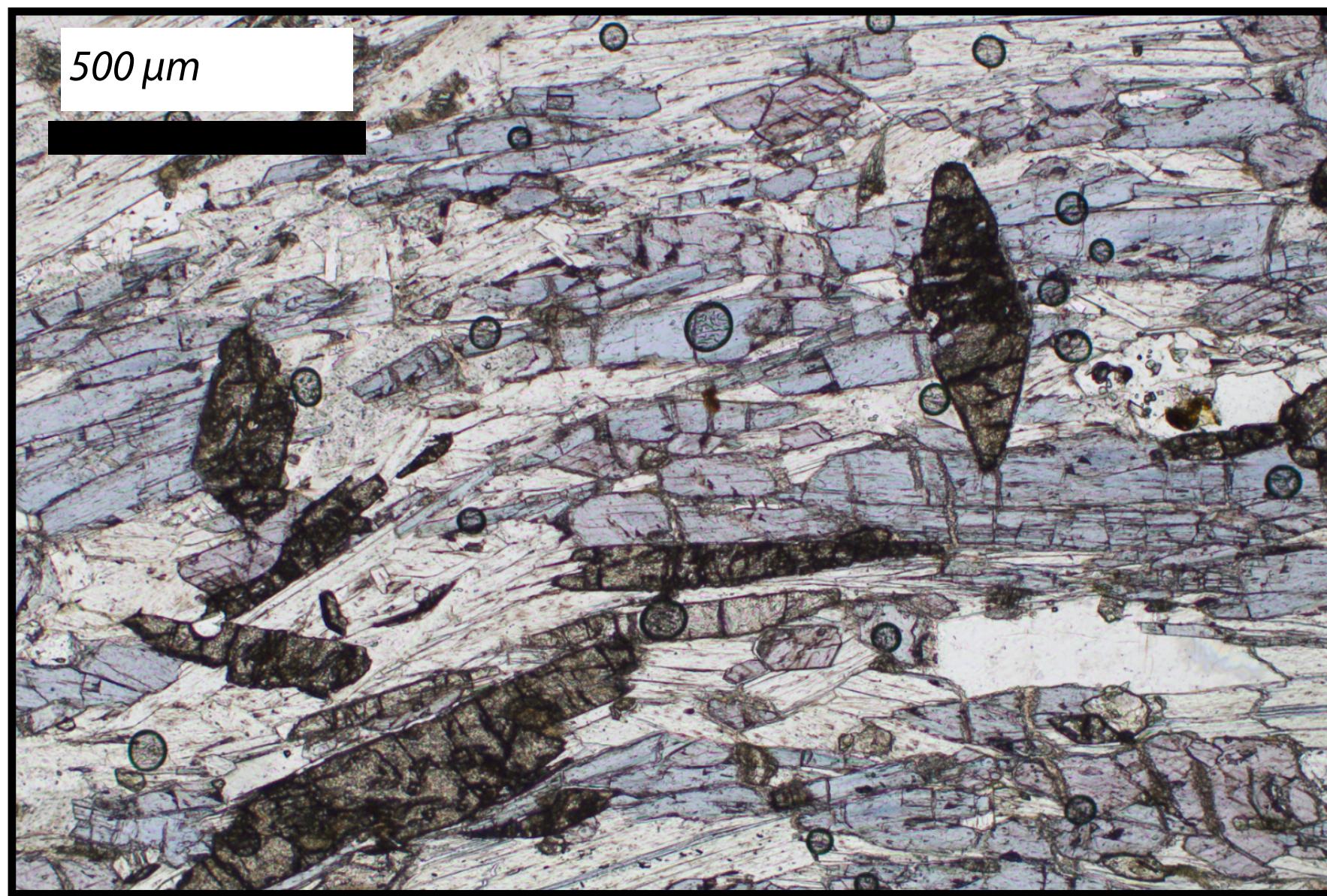
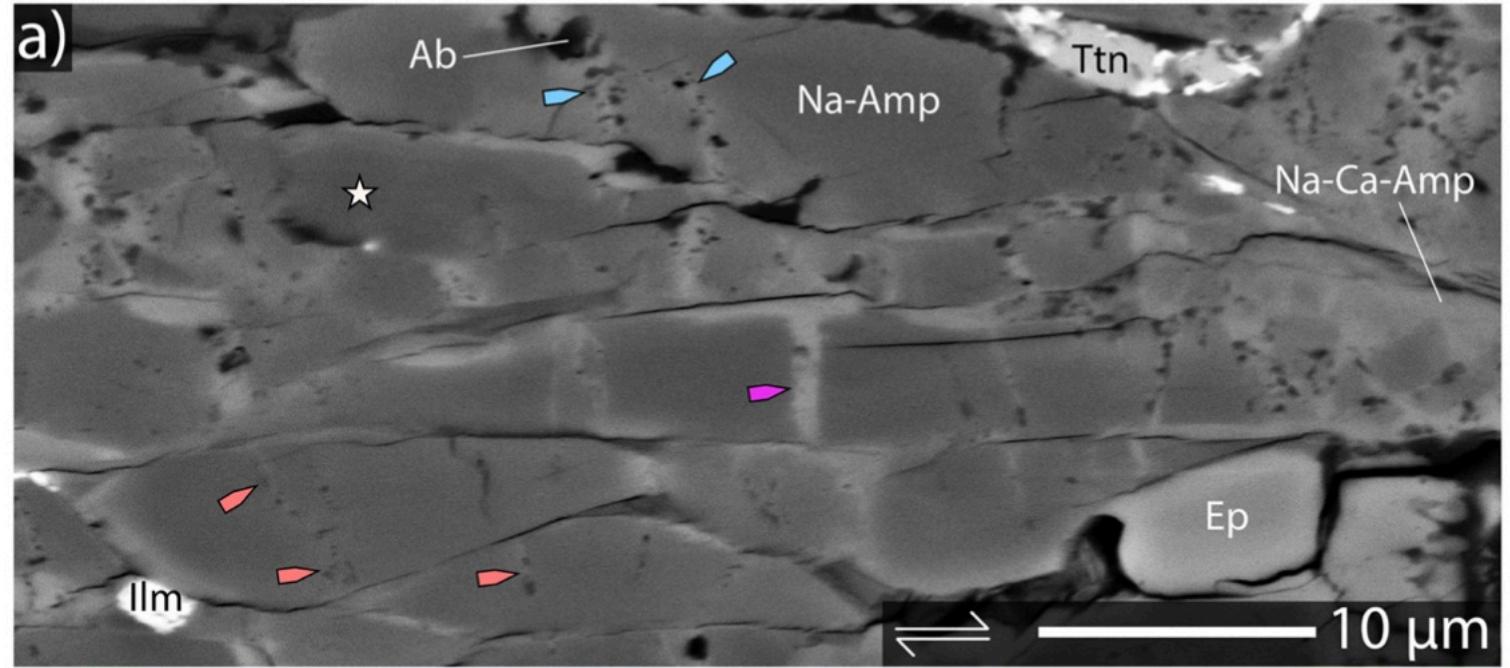
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Later microboudinage; new amphibole growth with fluid infiltration and chemical potential gradients - reactive fluids and/or shift in P-T conditions

Dislocation to diffusive mechanism shift in glaucophane

Seen experimentally by Tokle et al., 2023; n = 2



Later microboudinage; new amphibole growth with fluid infiltration and chemical potential gradients - reactive fluids and/or shift in P-T conditions

Seen during early exhumation on Sifnos, Island, GR
(~1.2 GPa; 400°C + H₂O & CO₂ fluid)

Ott, Condit et al., 2025 Geology